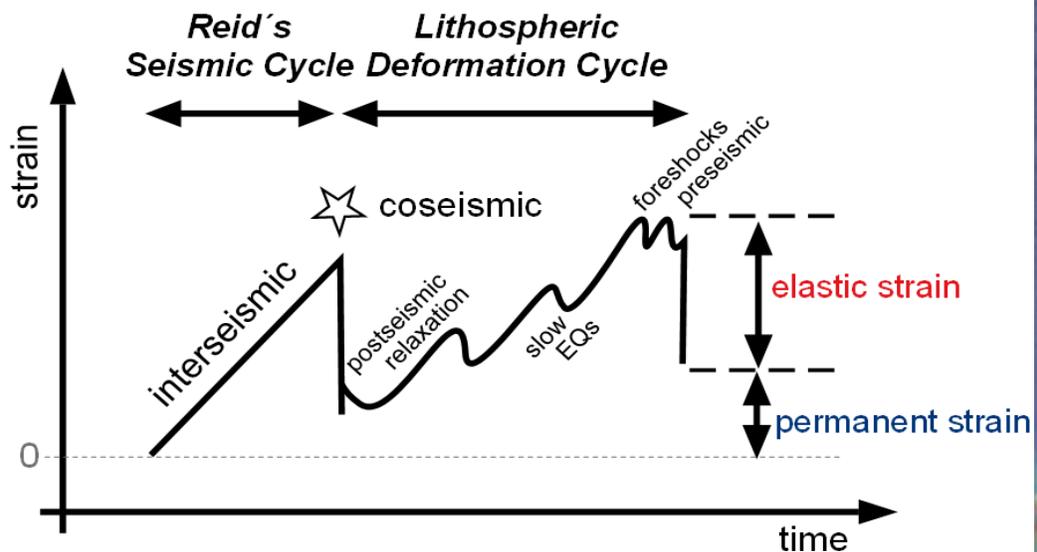
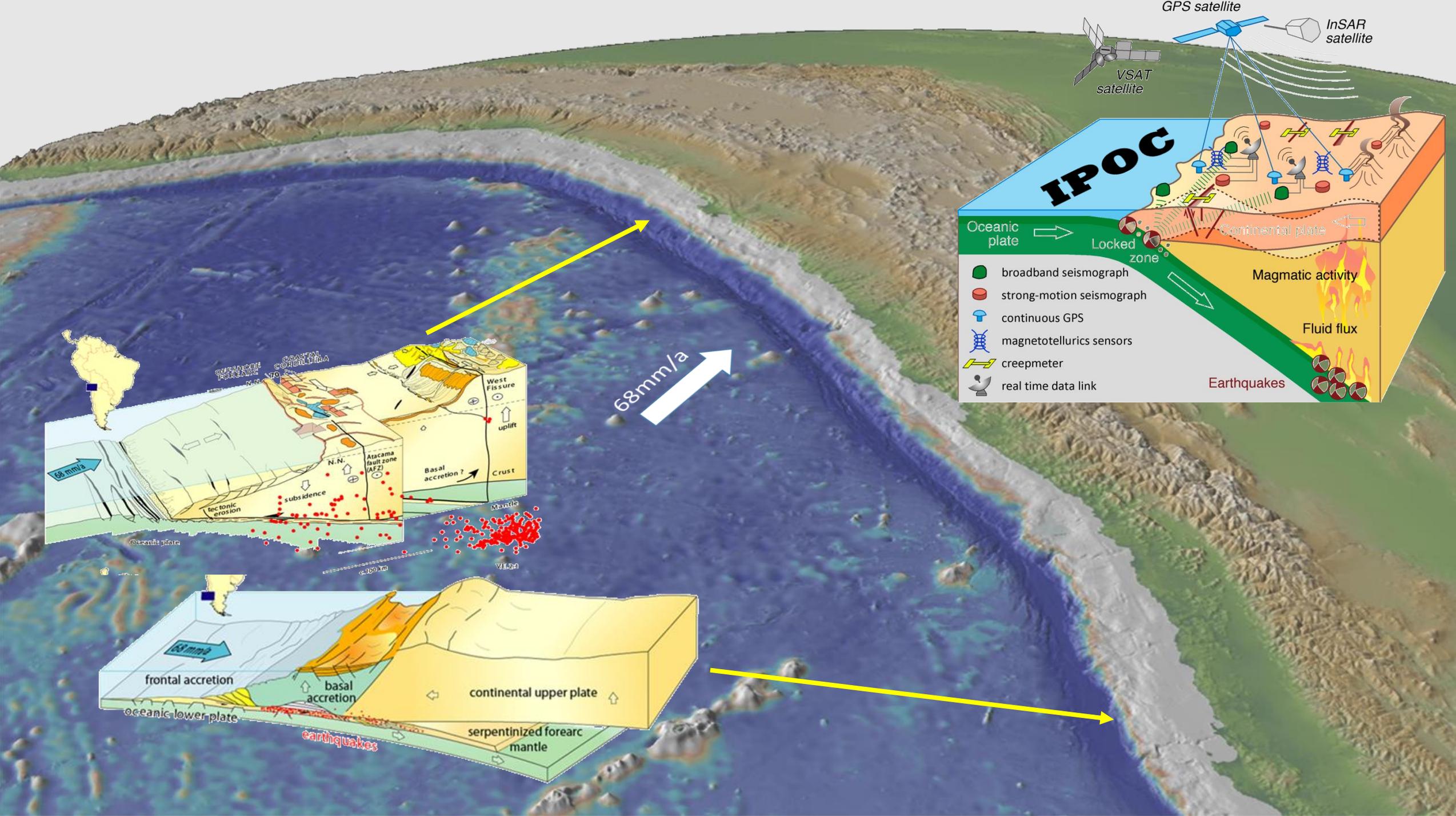


Forearc deformation across time scales - a tale of fluids, locking and transients

Onno Oncken, GFZ Potsdam

with M. Moreno, P. Victor, B. Schurr, I. Ioannidi, J. Bedford, V. Mouslopoulou, I. Urrutia, M. Rosenau, F. Hoffmann, S. Angiboust, S. Li ... and many more

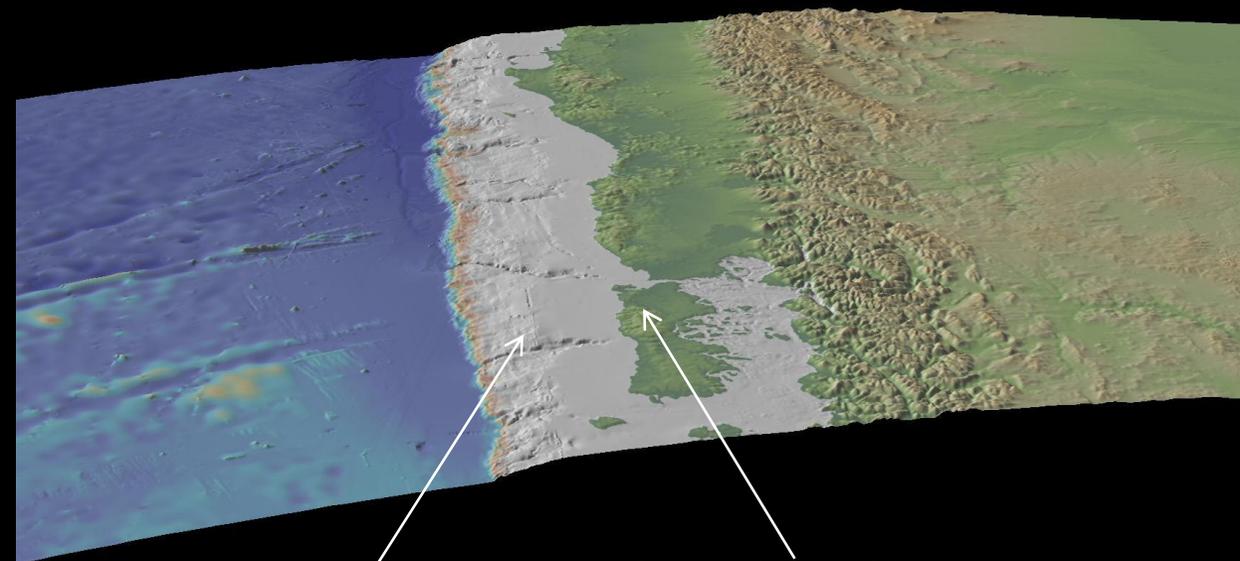
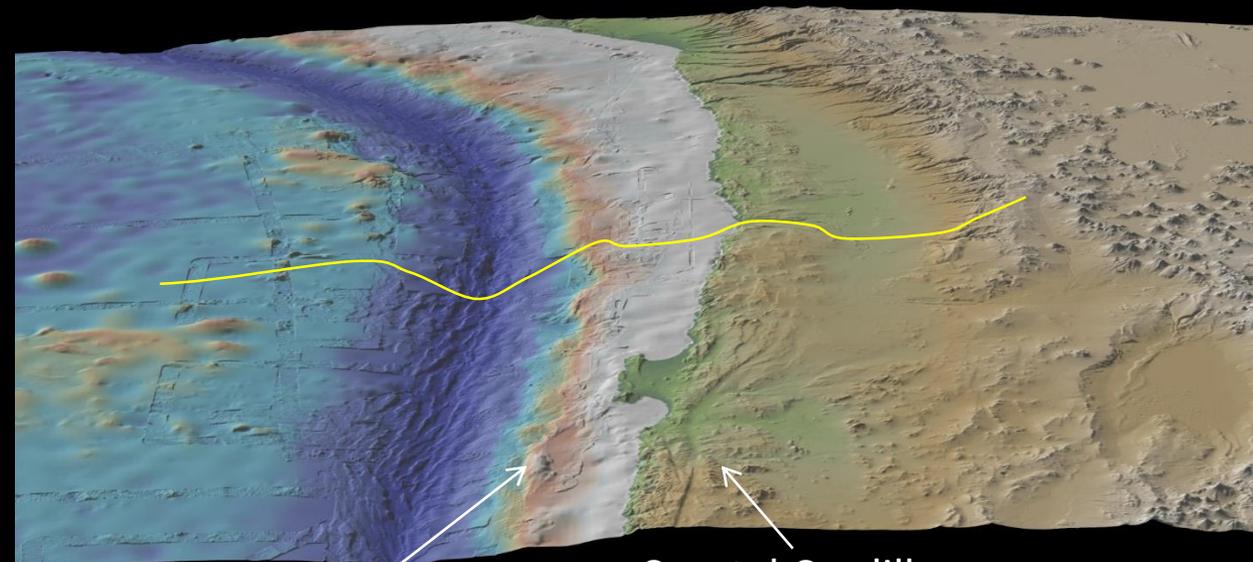




Topographic architecture of the Chile plate margin

tectonically erosive north

accretionary south



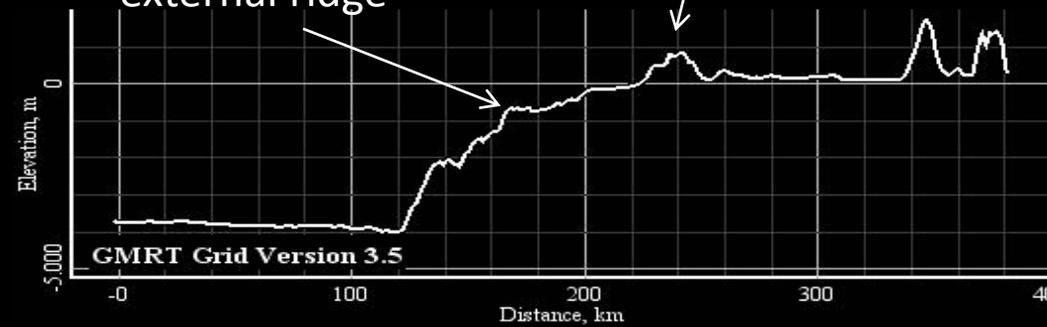
external ridge

Coastal Cordillera

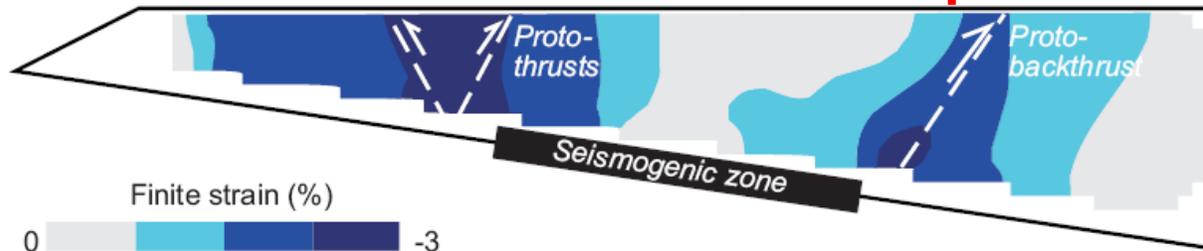
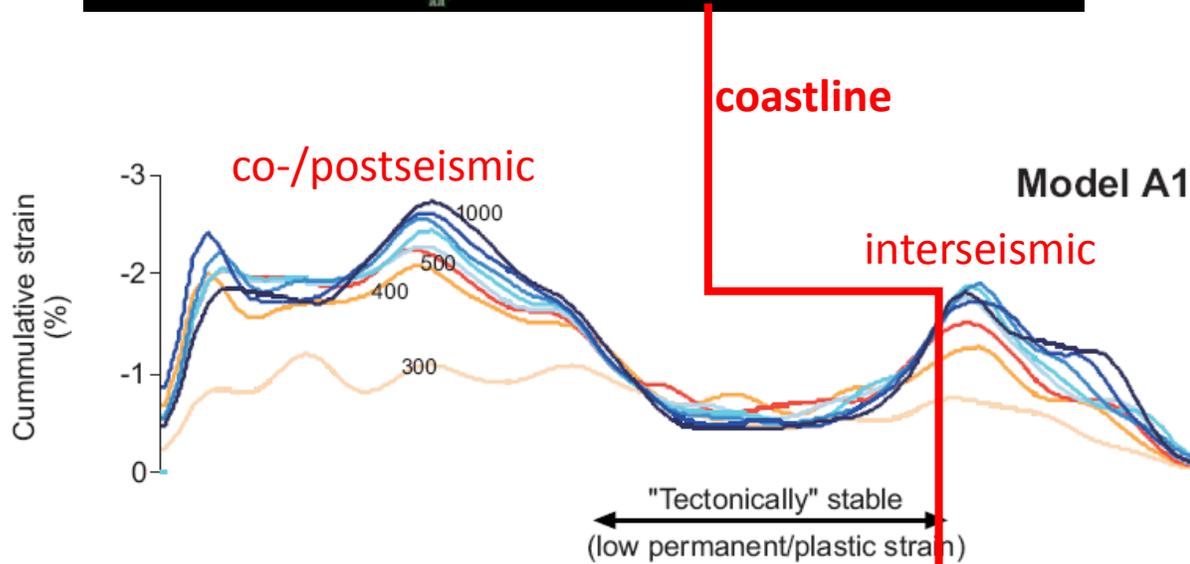
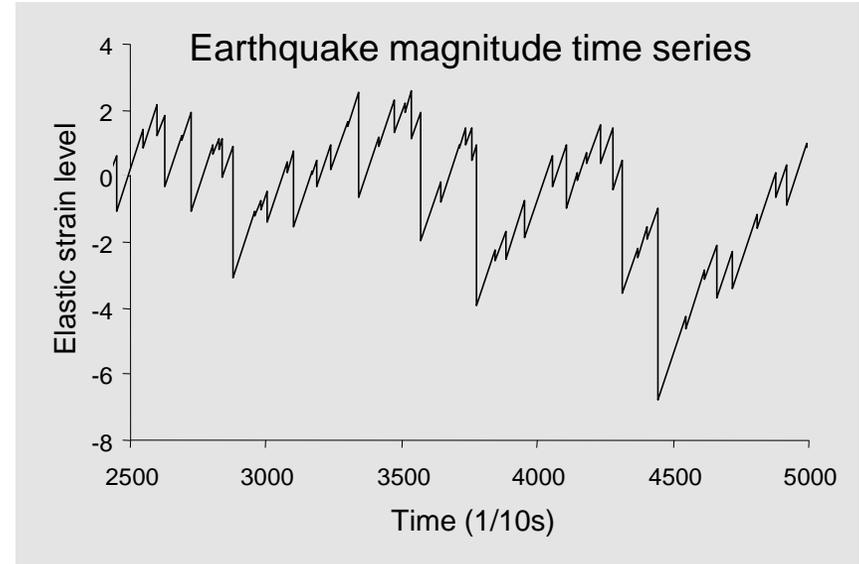
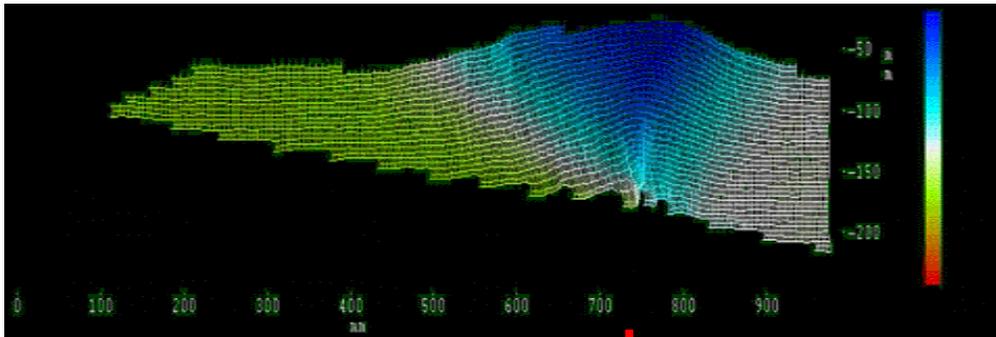


external ridge

Coastal Cordillera



Experimental seismotectonic deformation (10 – 10⁵ years)



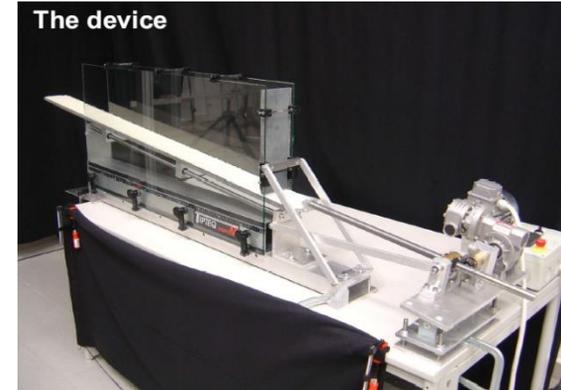
Simulation at large scale
(10² km, 10³ years)

Monitoring at high resolution
(decimeter, minutes)

Nature scale

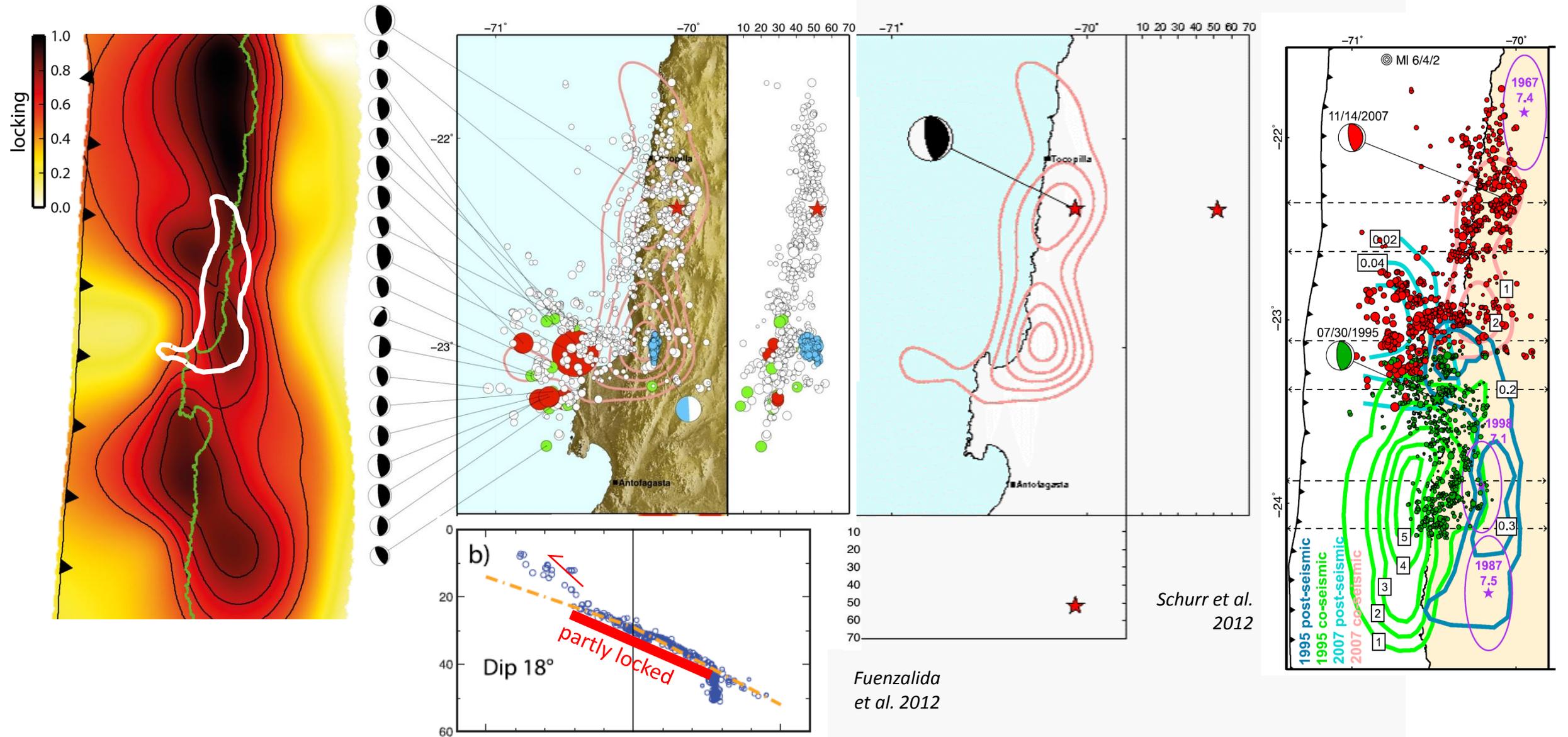
Dynamically scaled for:

- Strength
- Elasticity
- Viscosity
- Inertia
- Gravitation

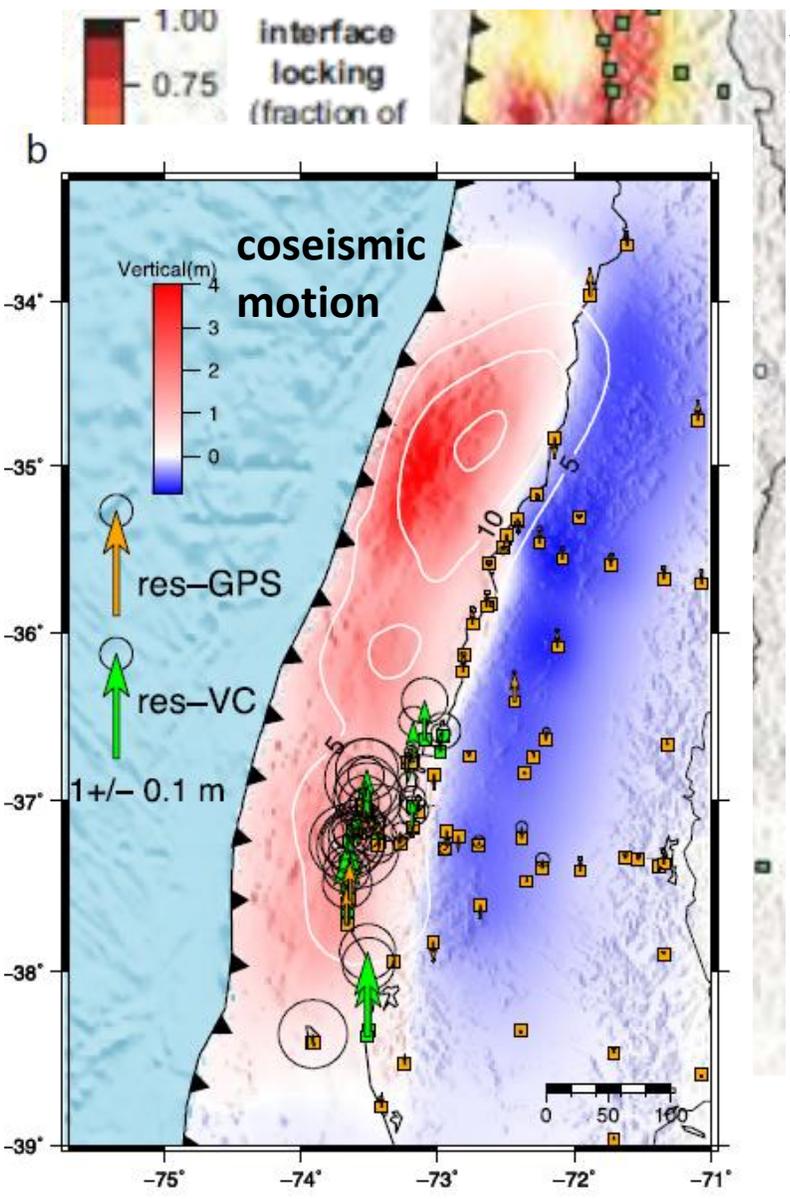


The Mw = 7.8, Tocopilla earthquake, 2007

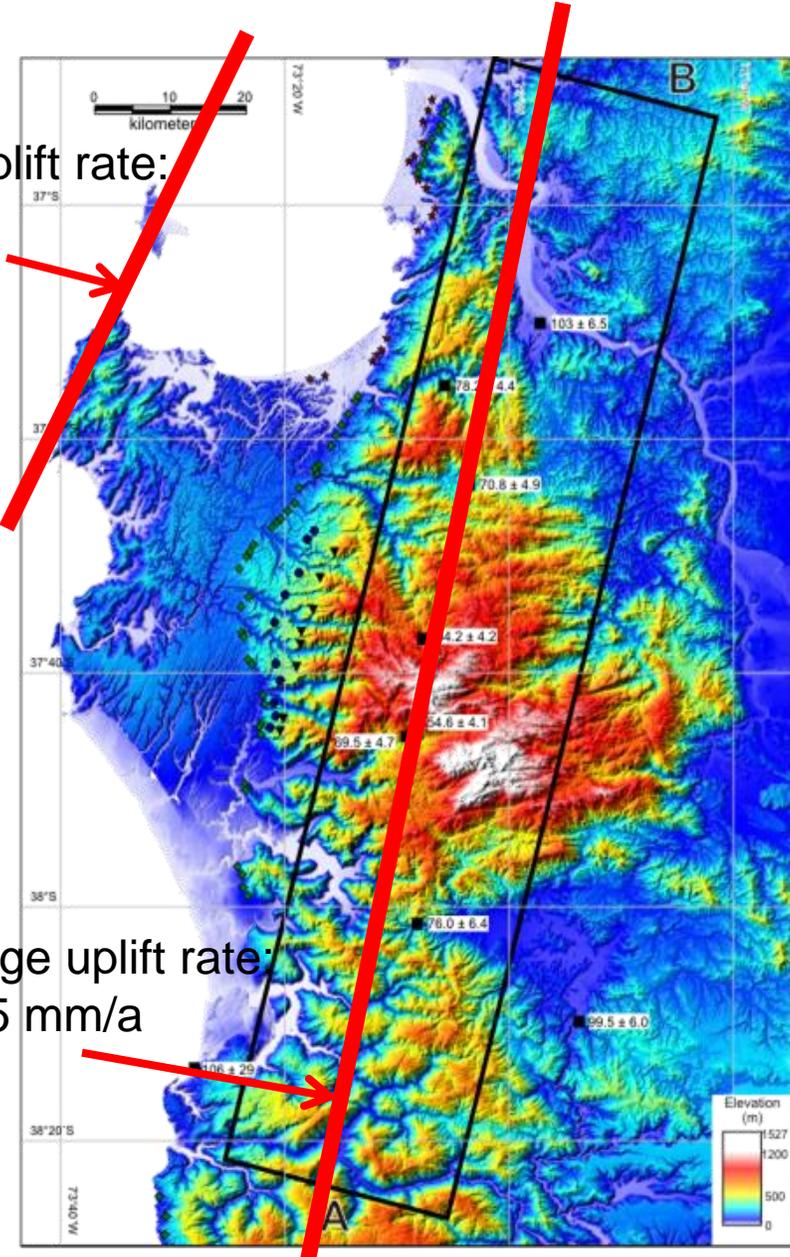
- aftershocks at updip end of locking drive uplift



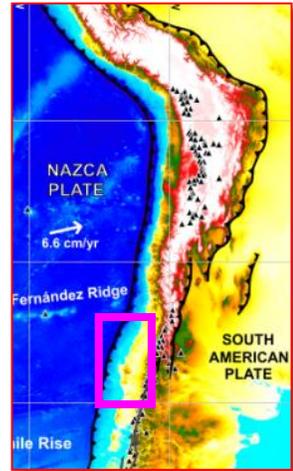
Arauco peninsula - Kinematics since c.3 Ma



Average uplift rate:
2.3 mm/a

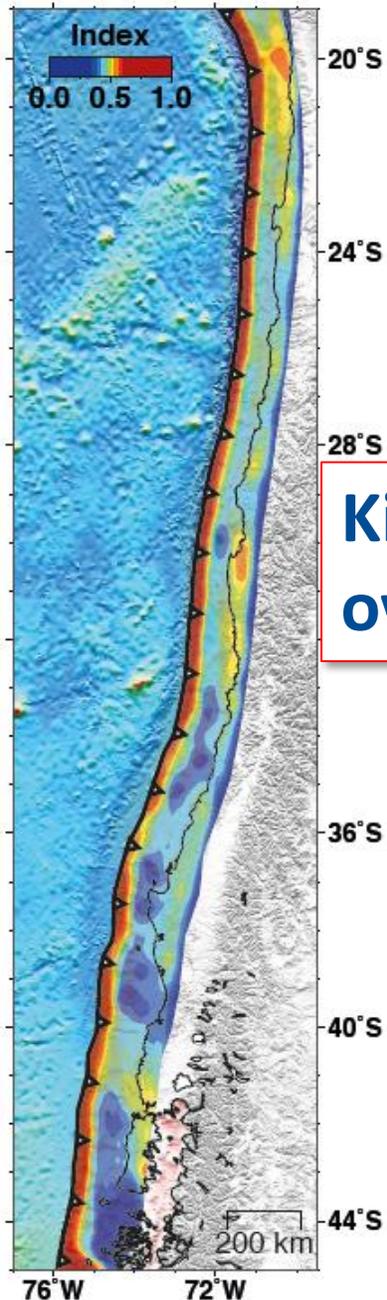


Average uplift rate:
1 - 1.5 mm/a

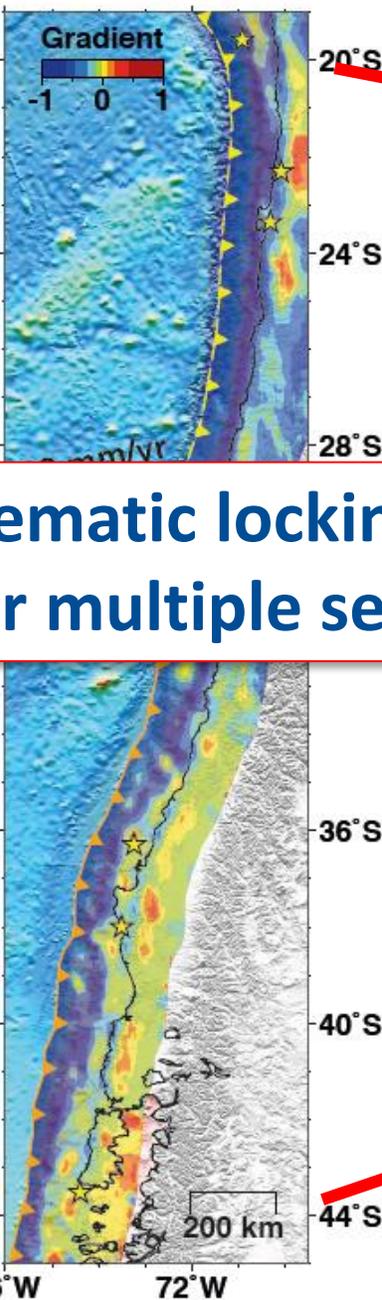


Bookhagen et al., Melnick et al, 2006, 2009

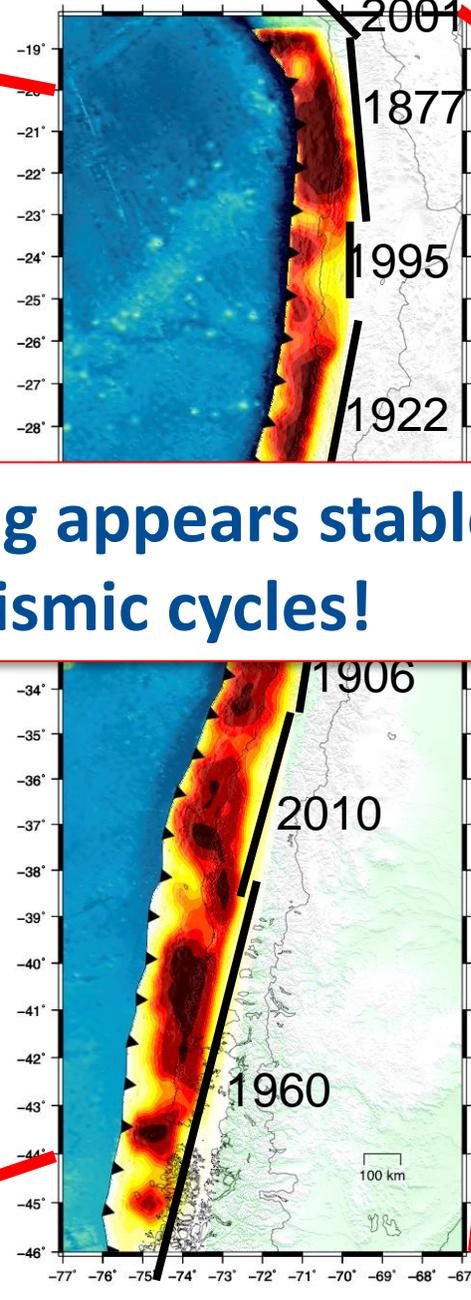
Correlation slope-locking



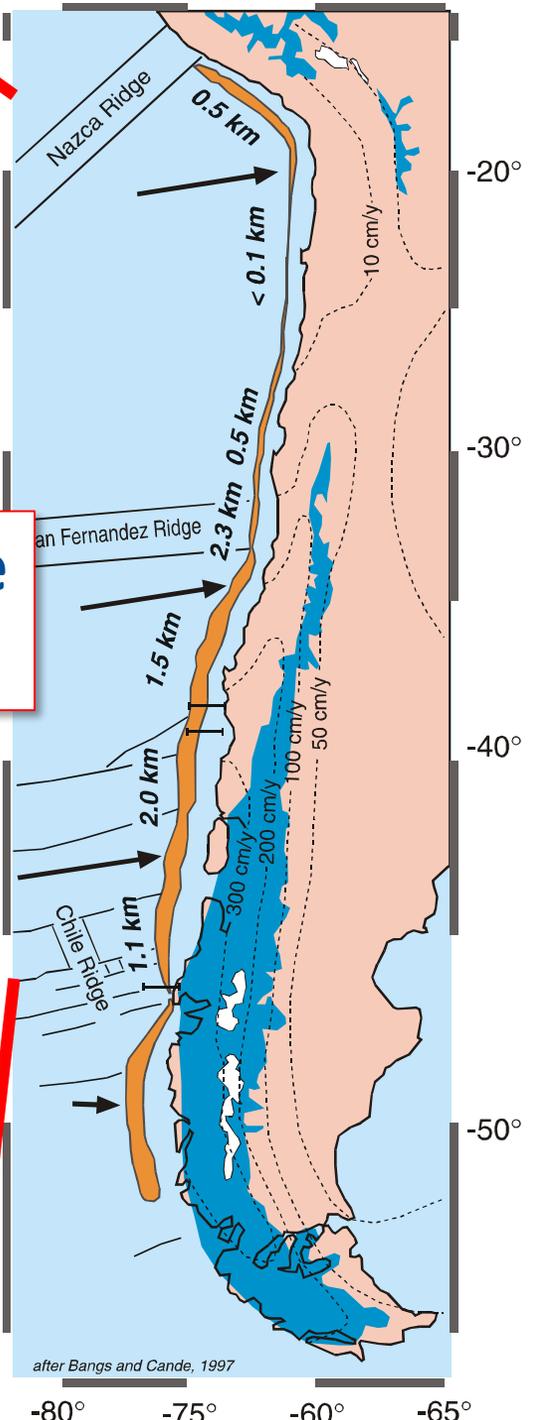
Slope gradient



Locking



Kinematic locking appears stable over multiple seismic cycles!



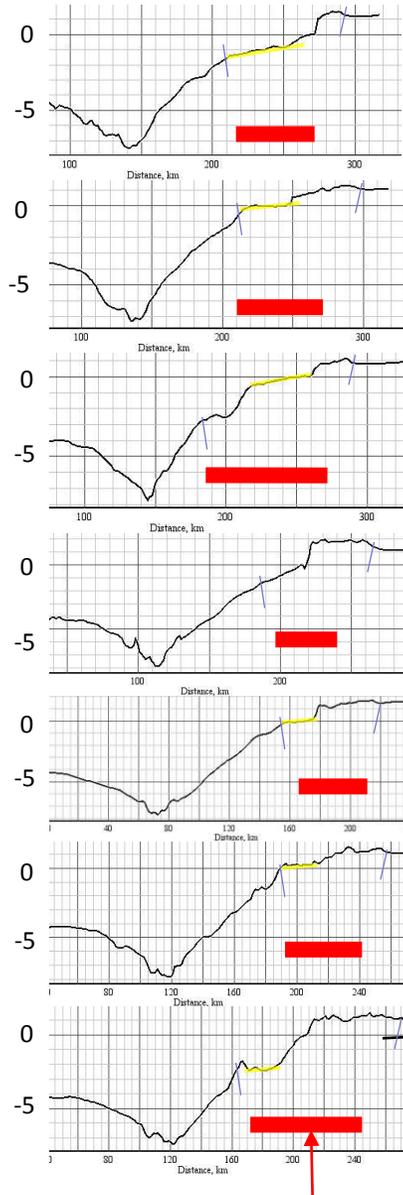
Urrutia et al, in prep.

Locking degree
0.00 0.25 0.50 0.75 1.00

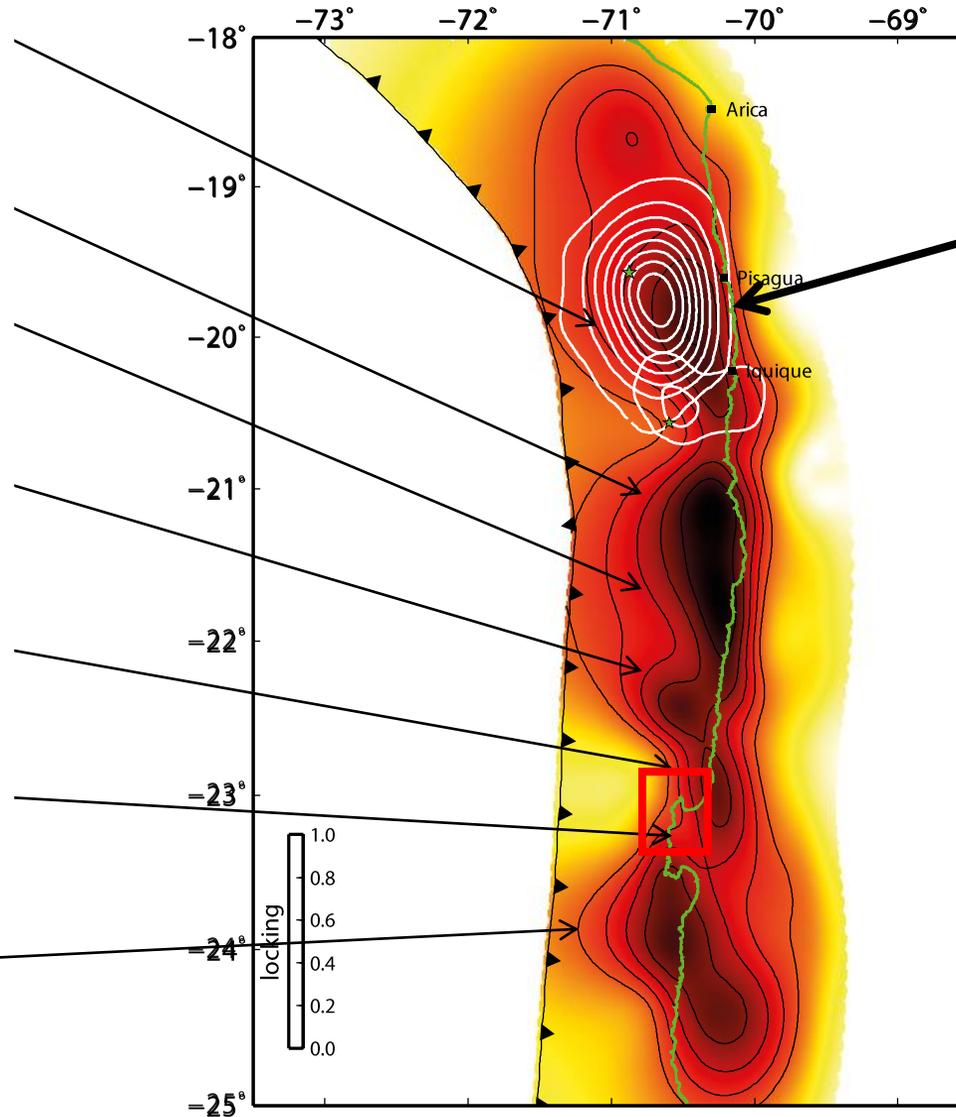
after Bangs and Cande, 1997

Are locked patches stable over 10^5 - 10^7 years?

And through what mechanism?



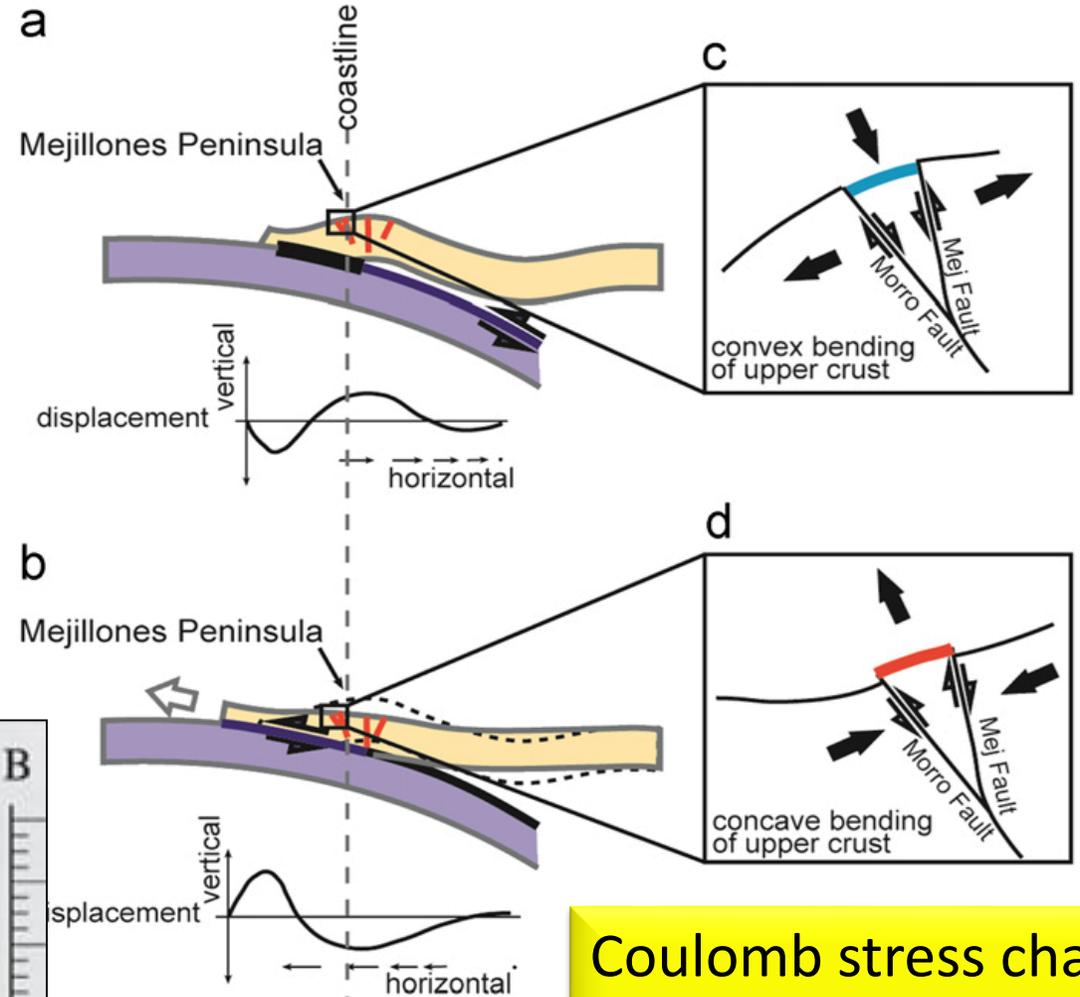
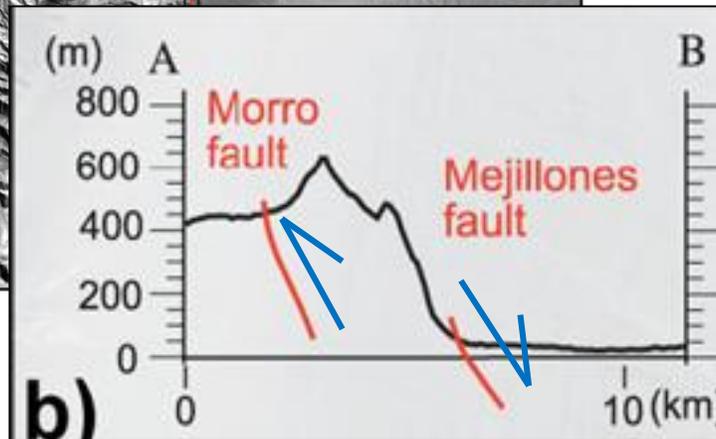
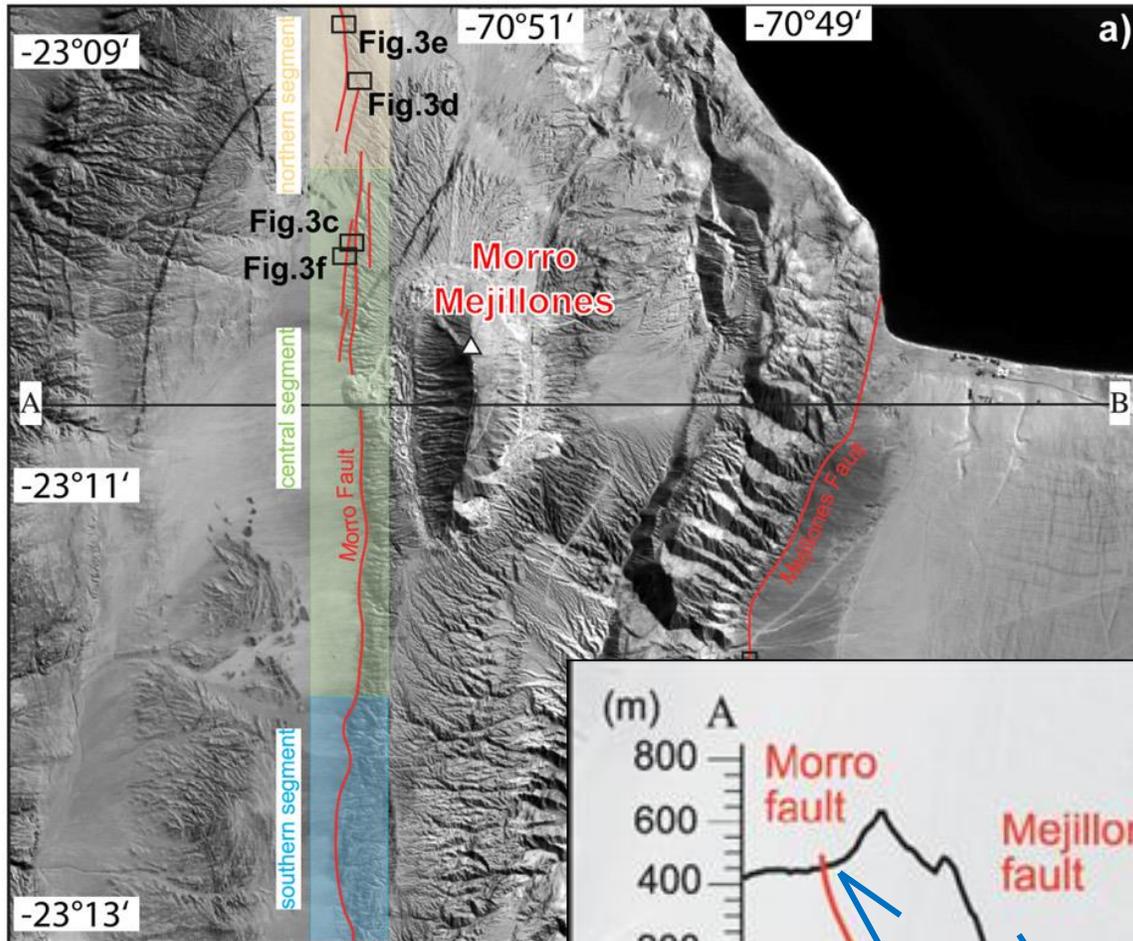
locked



Mw8.1 Iquique earthquake, 4-1-2014

Schurr et al., Nature, 2014

Slip reversal on upper plate faults during megathrust seismic cycle ?



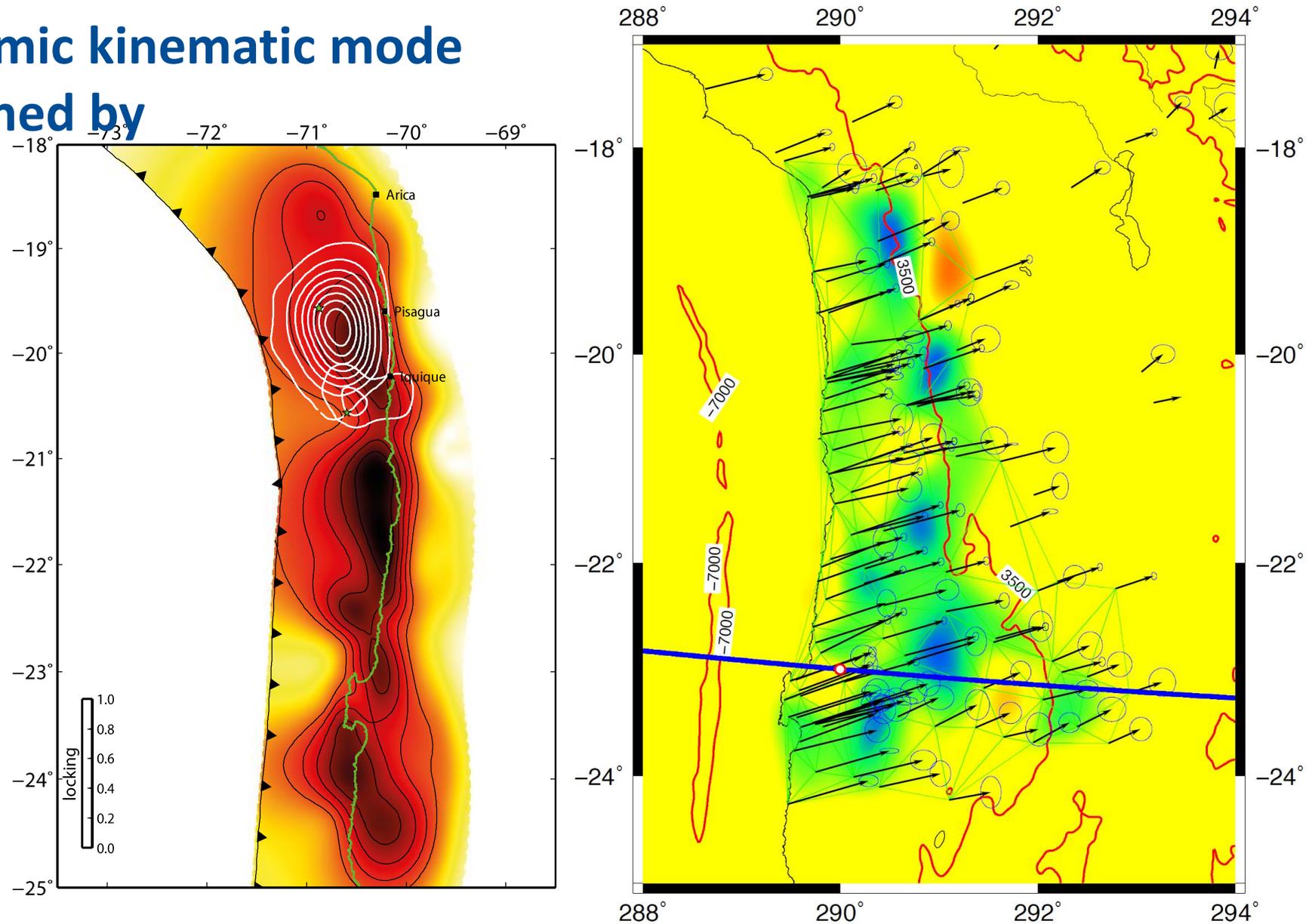
Coulomb stress change of several 10 kPa only!!



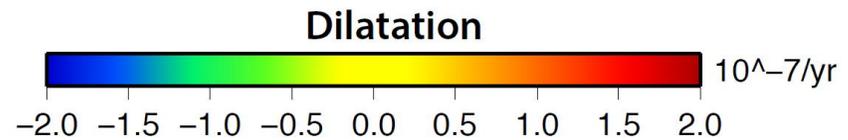
**extension during
postseismic
relaxation and
afterslip,
shortening
interseis-
mically?**



Late interseismic kinematic mode
indeed governed by
compression
landward of
the locked
zone

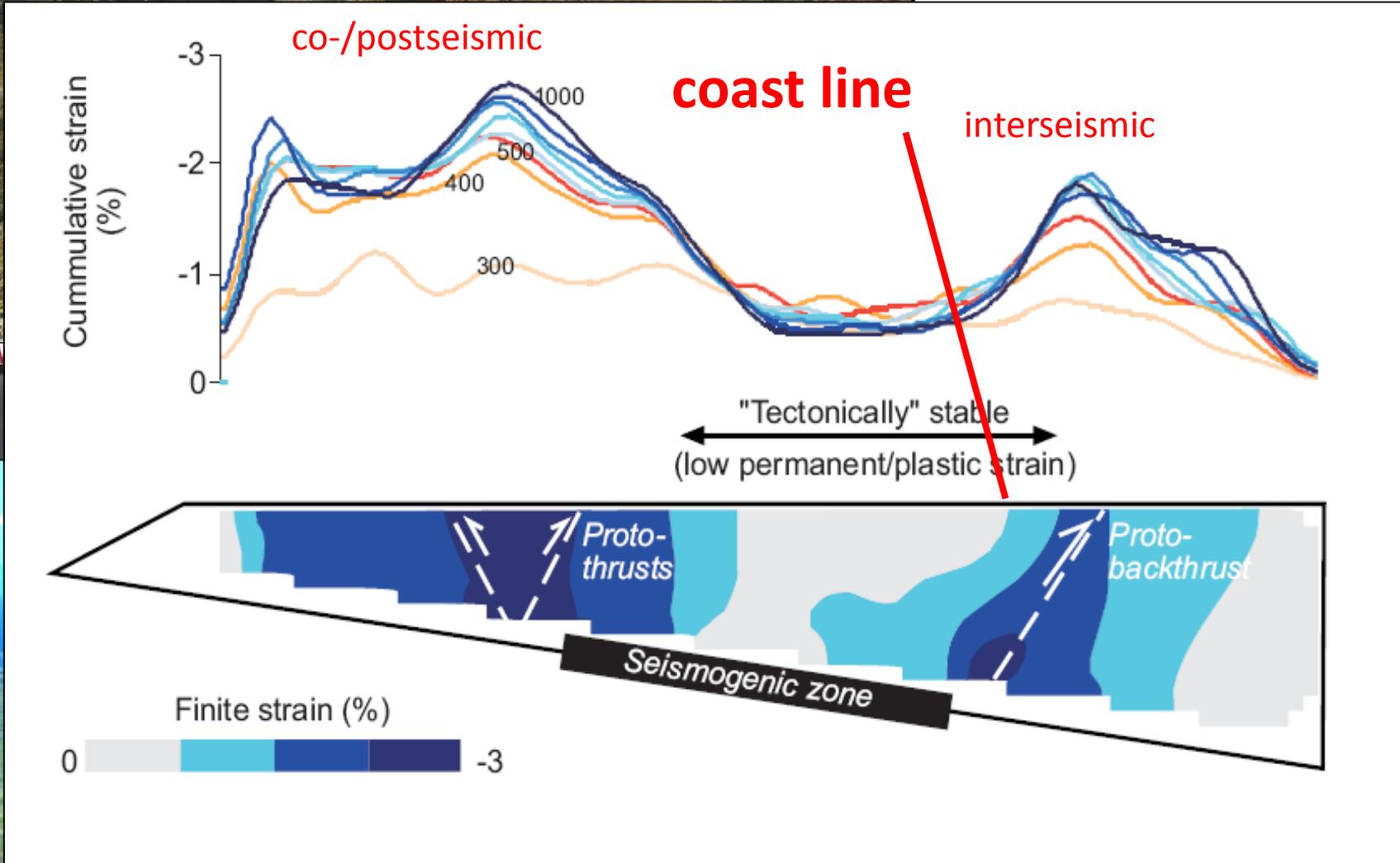
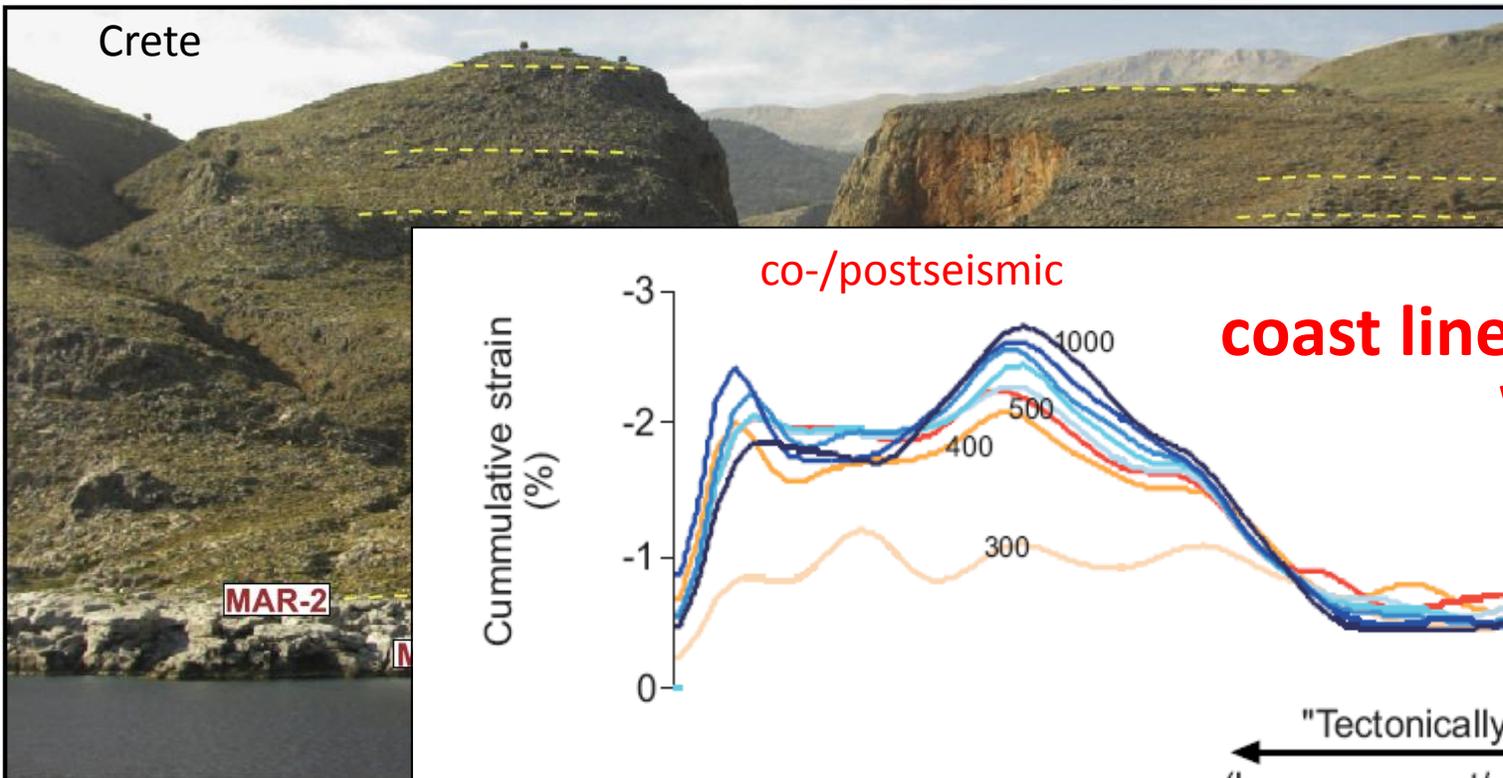


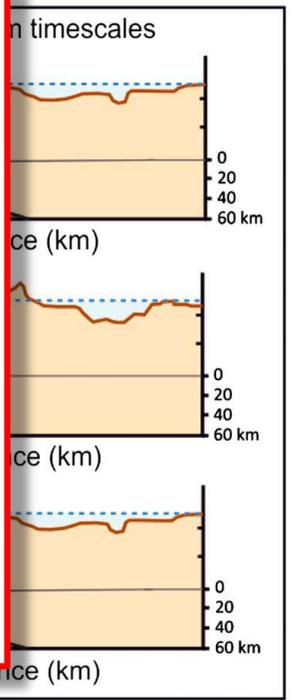
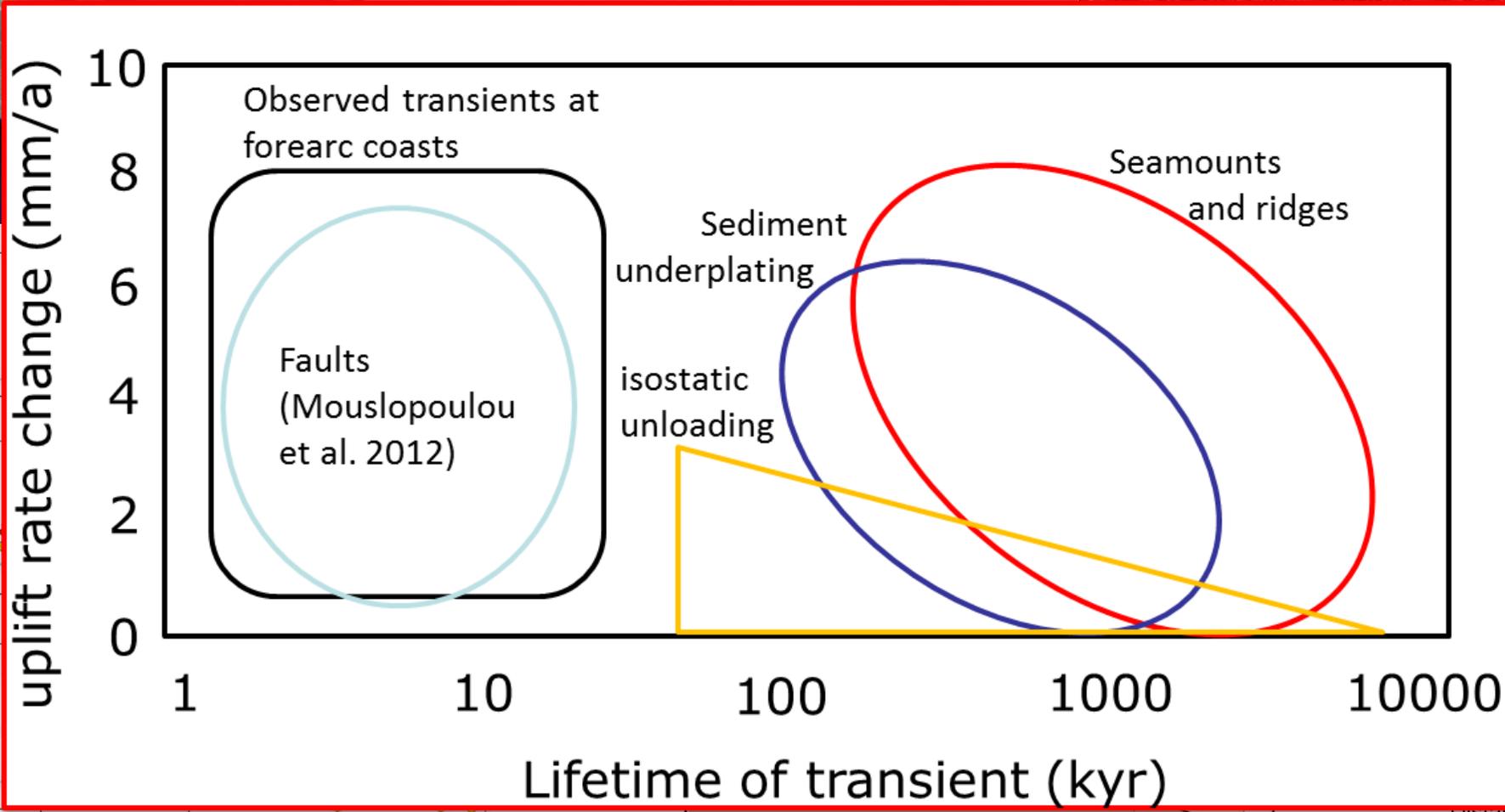
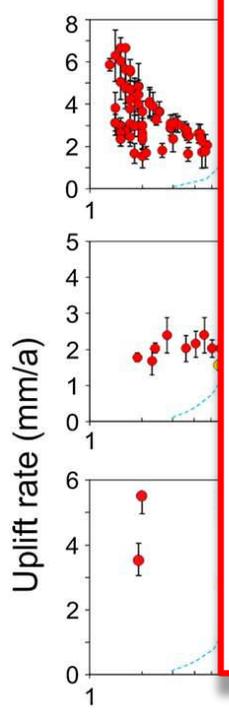
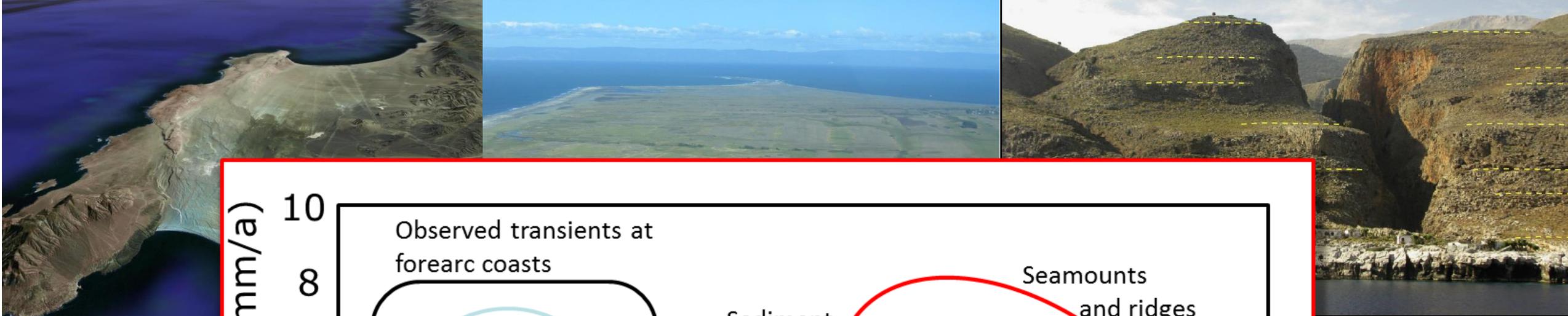
data in Hoffmann et al., 2018;
strain calculation courtesy of S. Lamb



Crete

Strandlines mark sea level – and past earthquakes



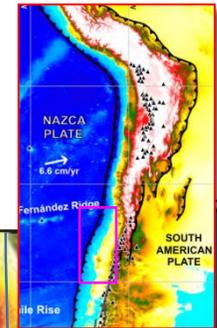


integration time (kyrs)

Mouslopoulou et al., 2015, 2016

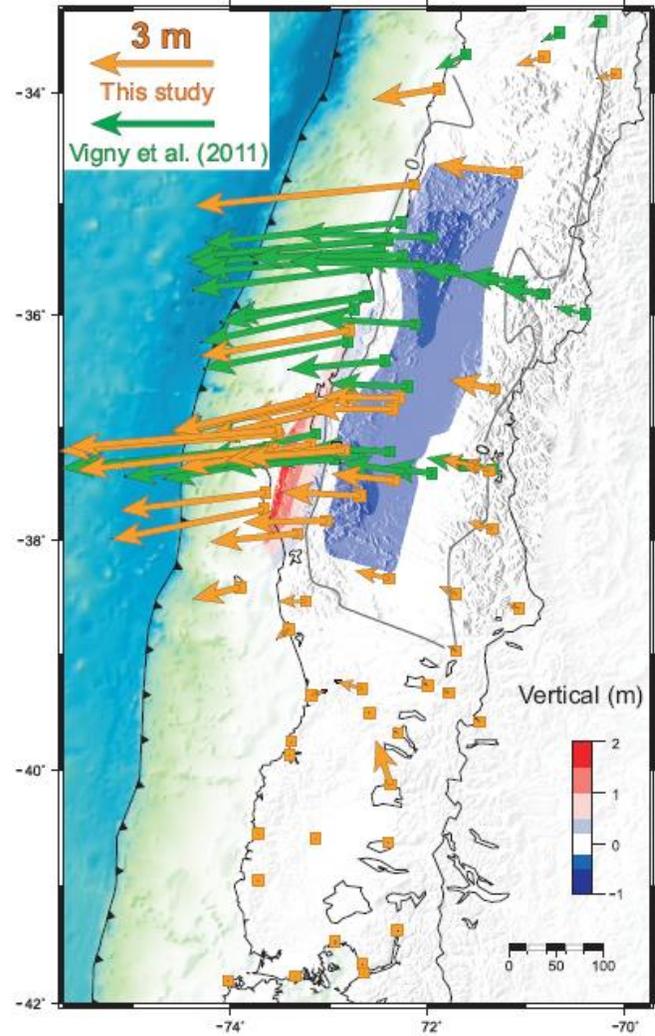
Long-term uplift transients in forearcs from earthquake clustering

Preseismic locking and coseismic slip – the 8.8. Maule earthquake of 27.2.2010



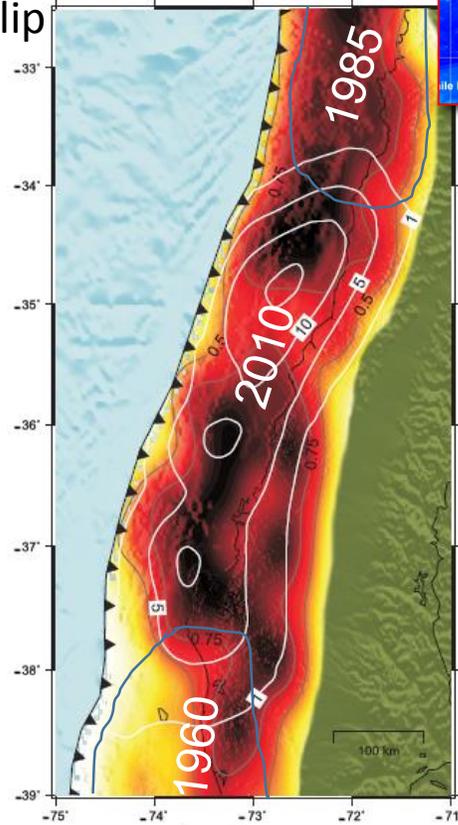
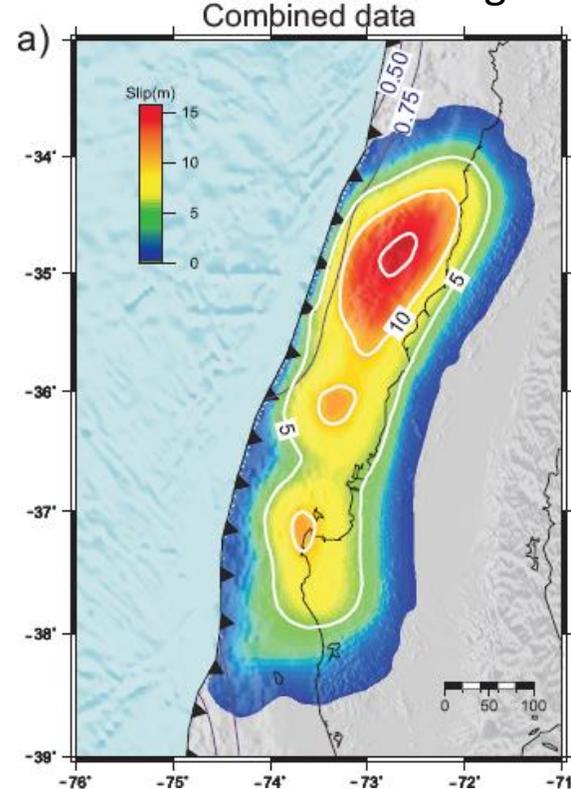
preseismic displacement

coseismic displacement

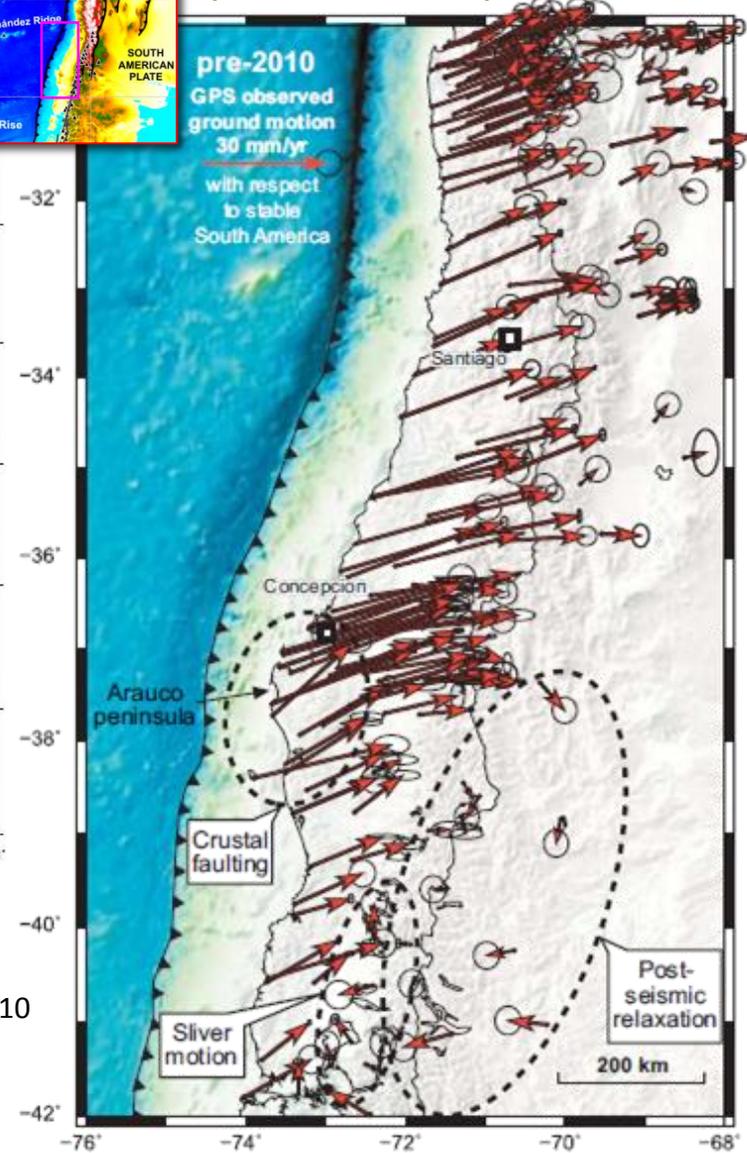
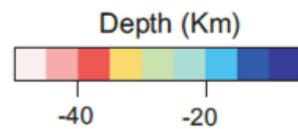
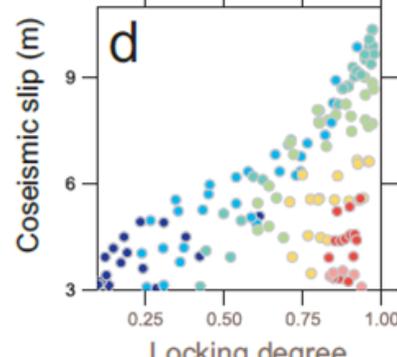


(Moreno et al. EPSL 2012)

locking and slip



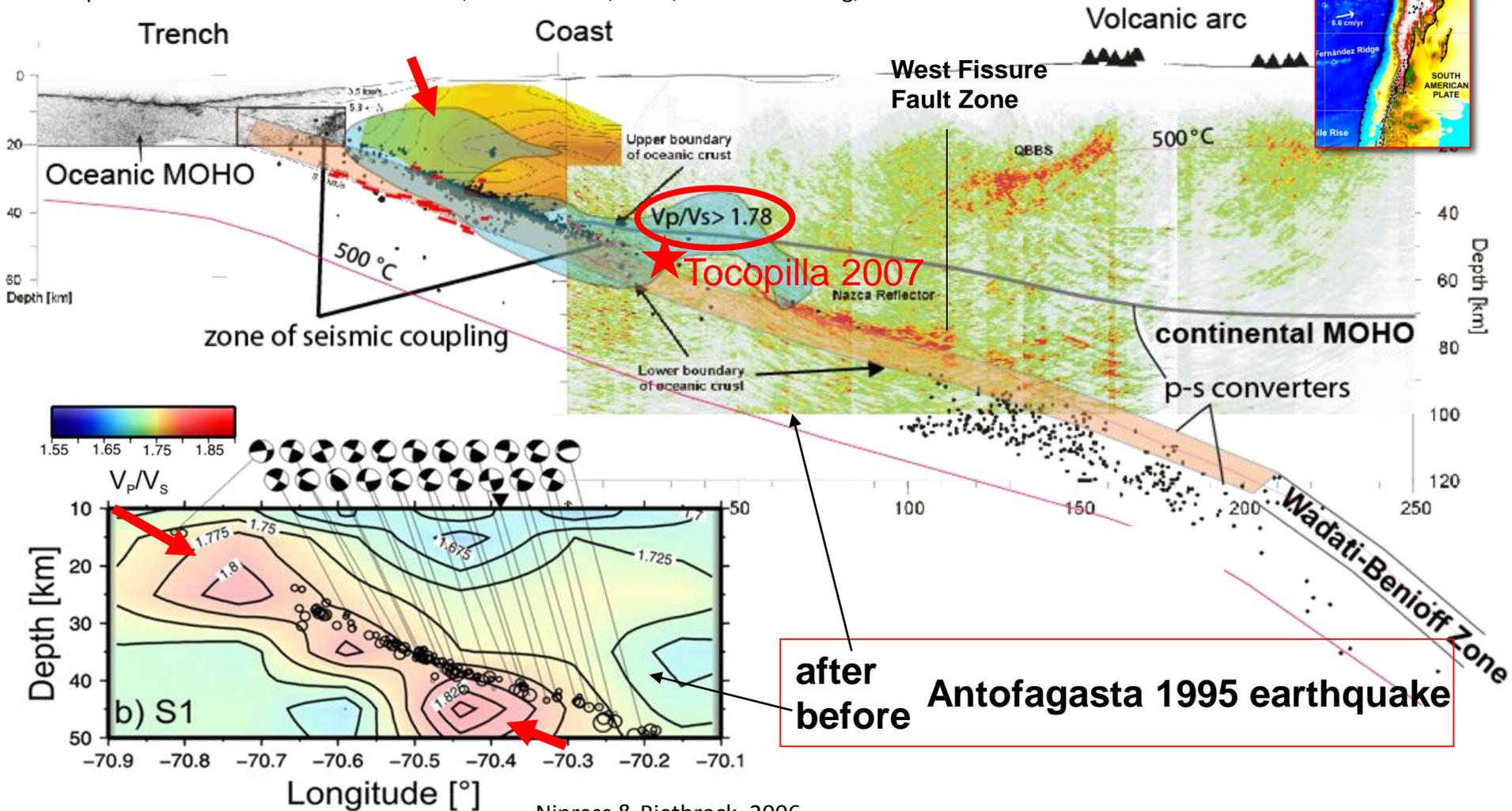
Moreno, Rosenau & Oncken, Nature 2010



Seismological image reveals hydraulic transients

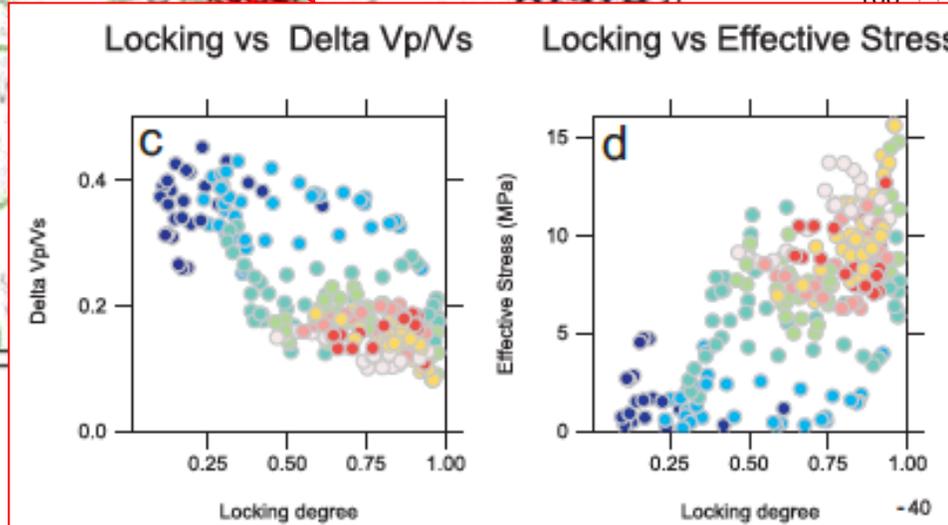
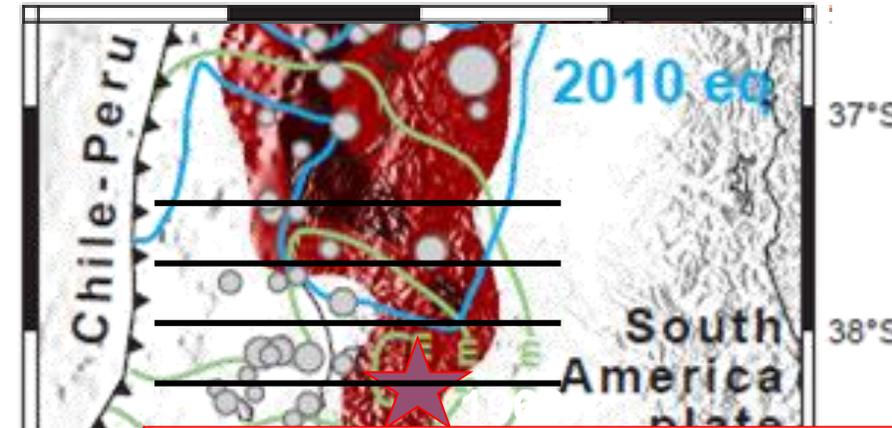
- the North Chilean case

Composite section of ANCORP and CINCA; Oncken et al., 2003; Husen & Kissling, 2001

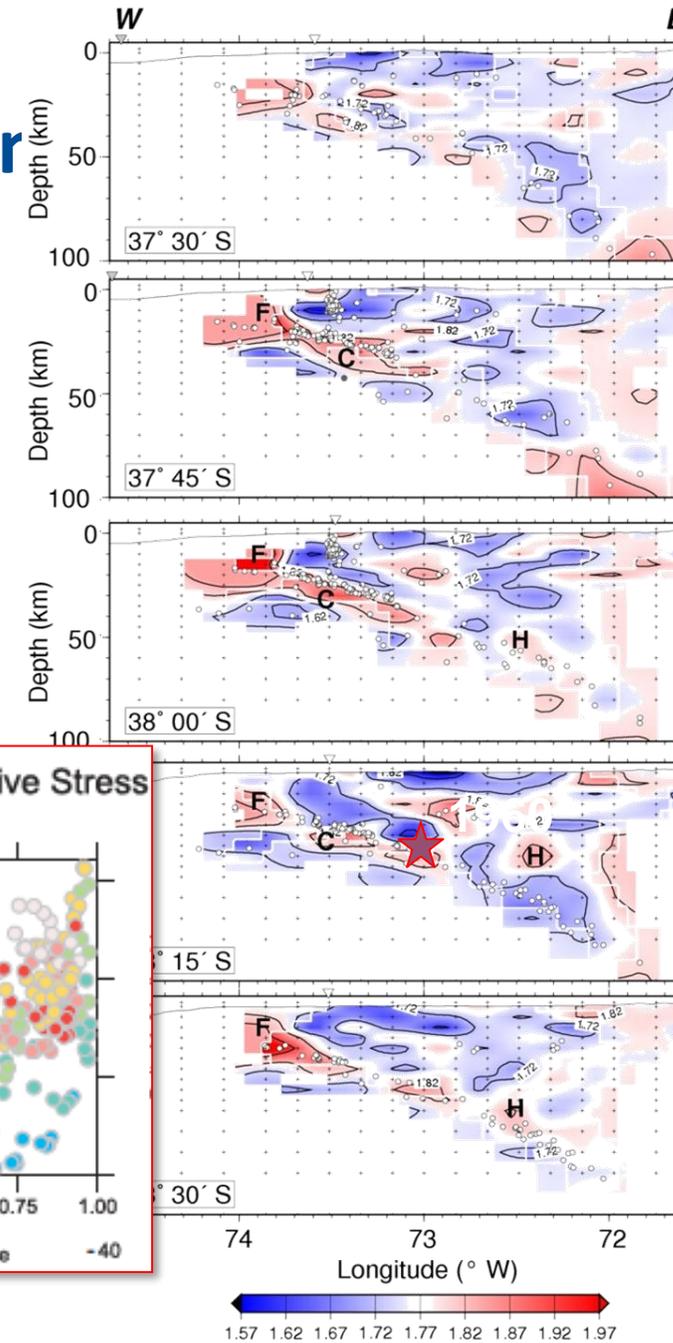


Nipress & Rietbrock, 2006

Background seismicity and V_p/V_s ratio show correlation with geodetic locking



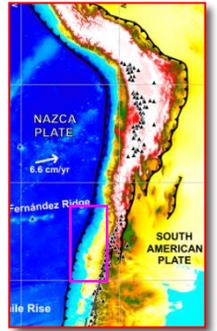
Moreno et al., 2014



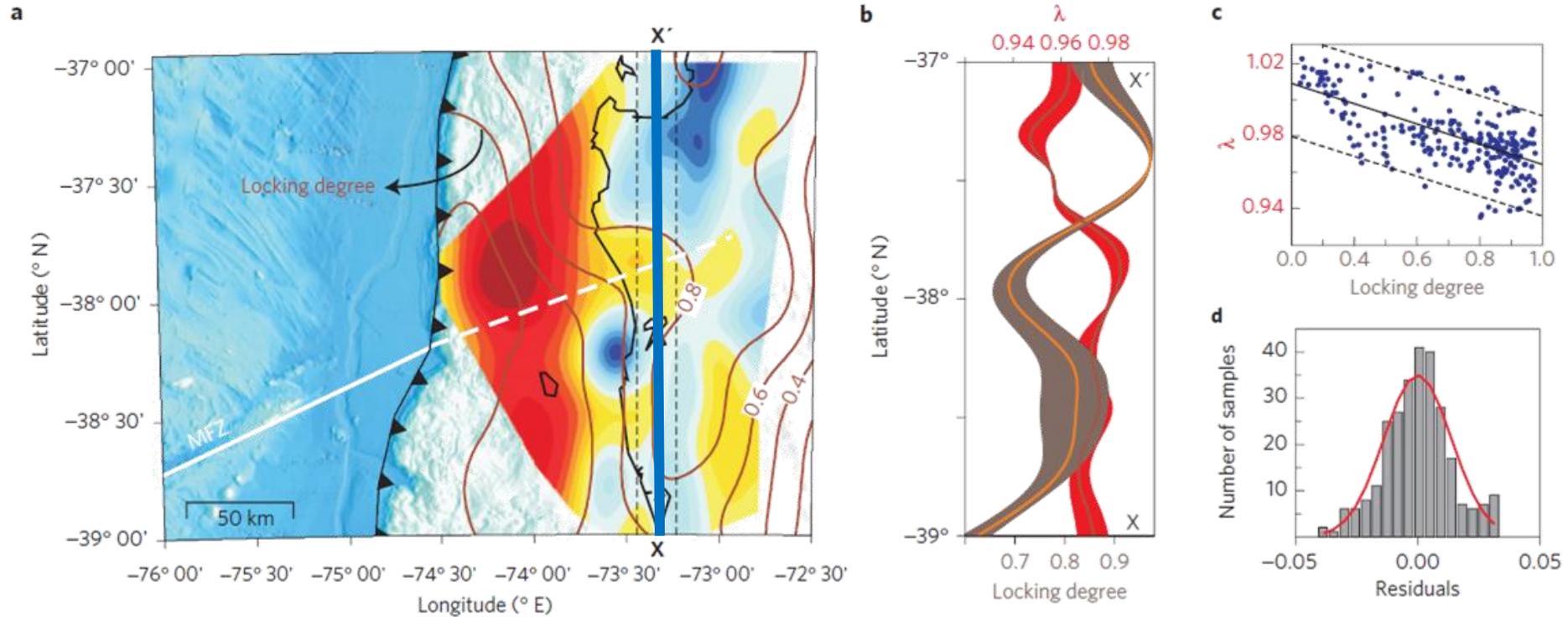
Haberland et al. 2009 v_p/v_s

2010 Maule rupture zone

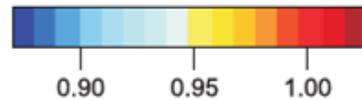
1960 Valdivia rupture zone



Pore pressure controls locking?



Overpressure ratio



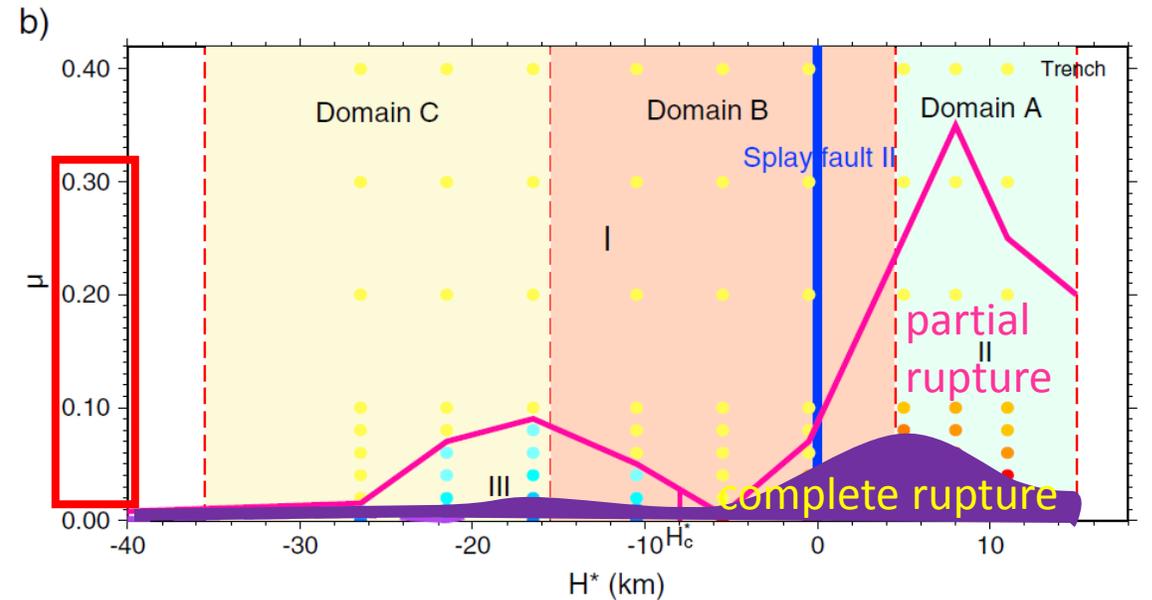
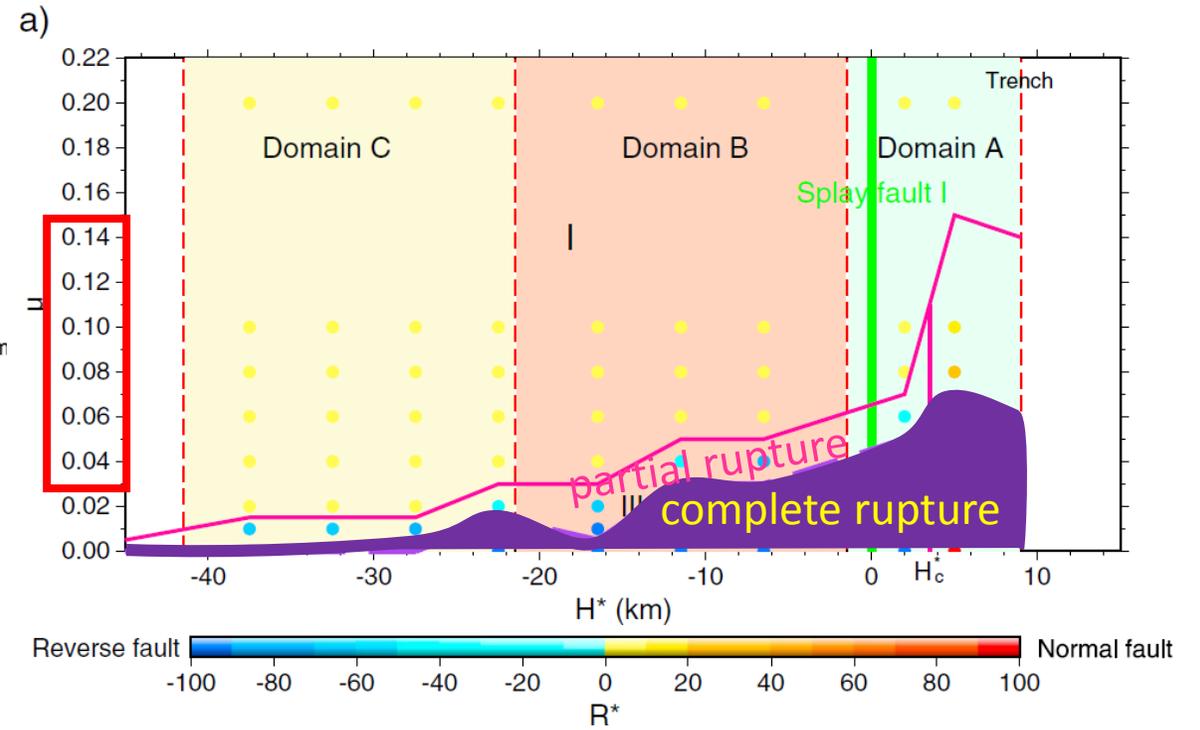
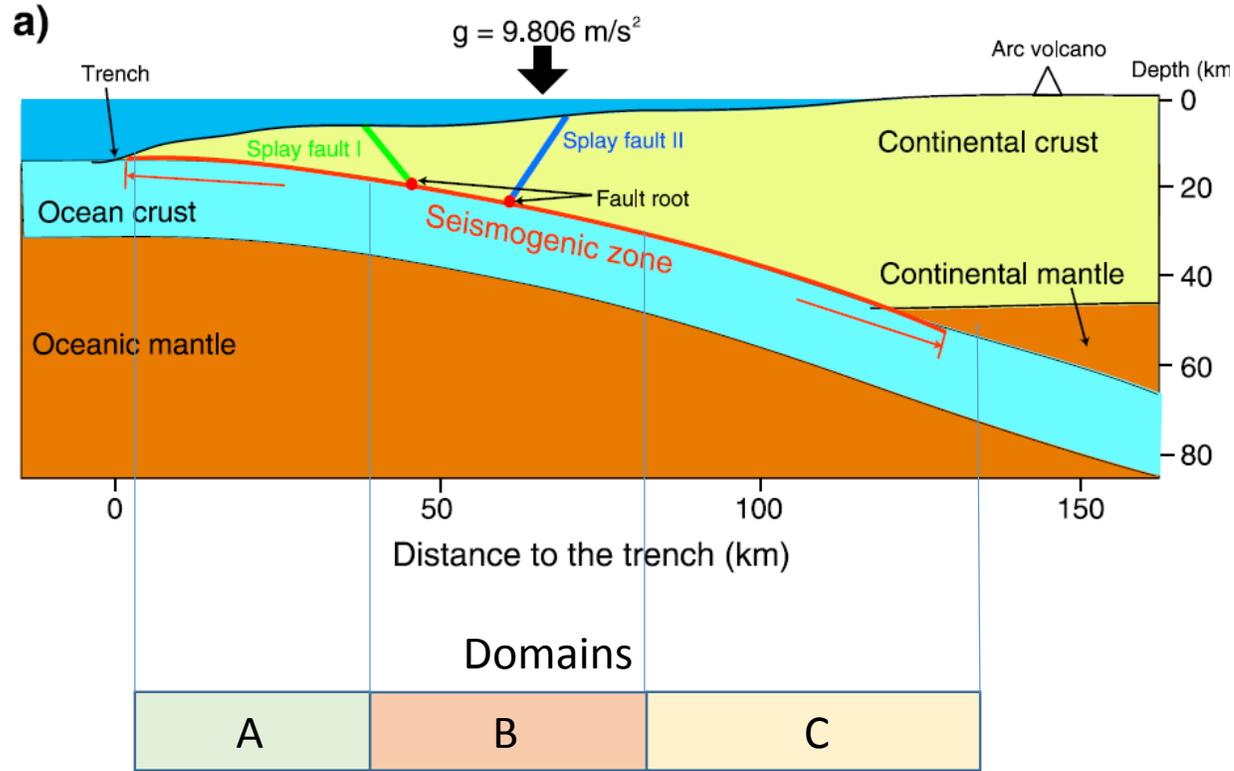
$$\lambda = p_f / p_l \quad \tau = C + \sigma_{zz} \mu (1 - \lambda)$$

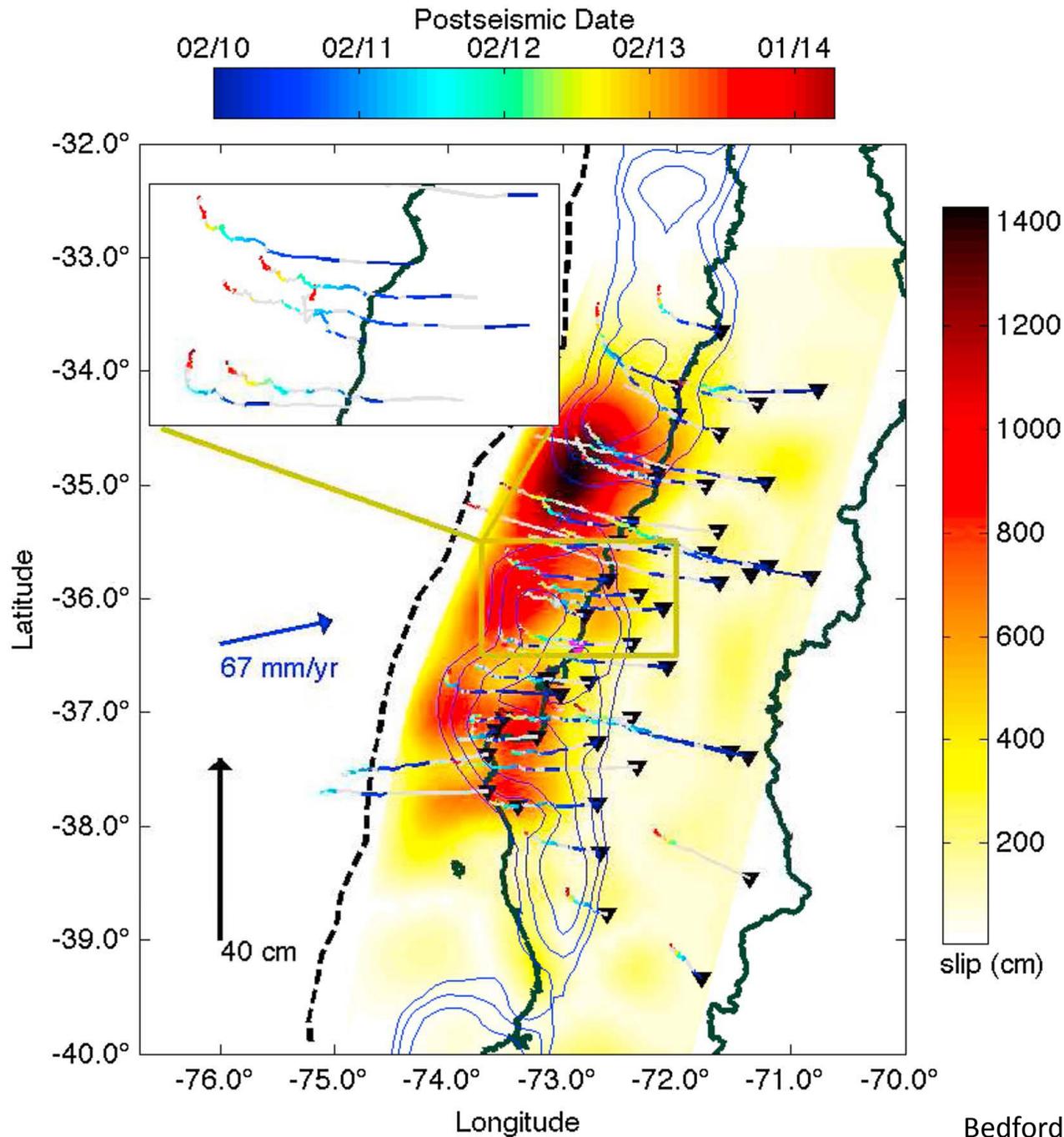
$$\lambda_b \approx 0.90 \text{ to } 1.00$$

$$\mu_{\text{eff}} \text{ on plate interface} \leq 0.05$$

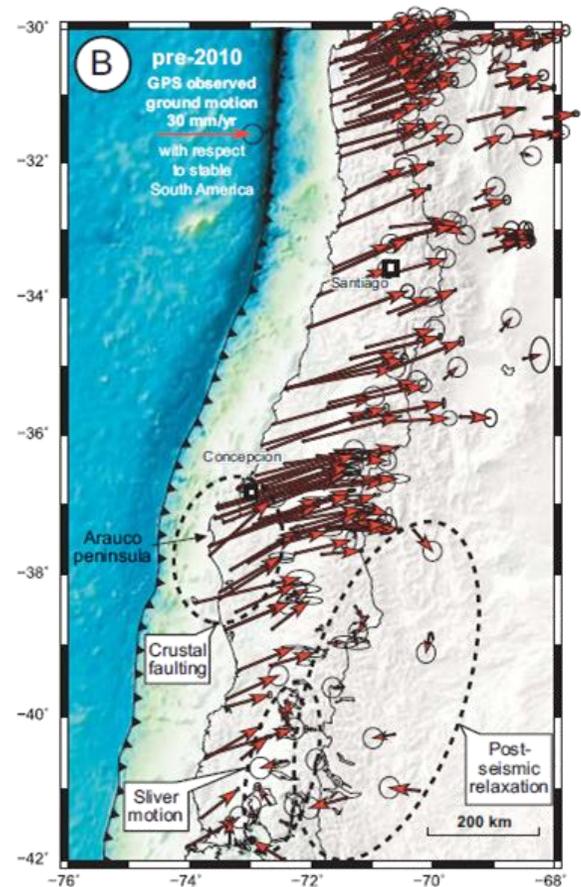
Moreno et al., Nature Geoscience 2014

Strength constraints for forearc faulting





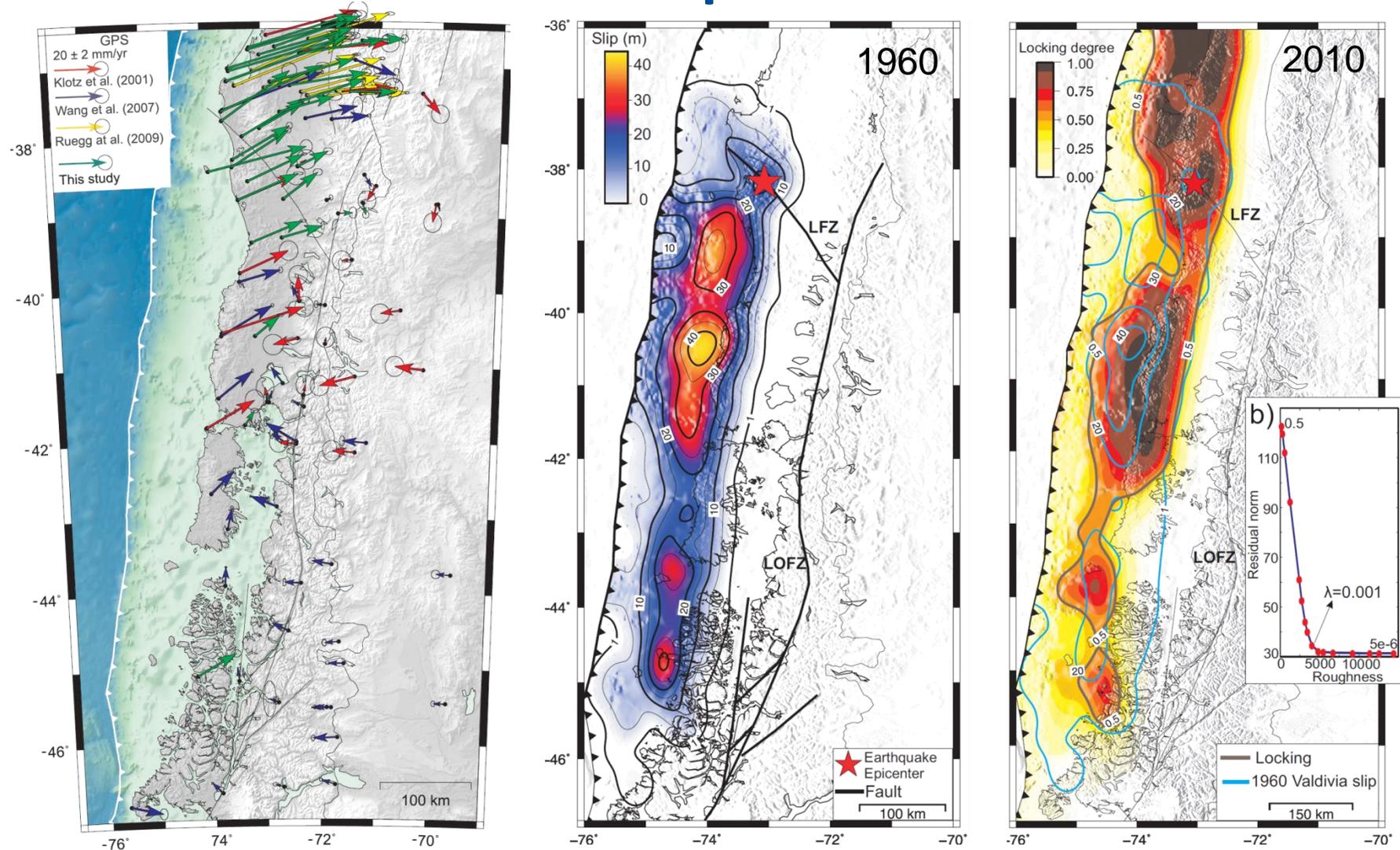
Fault healing and relocking after the Maule 2010 earthquake within 1 year



pre-Maule interseismic loading

Postseismic locking recovers at maximum slip domain

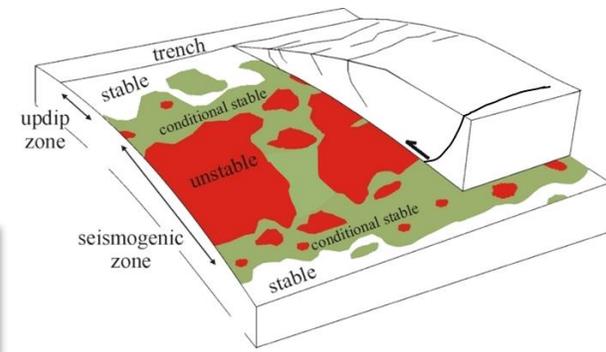
– the Mw 9.5 Chile earthquake of 22.5.1960



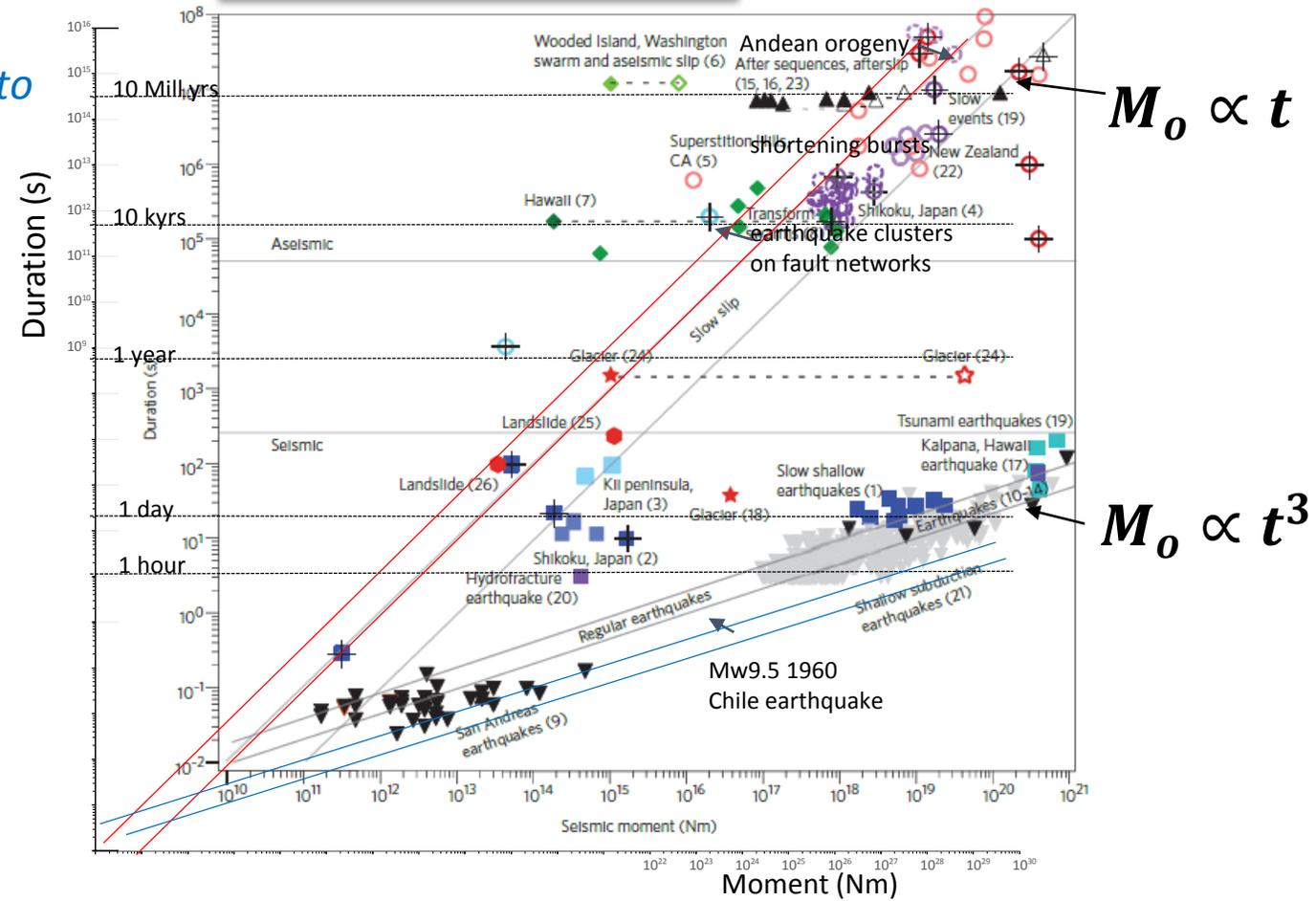
Moreno et al., 2009; Moreno et al, 2012

Conclusions and questions

- The plate interface is extremely weak and has transient strength evolution, *all due to hydraulic system and its transients?*
- Faults may creep **or** break **and** they respond to very small stress changes, *but what makes them respond to stress changes of just a few 10s kPa?*
- Locking and asperities at interface appear stable and control localization of upper plate deformation, *but over which time scales and is it always the hydraulic system and its variability?*
- Upper plate deformation and faulting occurs in ?cyclic mode driven by megathrust cycle, *but how well synchronized and why do faults break in earthquake clusters?*
- upper plate response may be compressional or extensional postseismically, *dependent on ratio of slip/slip-deficit every cycle?*



Energy release time \approx const.



Peng & Gomberg 2010

What this means



effective coefficient of friction
on plate interface ≤ 0.05

friction of banana peel ≈ 0.07

Awarded with 2014 Ig Nobel Prize

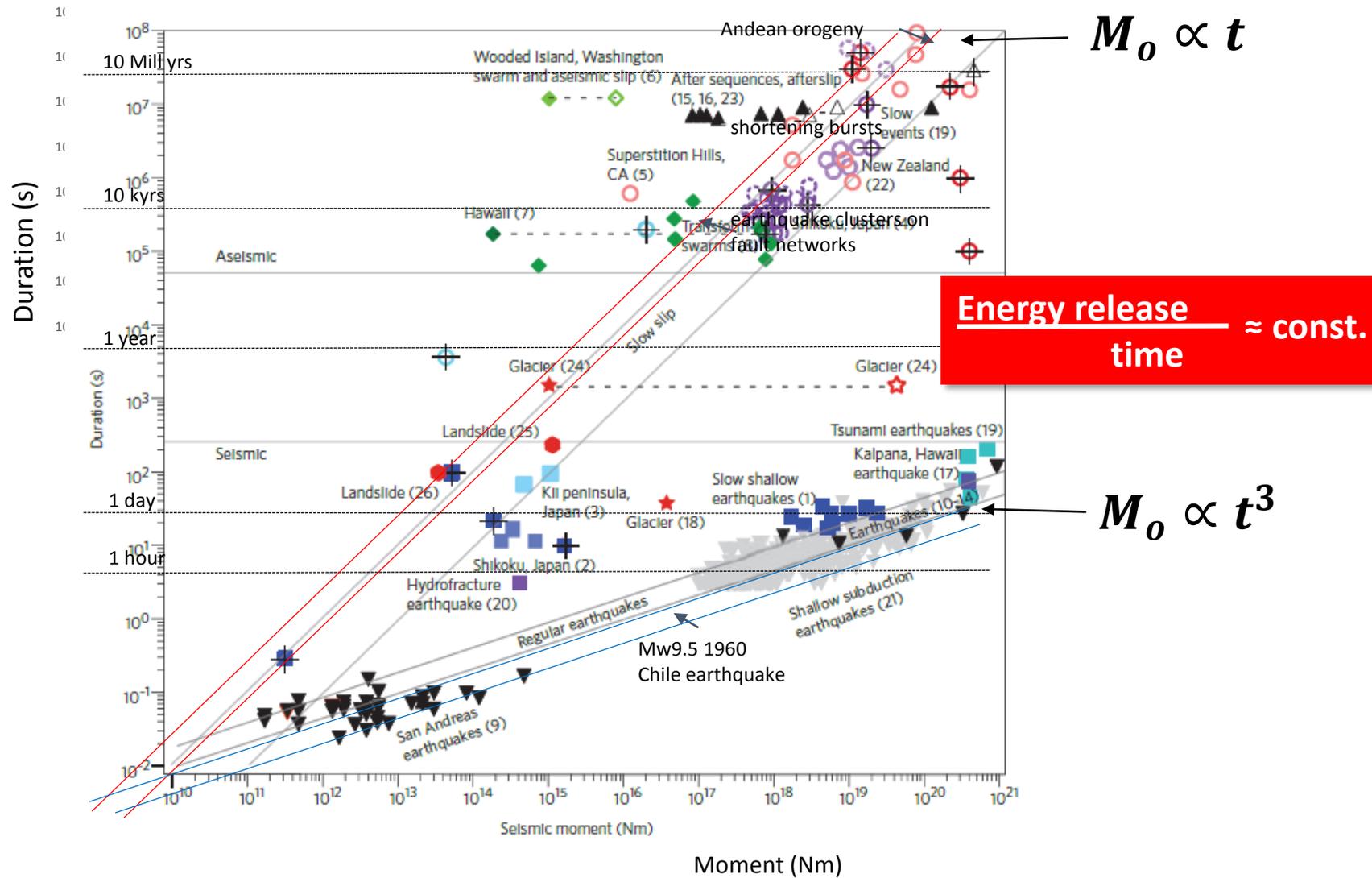
PHYSICS PRIZE:

for measuring the amount of friction between a shoe and a banana skin, and between a banana skin and the floor, when a person steps on a banana skin that's on the floor.

REFERENCE: "[Frictional Coefficient under Banana Skin](#)," Kiyoshi Mabuchi, Kensei Tanaka, Daichi Uchijima and Rina Sakai, Tribology Online 7, no. 3, 2012, pp. 147-151.

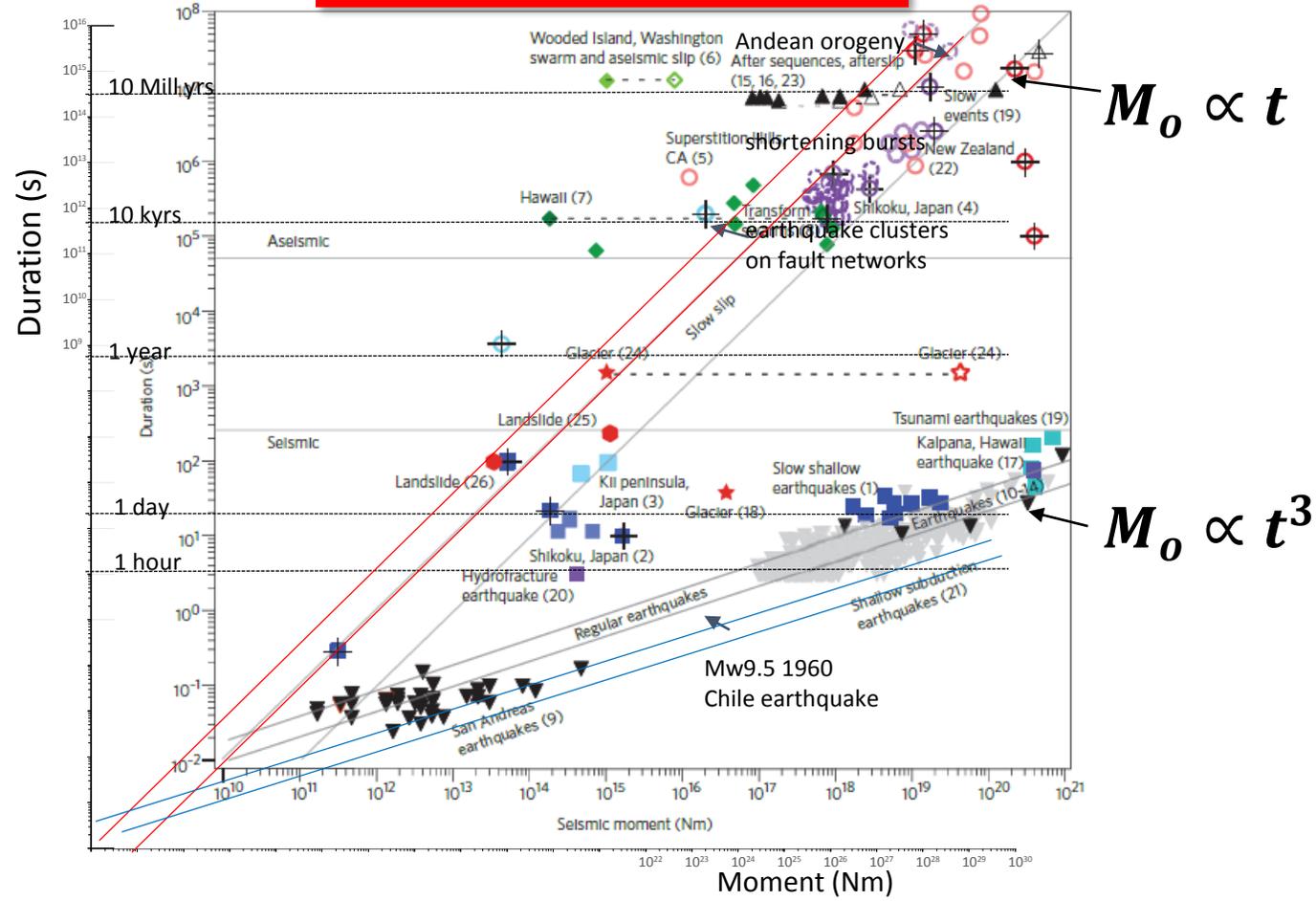
... plate interfaces are weaker than banana peels !

Extending the slow earthquake scaling law

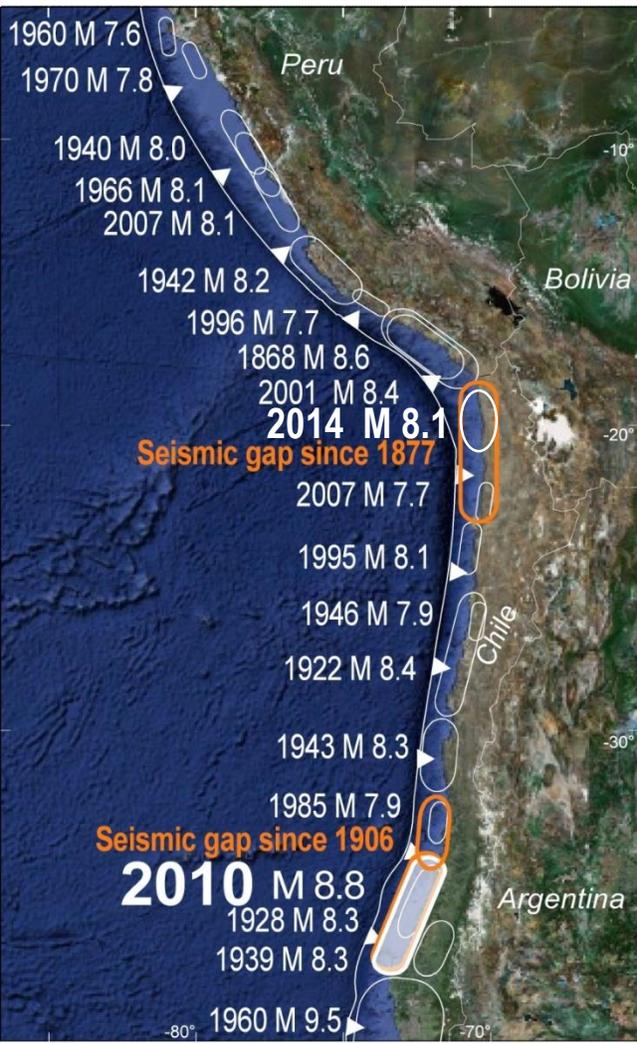


Peng & Gomberg 2010

**Energy release
time \approx const.**



Peng & Gomberg 2010





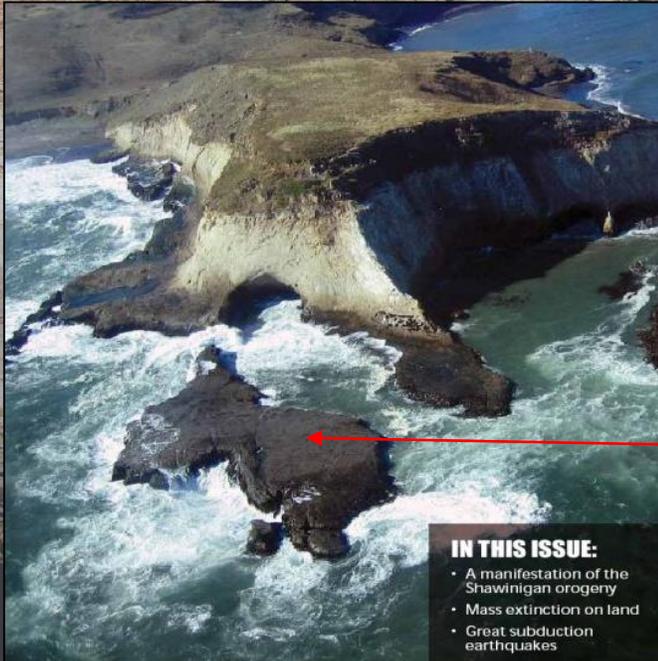
Gehobene Küste

Maule Erdbeben, 27.2.2010



M>8 Erdbeben
1751 ?
(Lomnitz, 1970)
~7 m

Historische koseismische
Hebung
1835 M>8 earthquake
2.4 – 3.0 m
(Darwin, 1851)

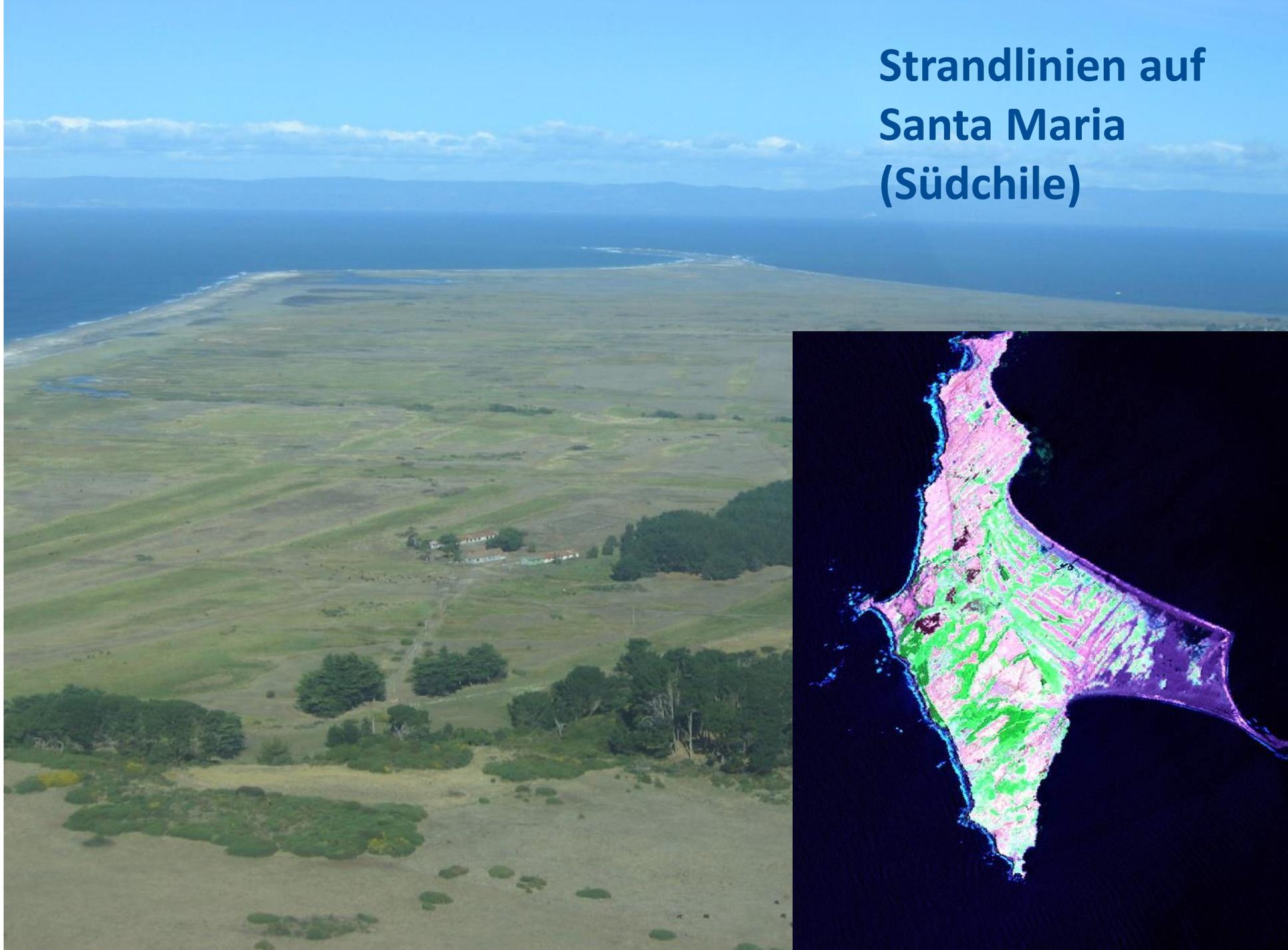


IN THIS ISSUE:

- A manifestation of the Shawinigan orogeny
- Mass extinction on land
- Great subduction earthquakes

Wellenplattform
gehoben 1960 M=9.5
c. 2.5 m

Strandlinien auf Santa Maria (Südchile)



... und vor 58 Jahren



1960

1837

1737

Moreno et al.
2009

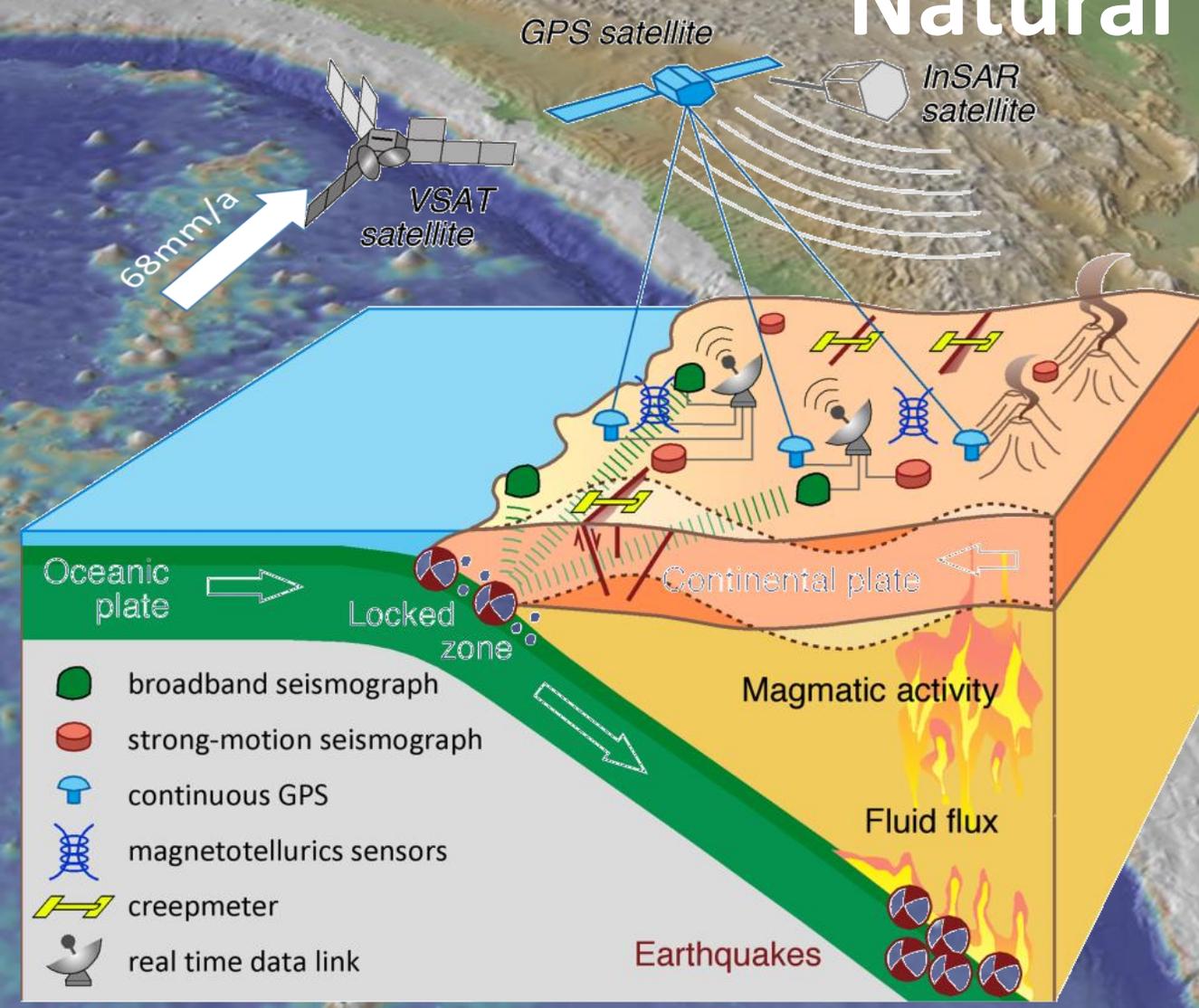
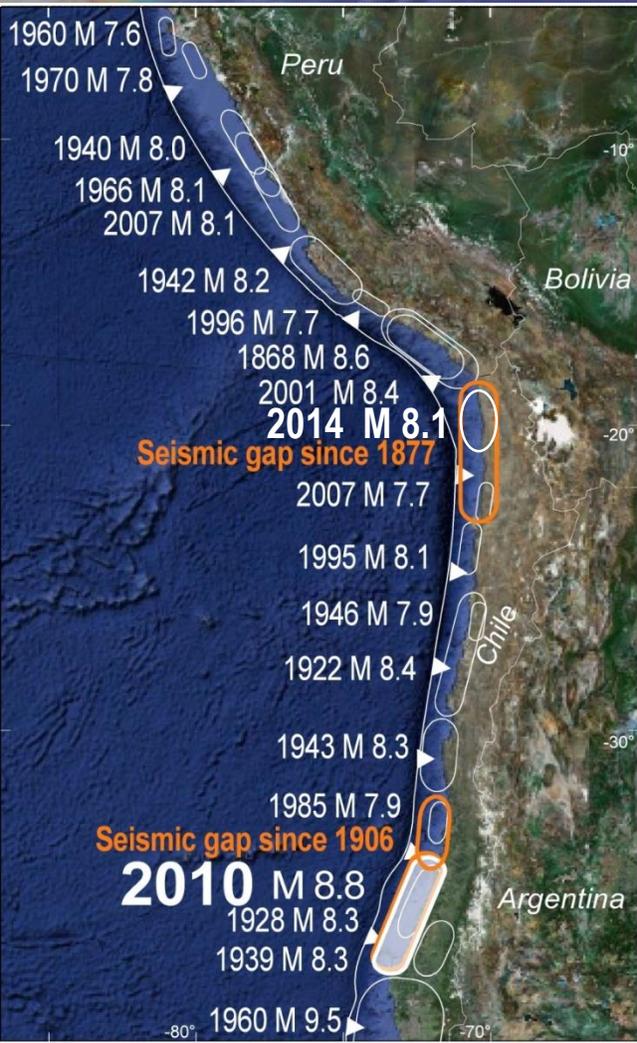


Zeiger 39°44'20.67" S 73°15'19.51" W Höhe 2 m Übertragung 100% Sichthöhe 3.56 km

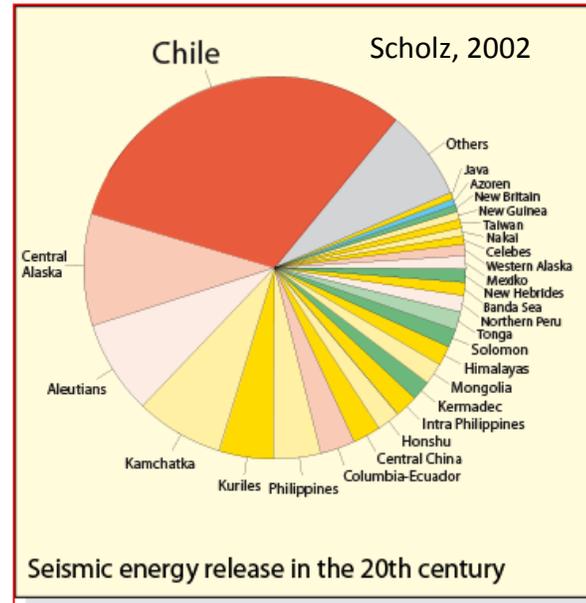
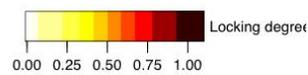
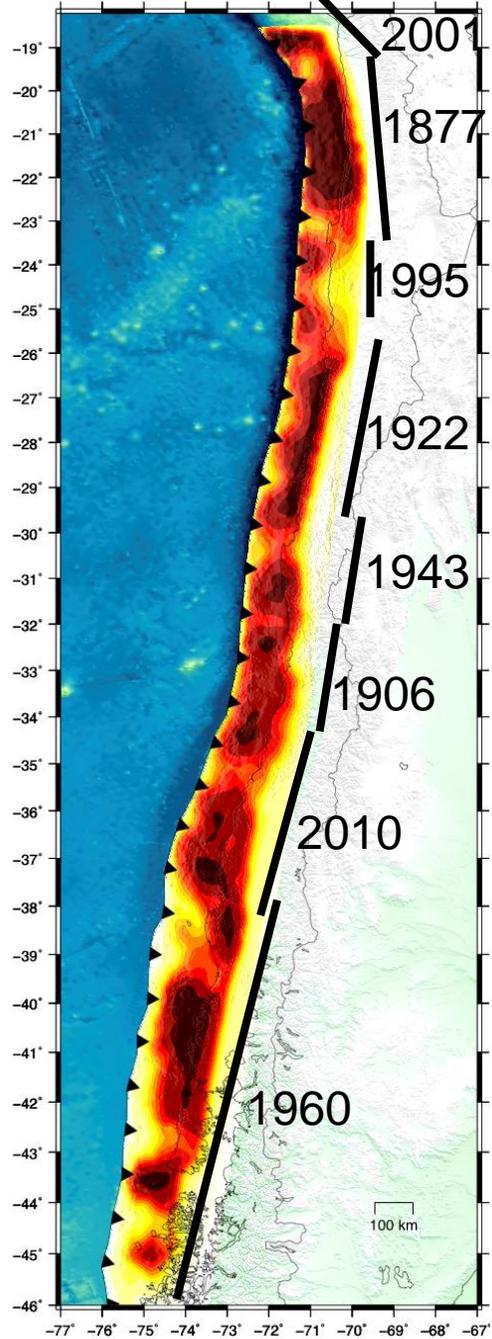
© 2007 Europa Technologies
Image © 2007 DigitalGlobe

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Natural lab Chile



Locking, earthquakes and why Chile is a perfect natural lab



Moreno et al., in prep.

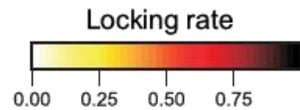
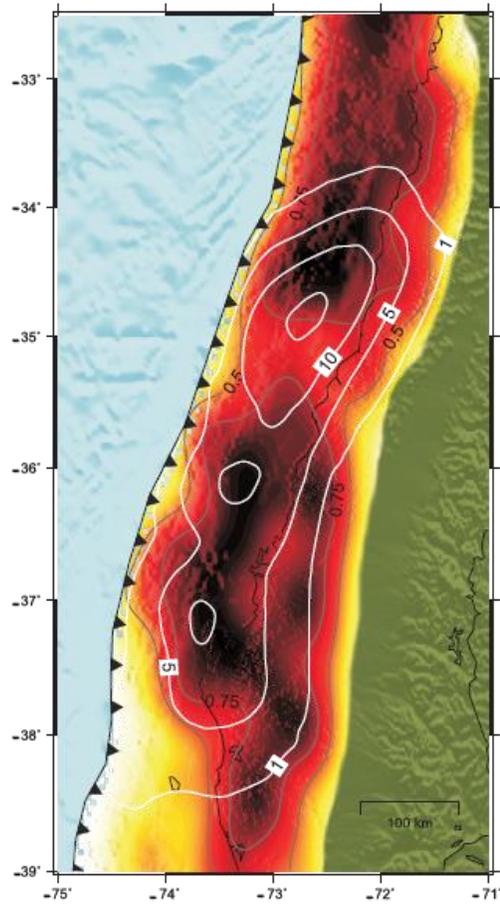
lpoc-network.org



 >100 yrs since last event

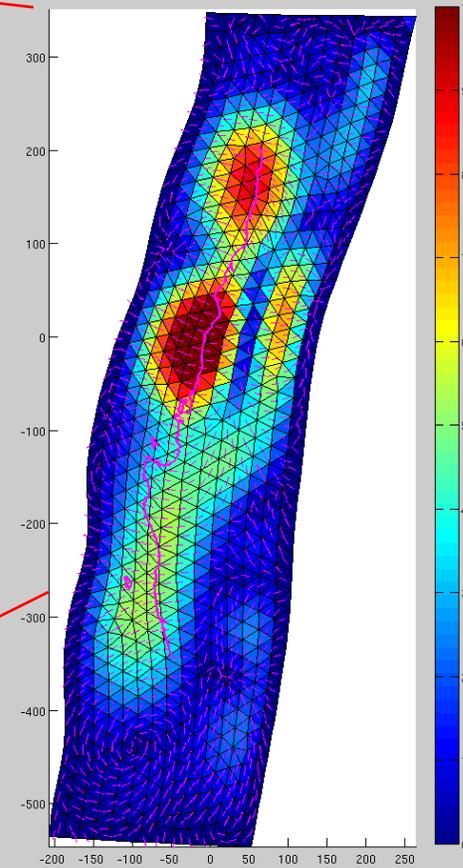
7 months of afterslip and aftershocks – the 8.8. Maule earthquake of 27.2.2010

Slip and locking



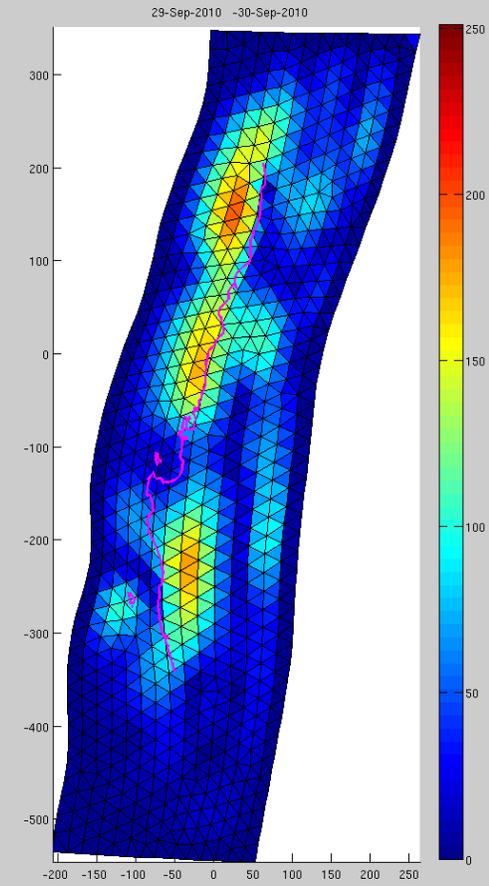
Moreno et al., 2010

Daily slip solution

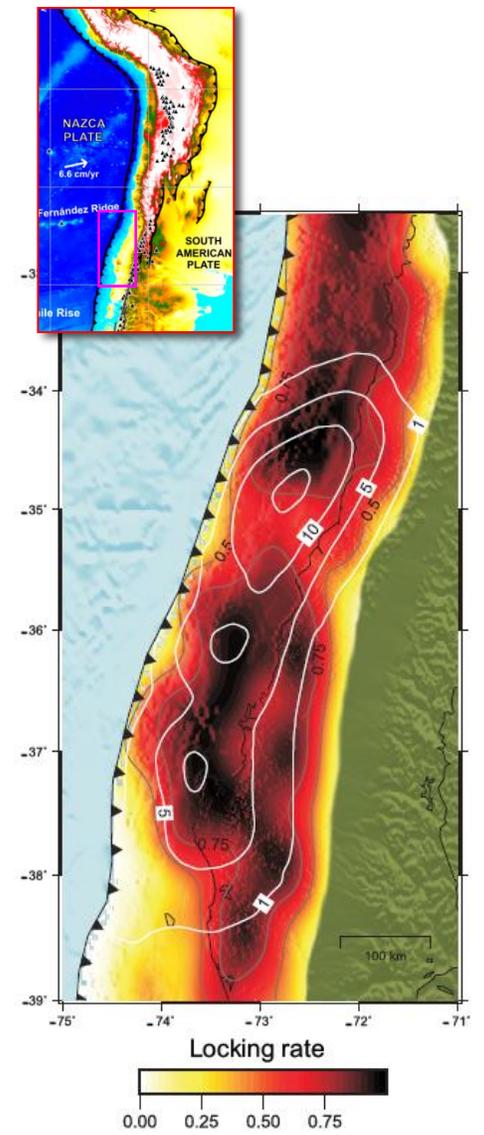
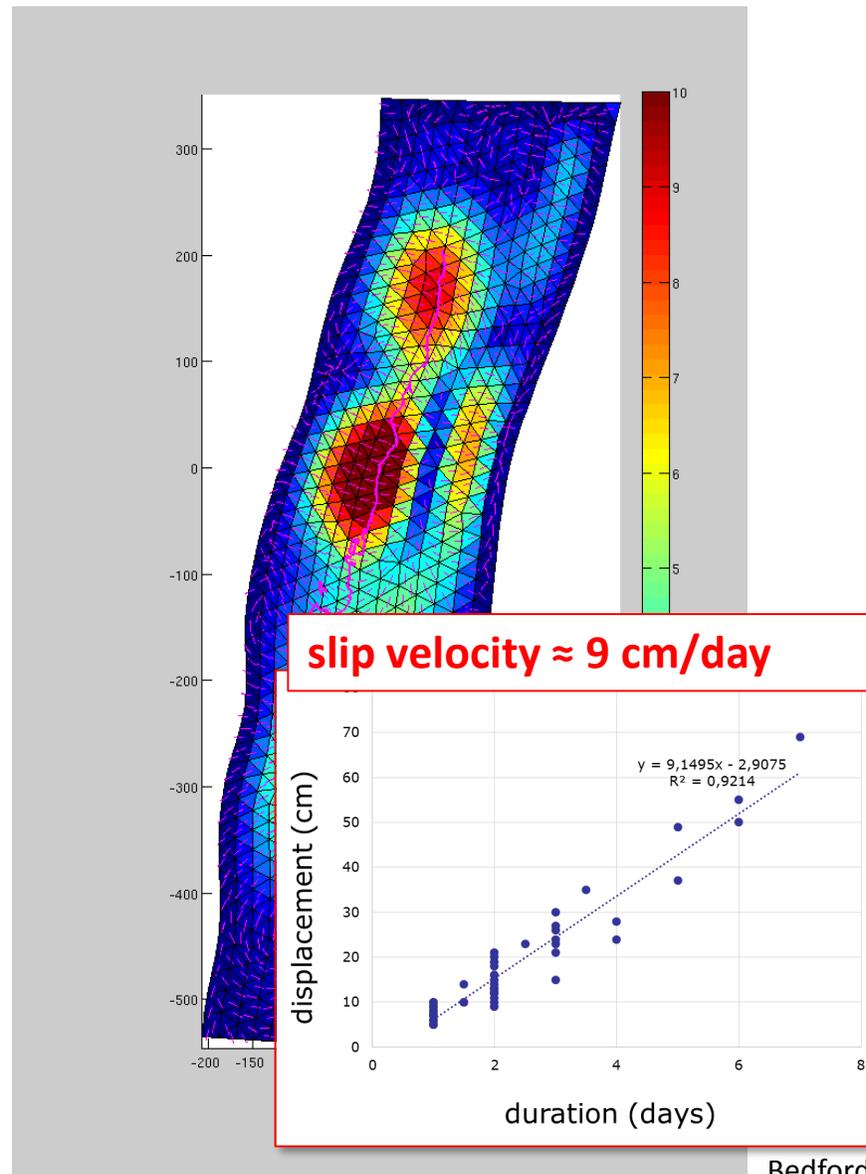
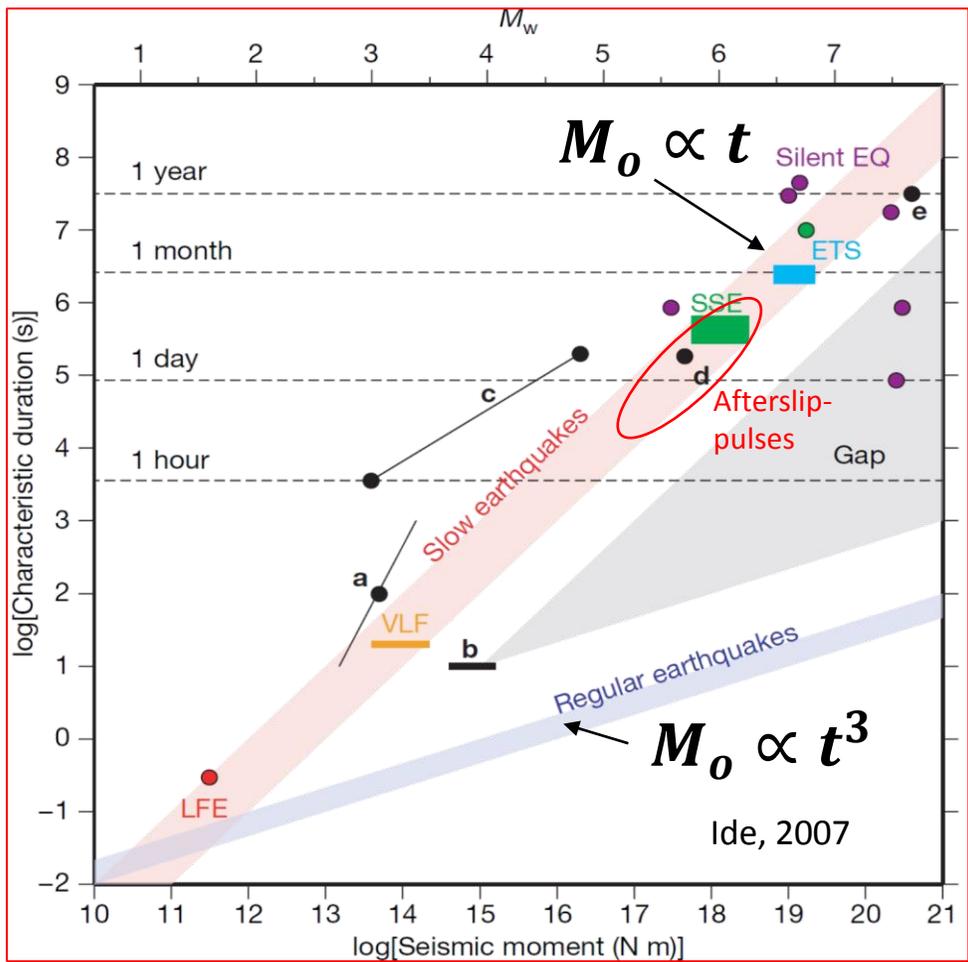


Bedford et al., 2013

Cumulative slip

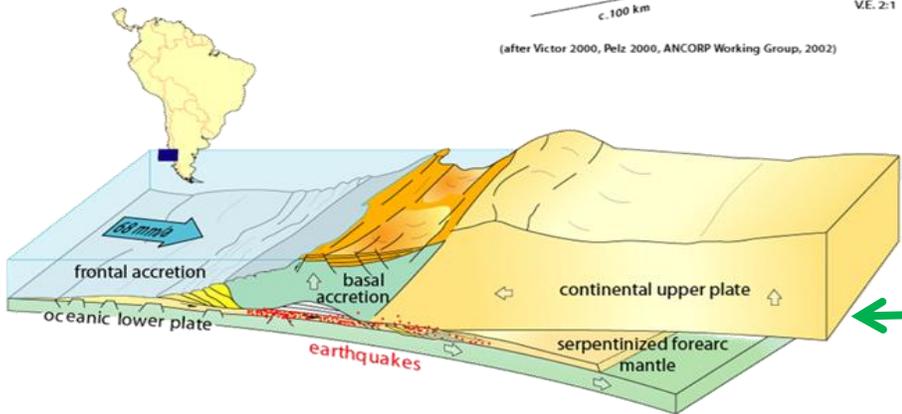
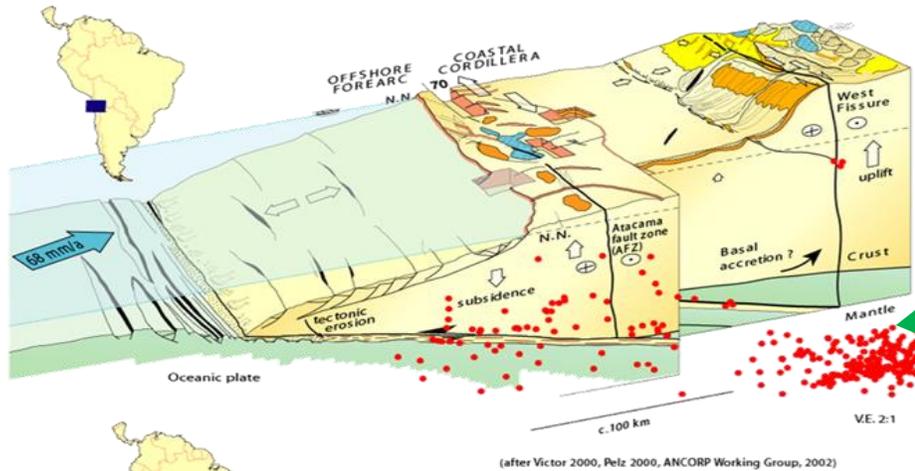


Daily slip solutions reveal two velocity regimes for M8.8., 2010 Maule earthquake 7-month afterslip evolution

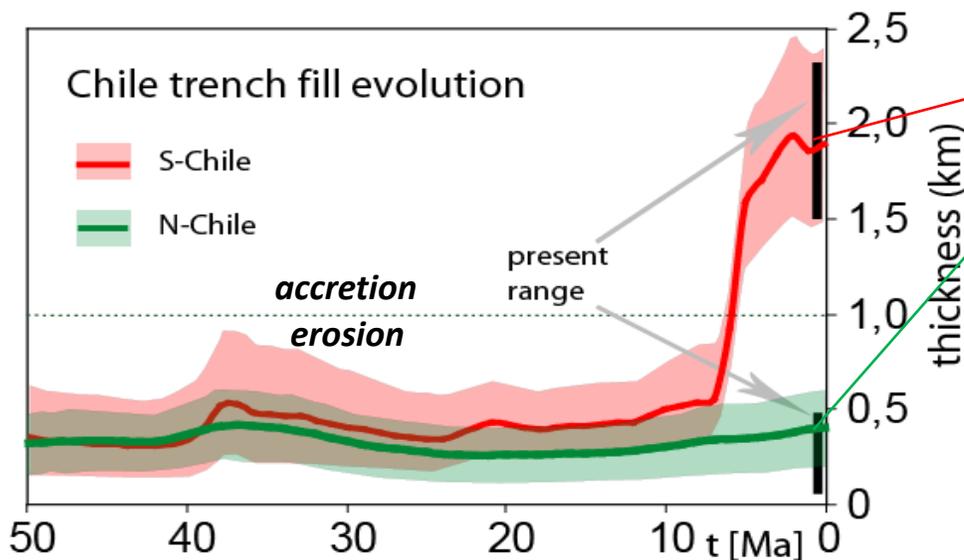


Moreno, Rosenau & Oncken, Nature 2010

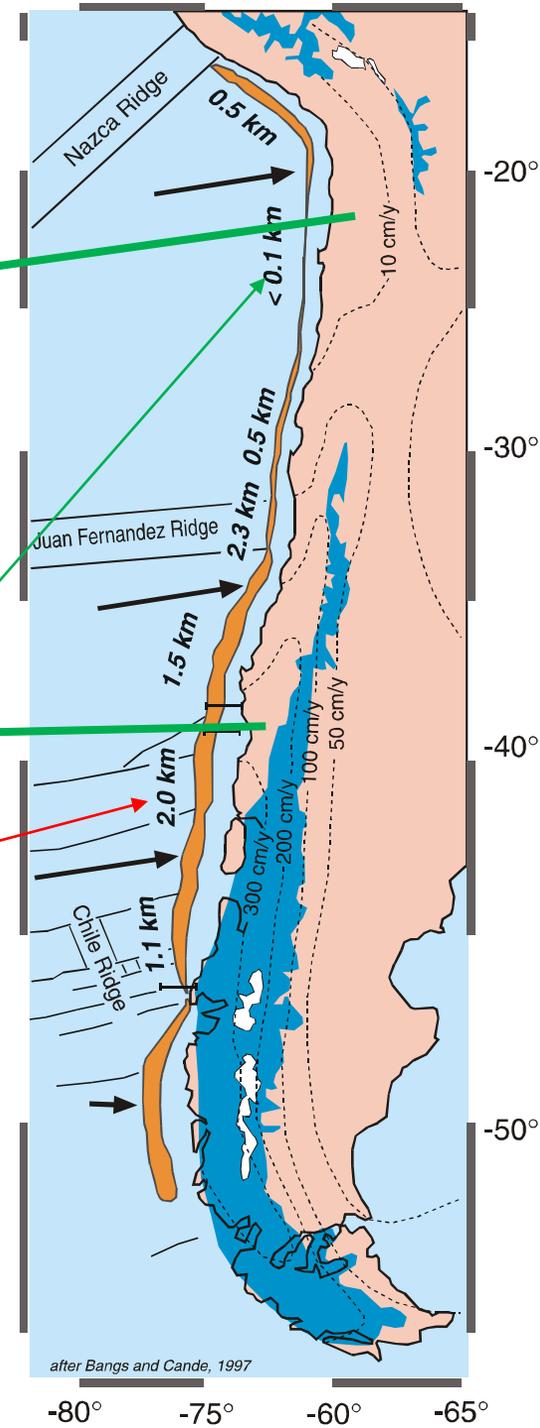
Bedford et al. EPSL, 2013



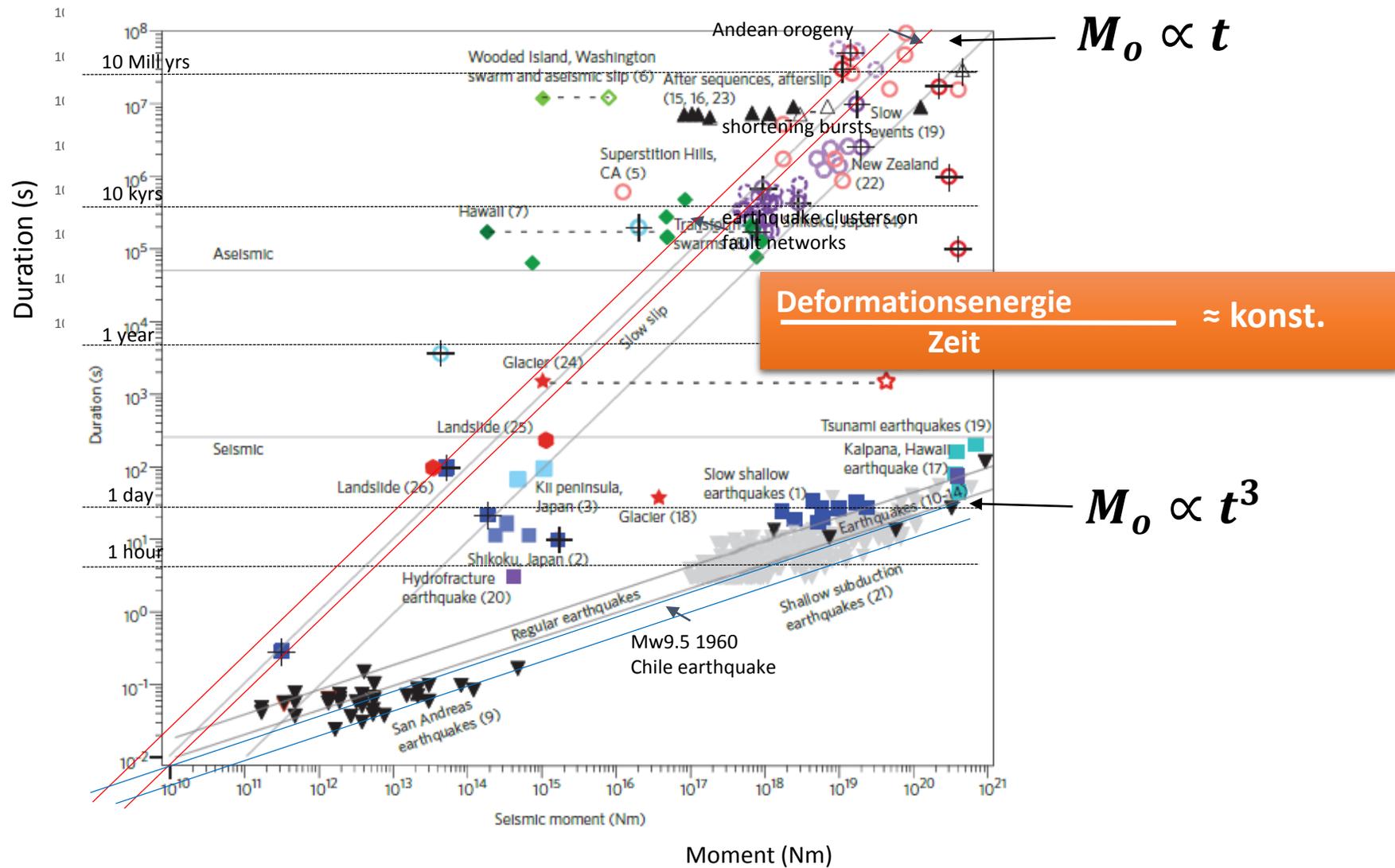
Controls on subduction erosion: the trench fill



Oncken et al., 2006

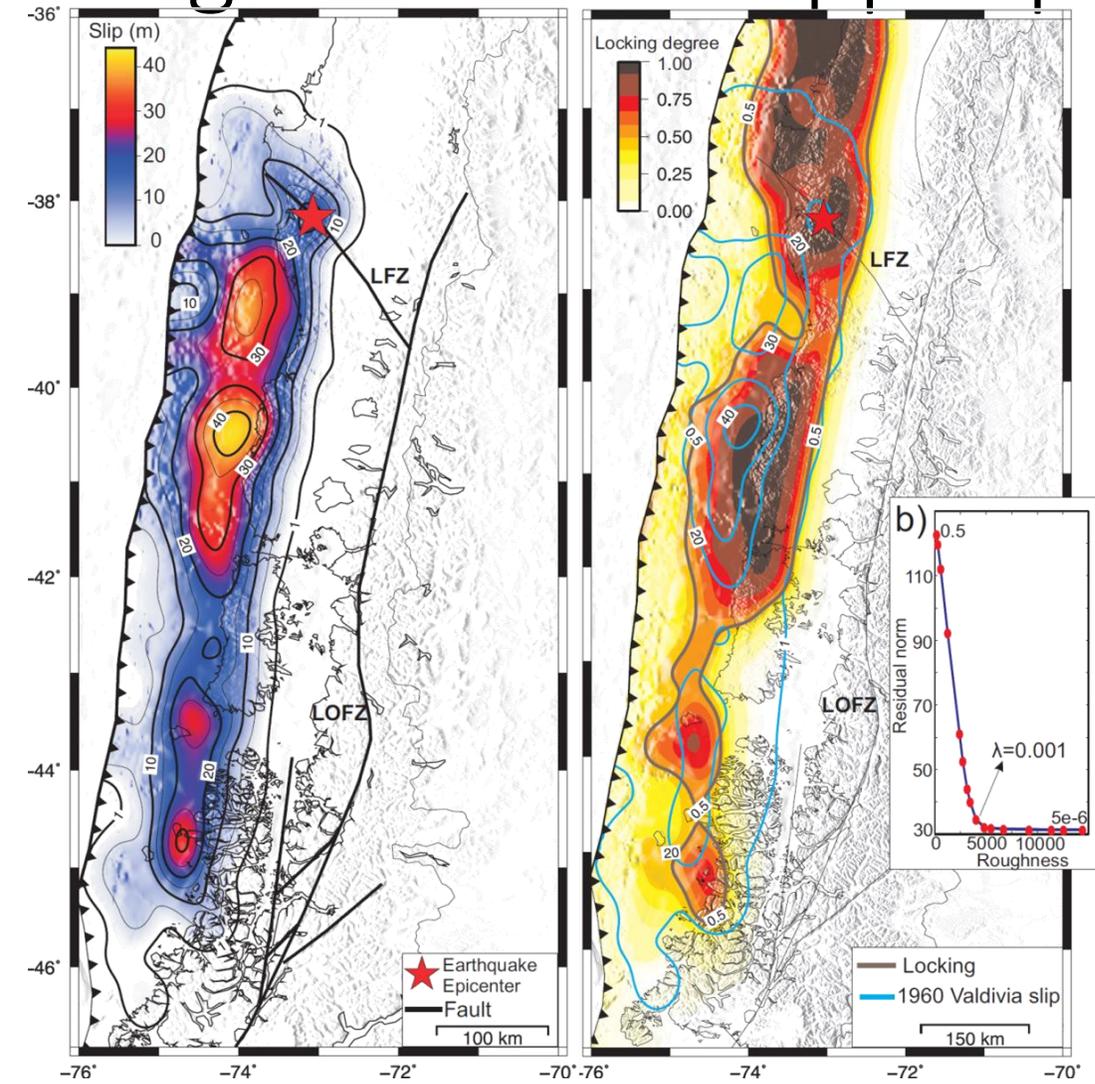


Der erweiterte Erdbeben-/‘Transienten‘-Zoo

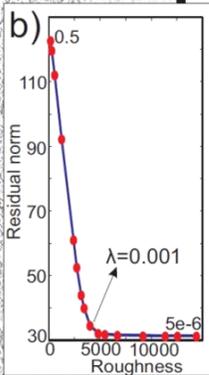
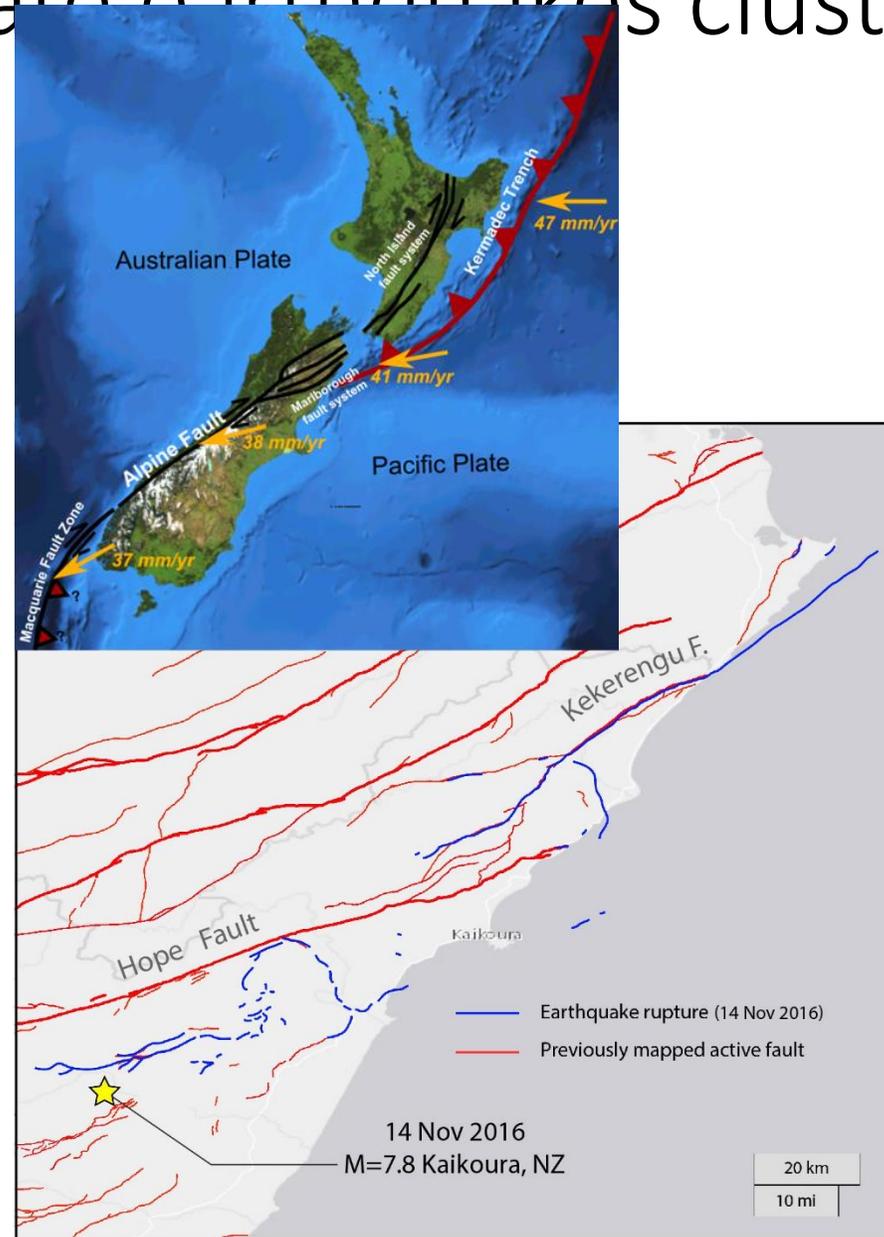


Peng & Gomberg 2010

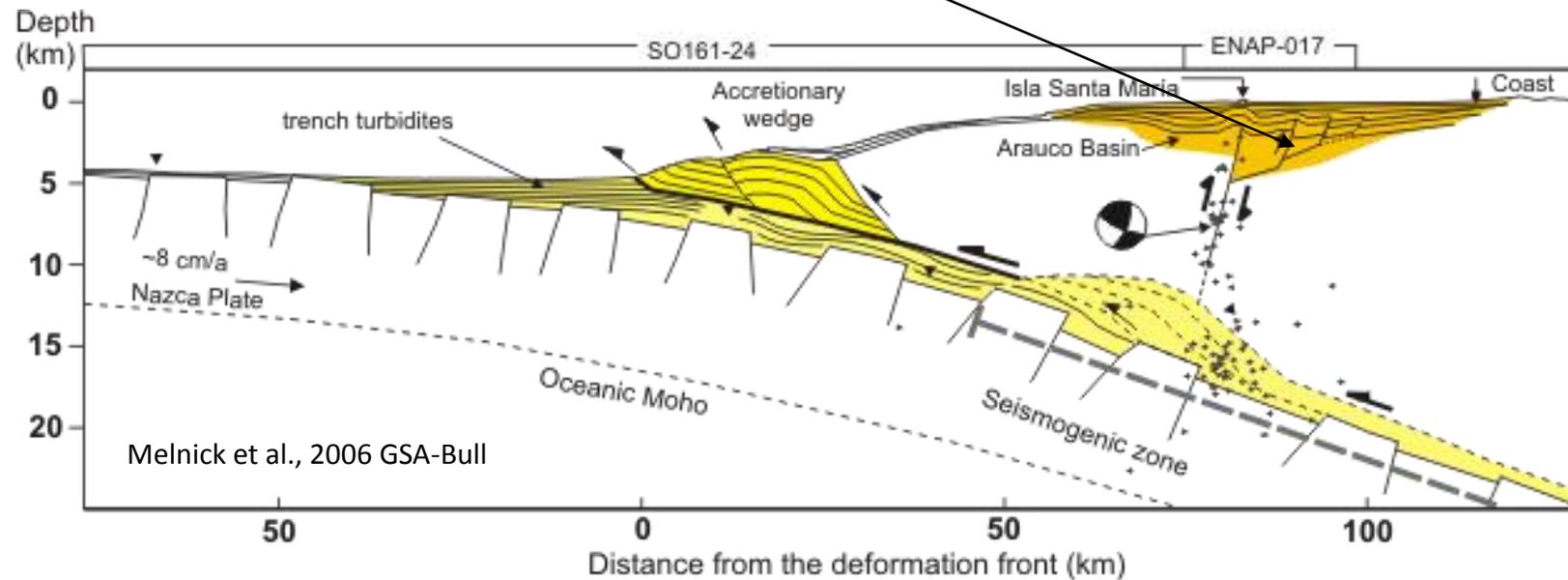
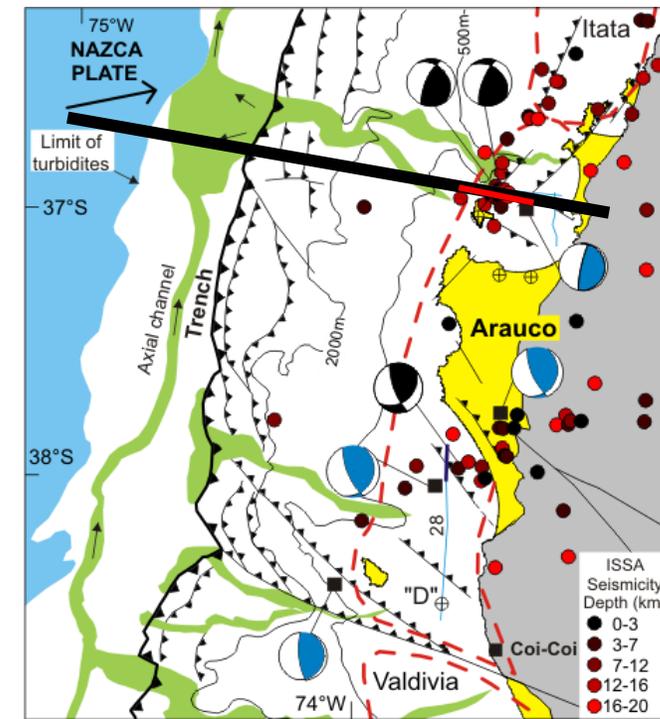
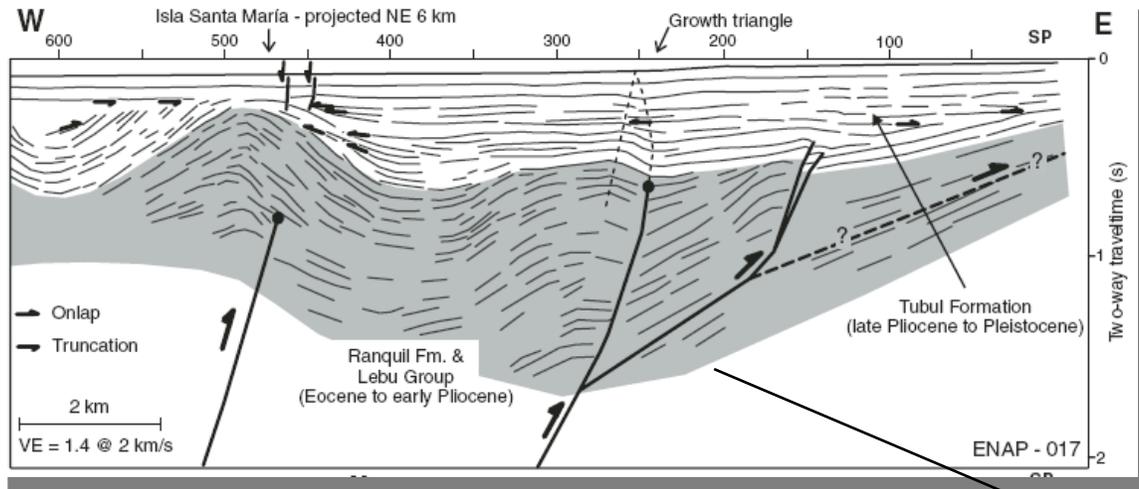
Megathrust and upper plate earthquakes cluster?



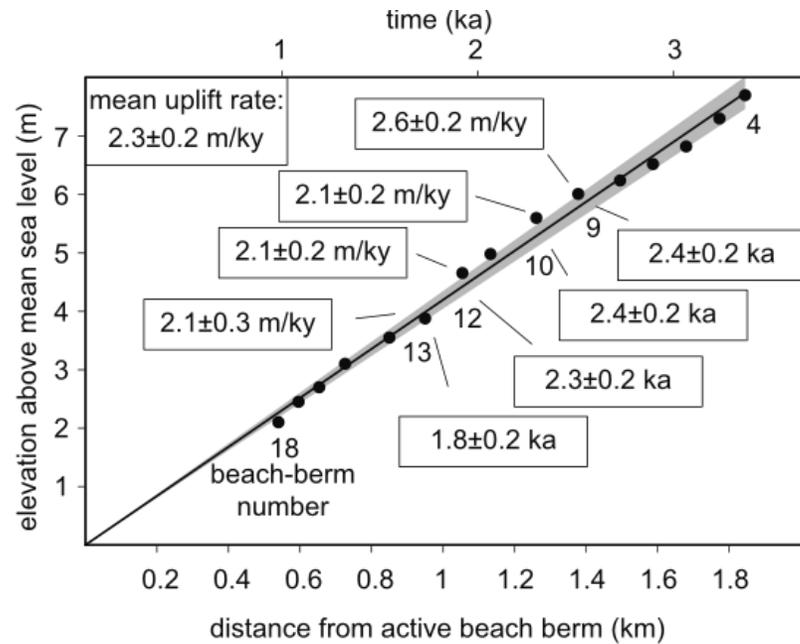
Moreno et al., 2009;
Moreno et al., 2012



Arauco Basin: Stop and go inversion & uplift since 3.6 Ma

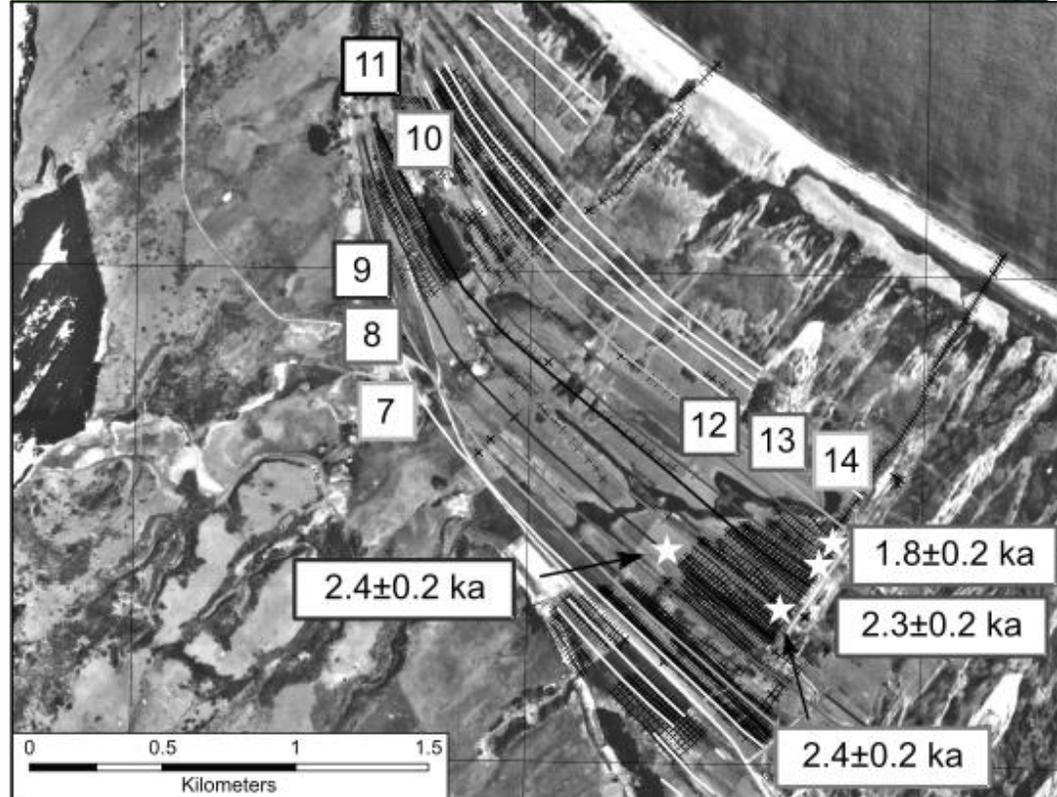
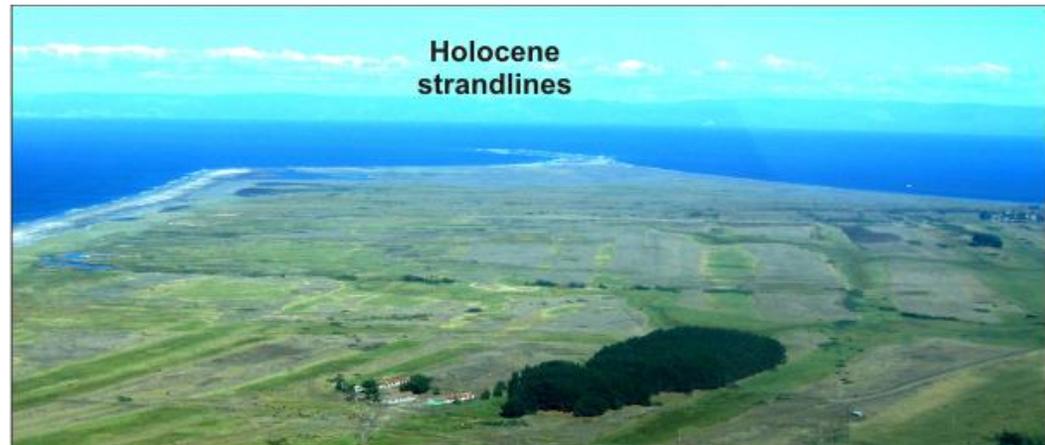


Seismic cycle-related evolution of uplift



Bookhagen et al., 2006

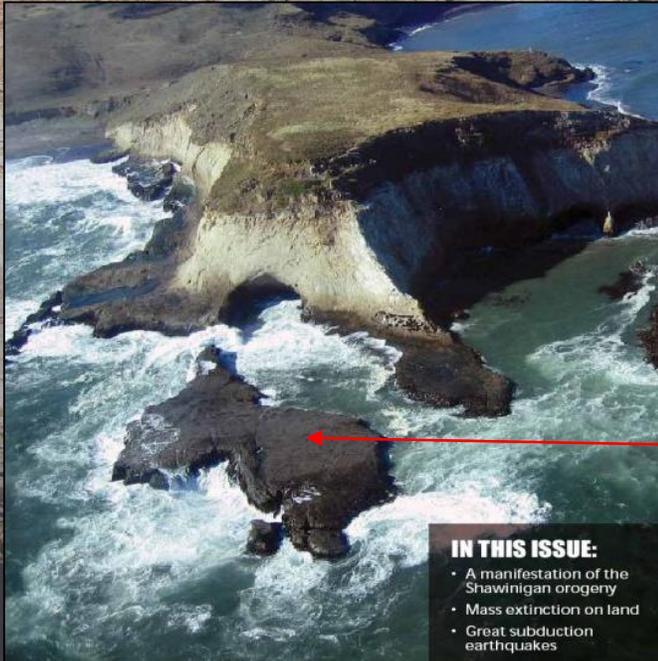
Earthquake recurrence
~ 185 yrs



M>8 earthquake
1751 ?
(Lomnitz, 1970)
~7 m

Historical coseismic
uplift
1835 M>8 earthquake
2.4 – 3.0 m
(Darwin, 1851)

Wave-cut platform
uplifted 1960 M=9.5
c. 2.5 m



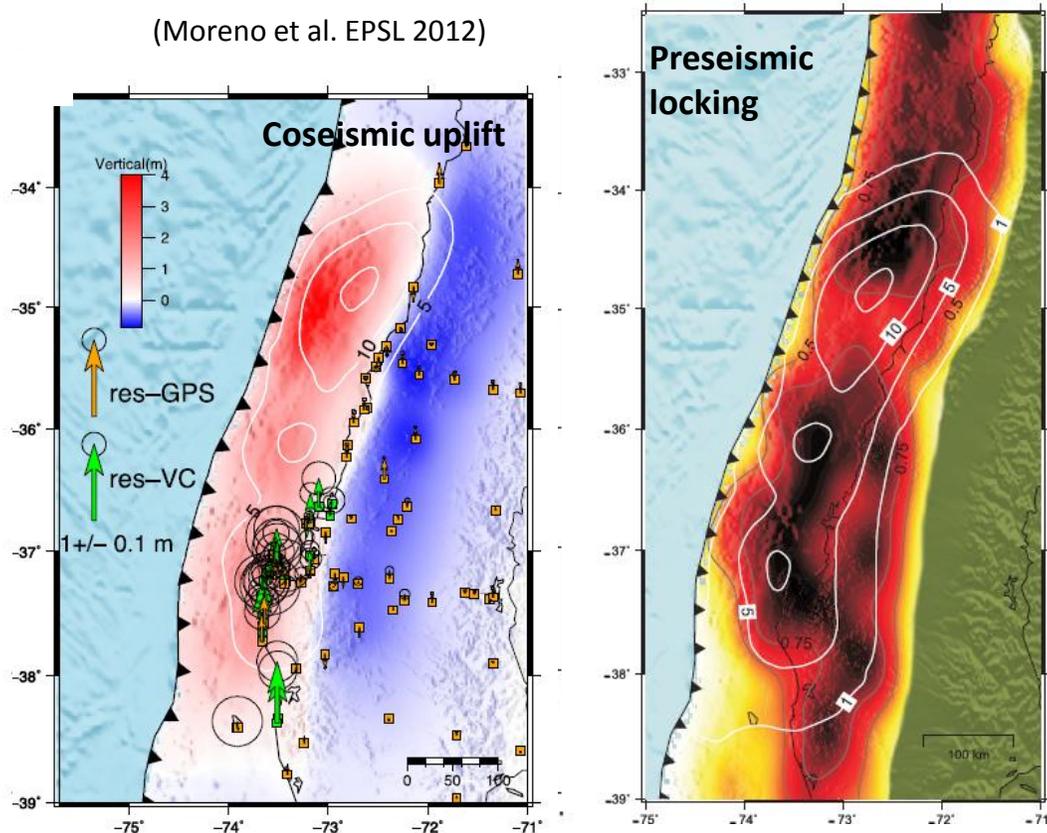
IN THIS ISSUE:

- A manifestation of the Shawinigan orogeny
- Mass extinction on land
- Great subduction earthquakes

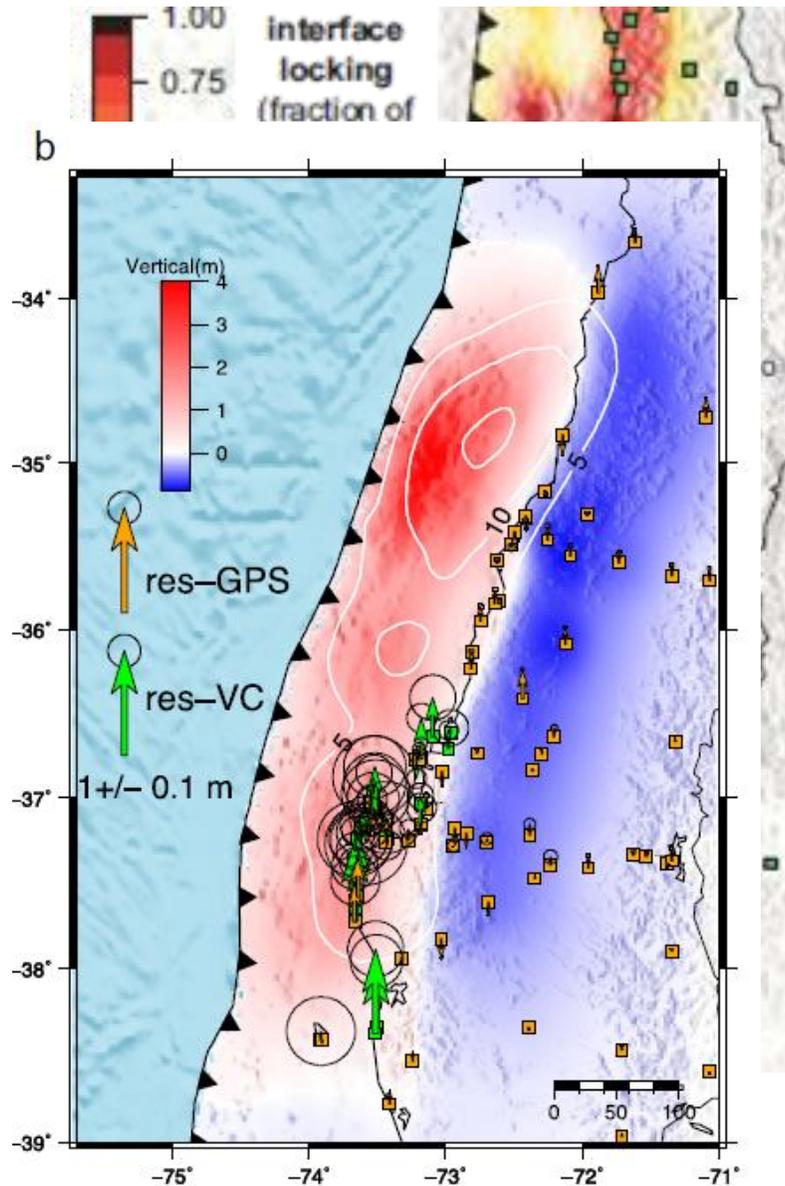
Coseismic slip, preseismic locking and coastal uplift – Mw8.8 Maule earthquake of 27-2-2010



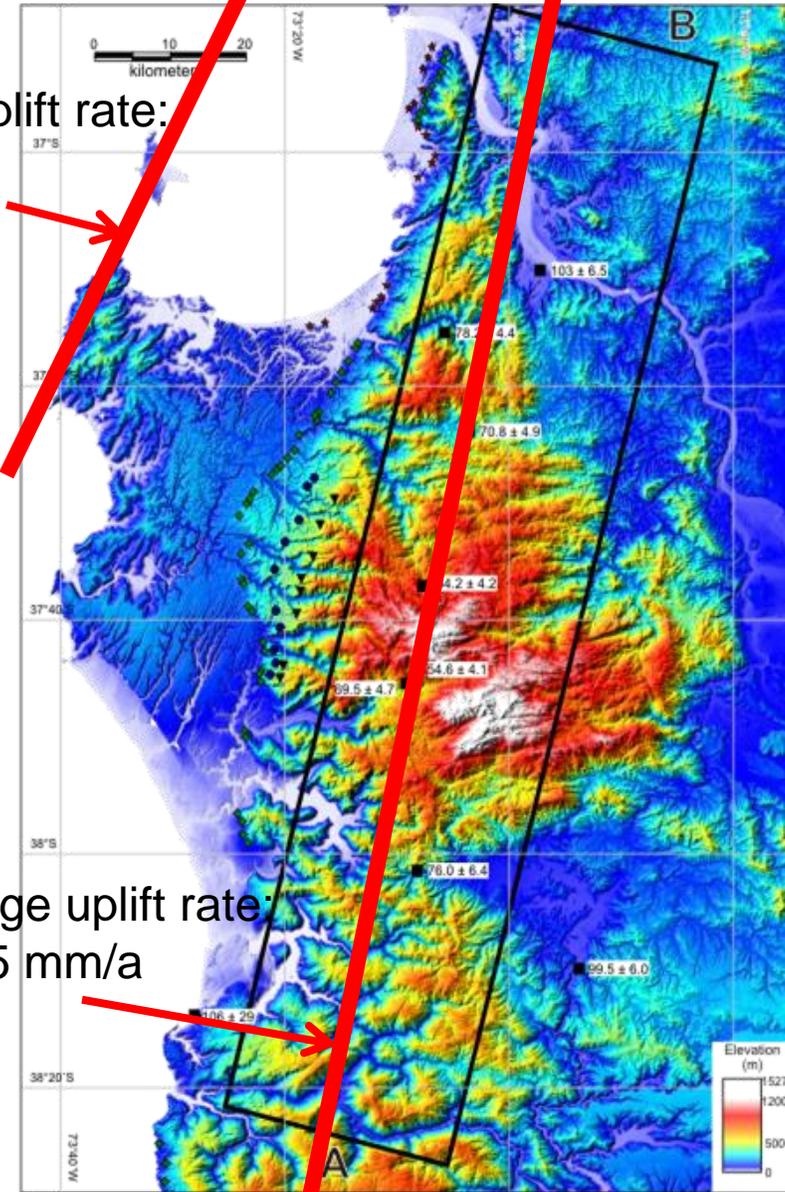
(Moreno et al. EPSL 2012)



Kinematics since c.3 Ma

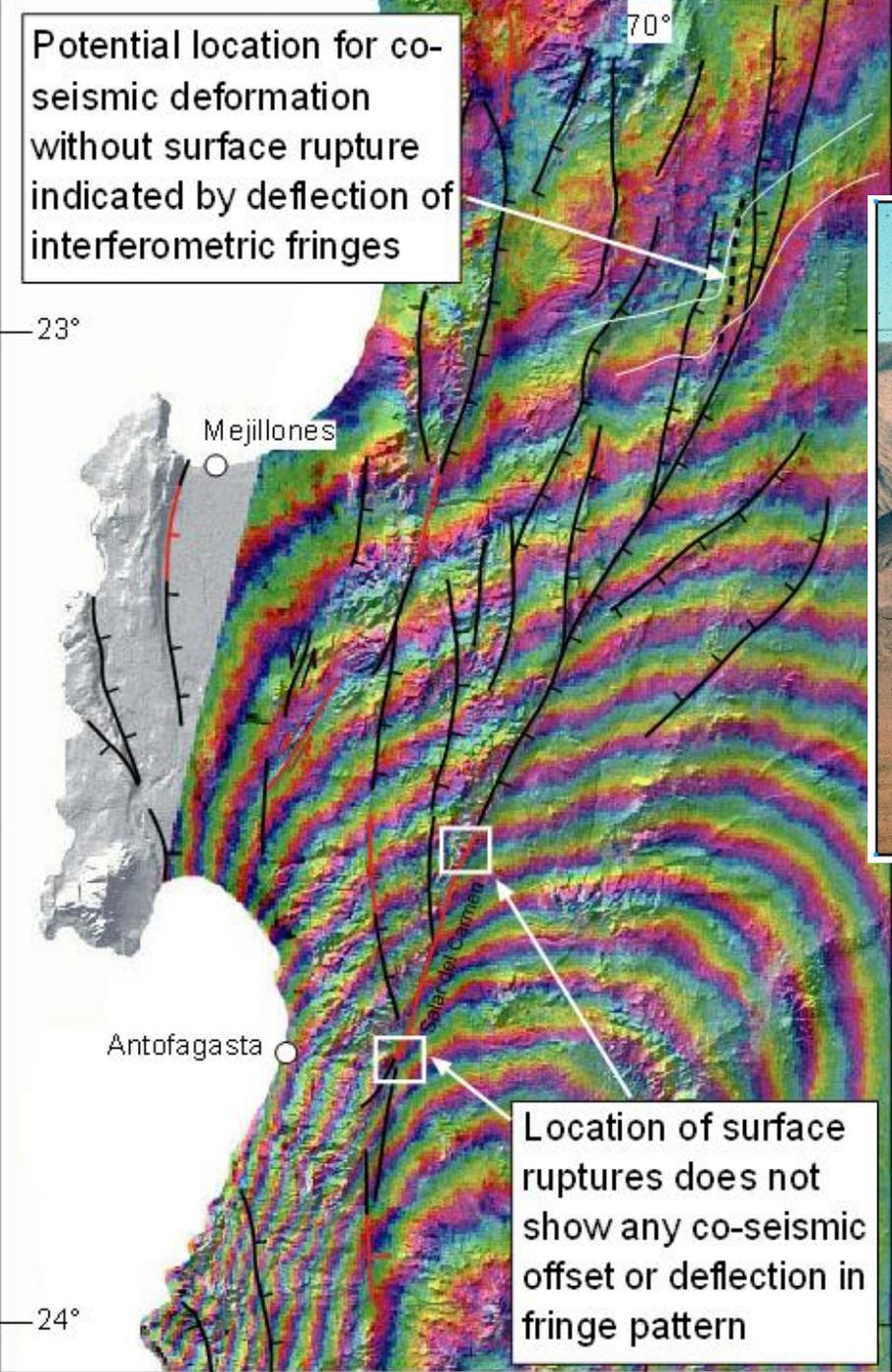


Average uplift rate:
2.3 mm/a

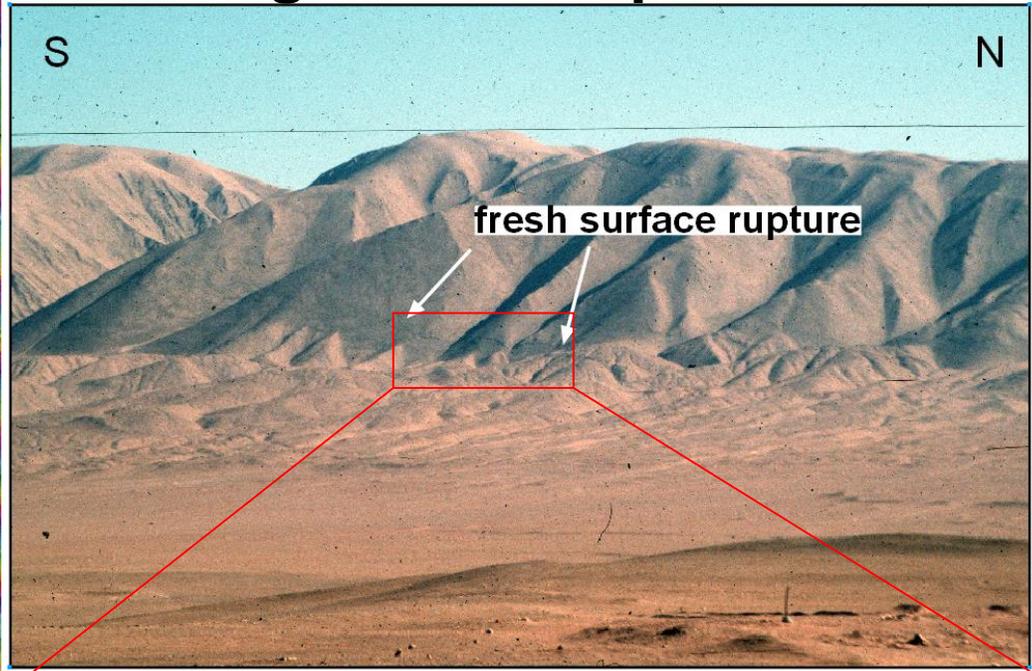


Average uplift rate:
1 - 1.5 mm/a

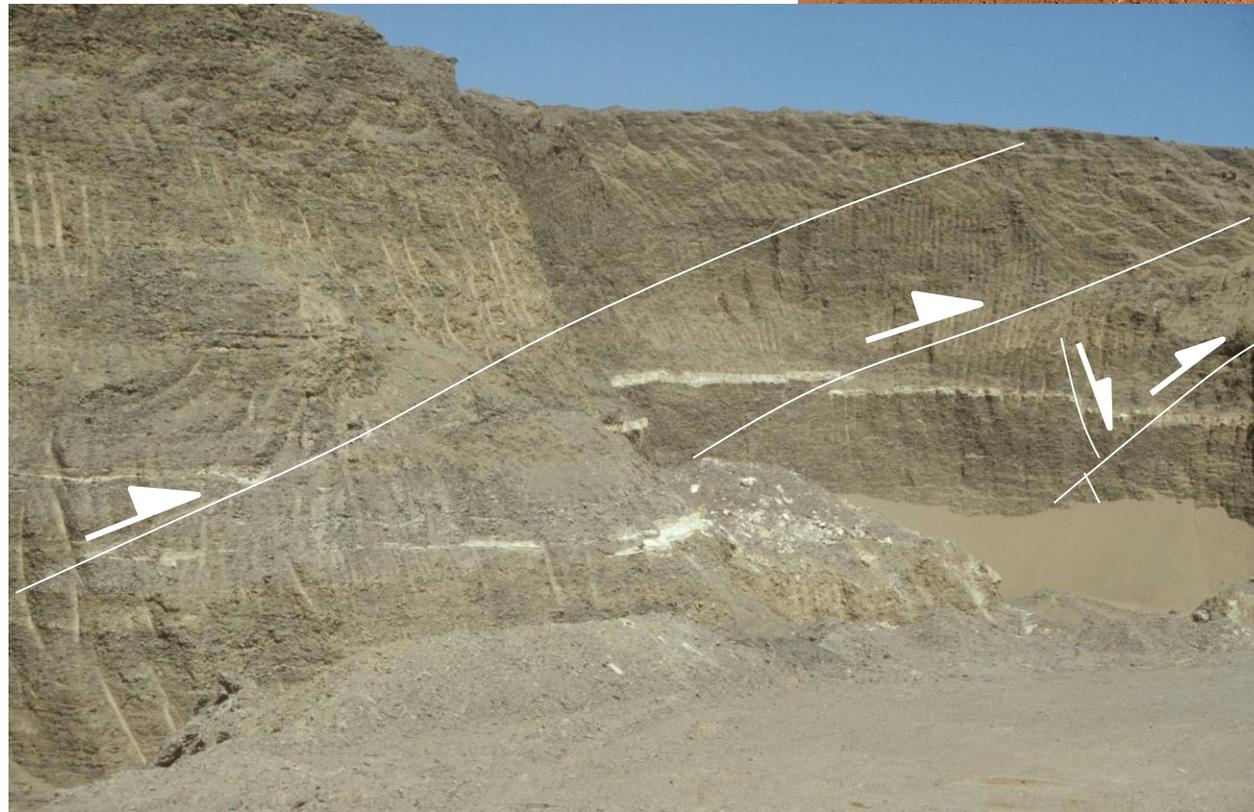
Co- and postseismic deformation of surface from Antofagasta earthquake



Location of surface ruptures does not show any co-seismic offset or deflection in fringe pattern



??Interseismic shortening at low slip rates



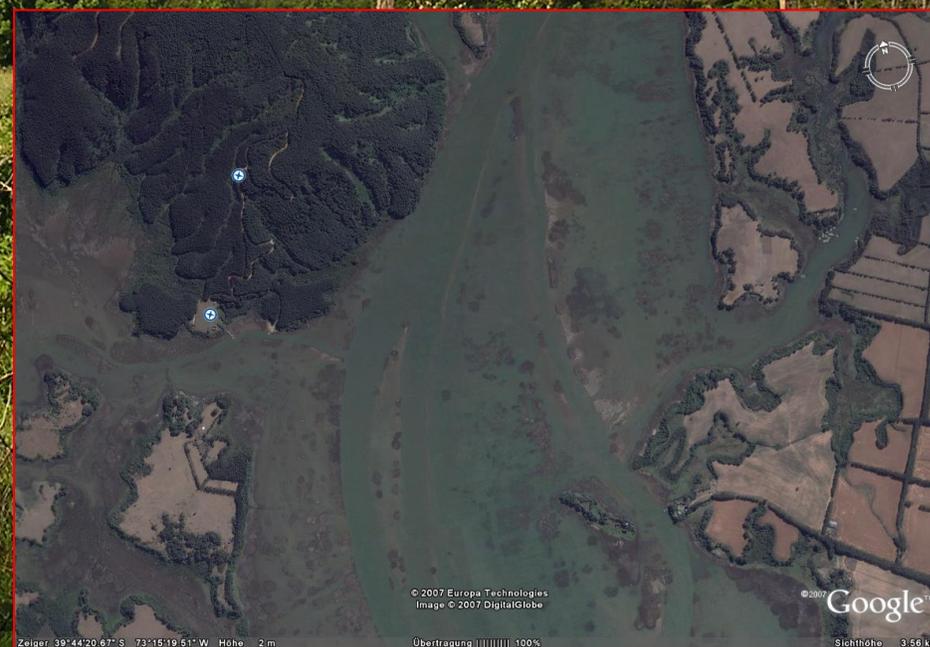
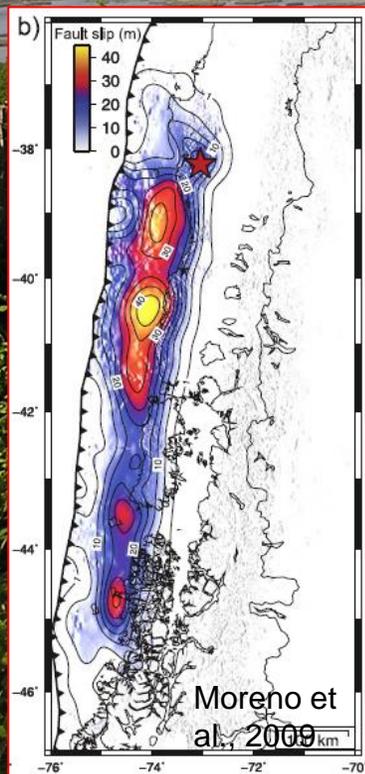
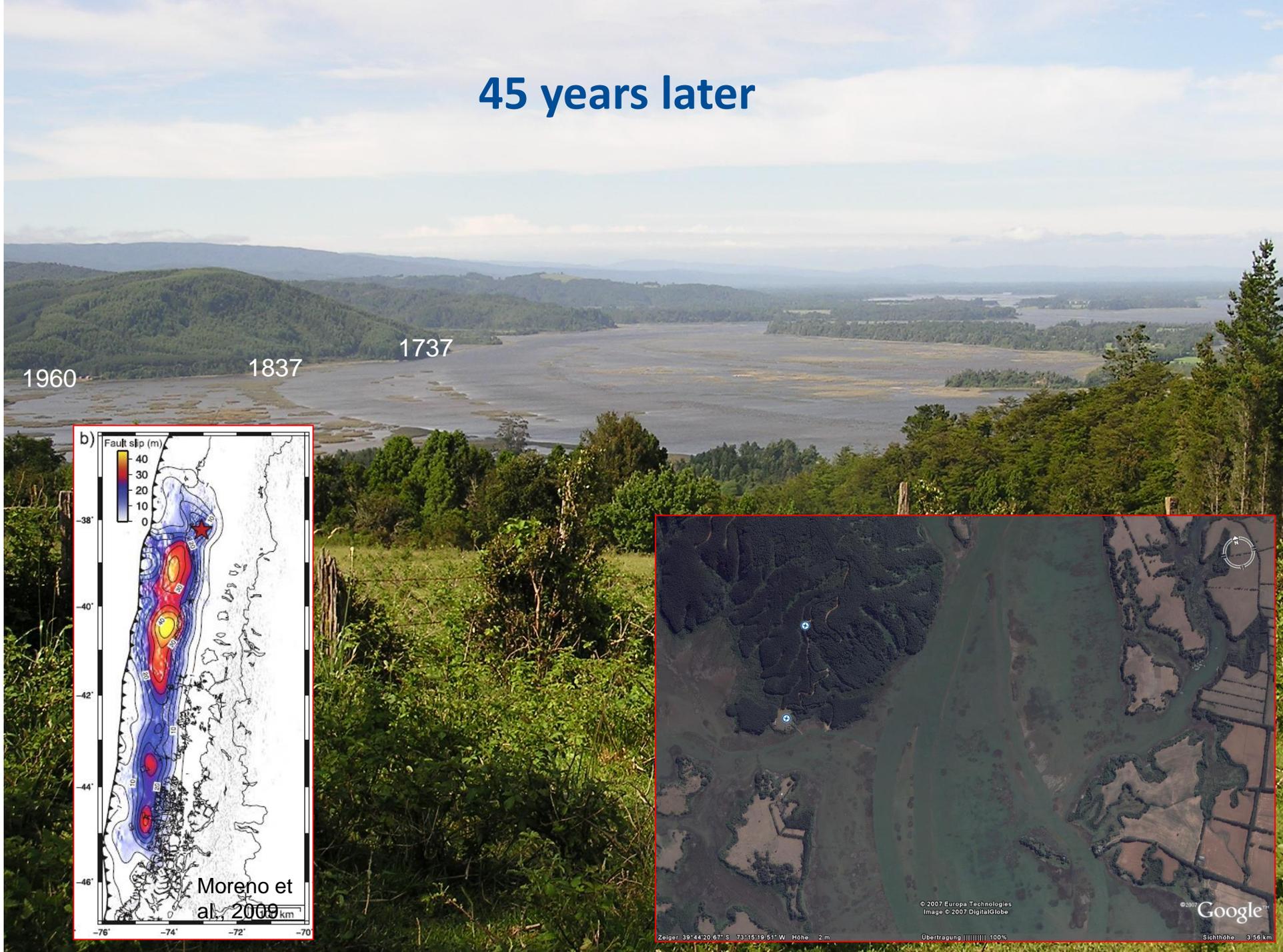


Uplifted coast

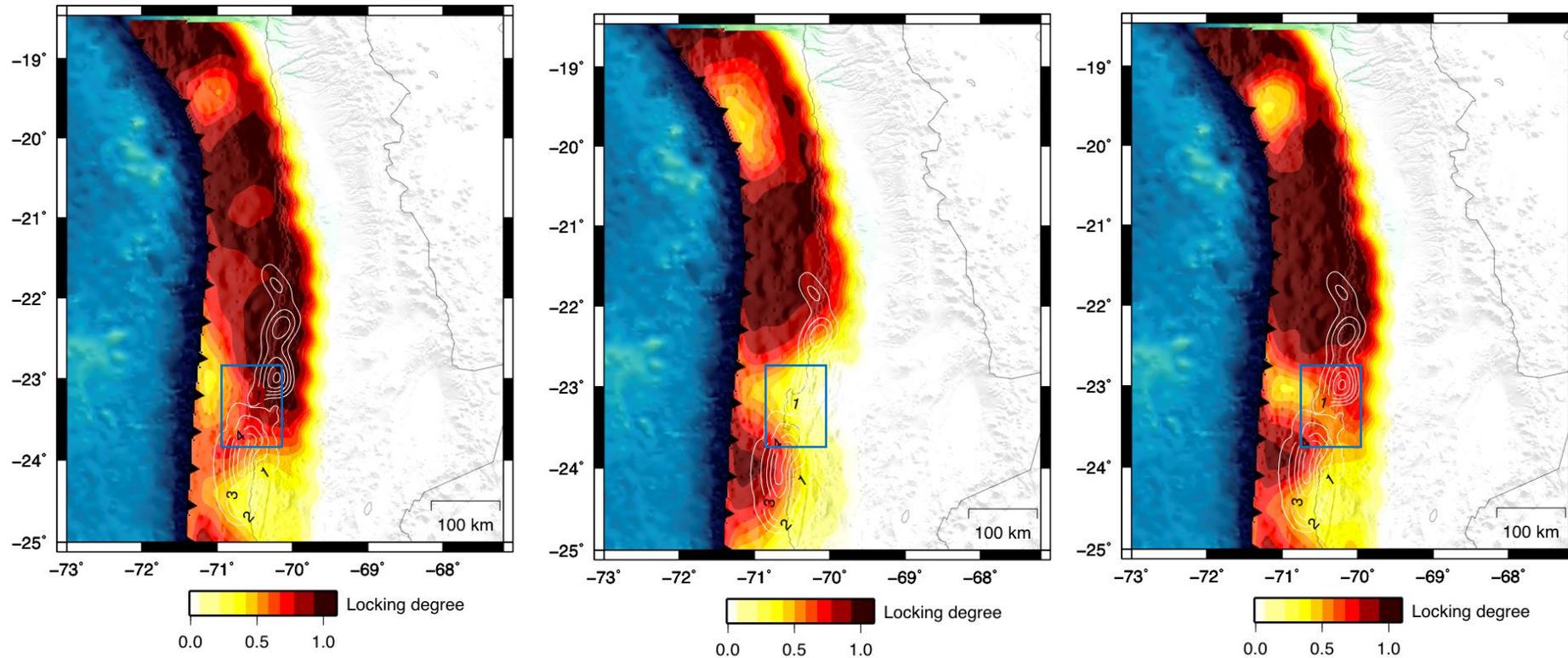
Maule earthquake, 27.2.2010



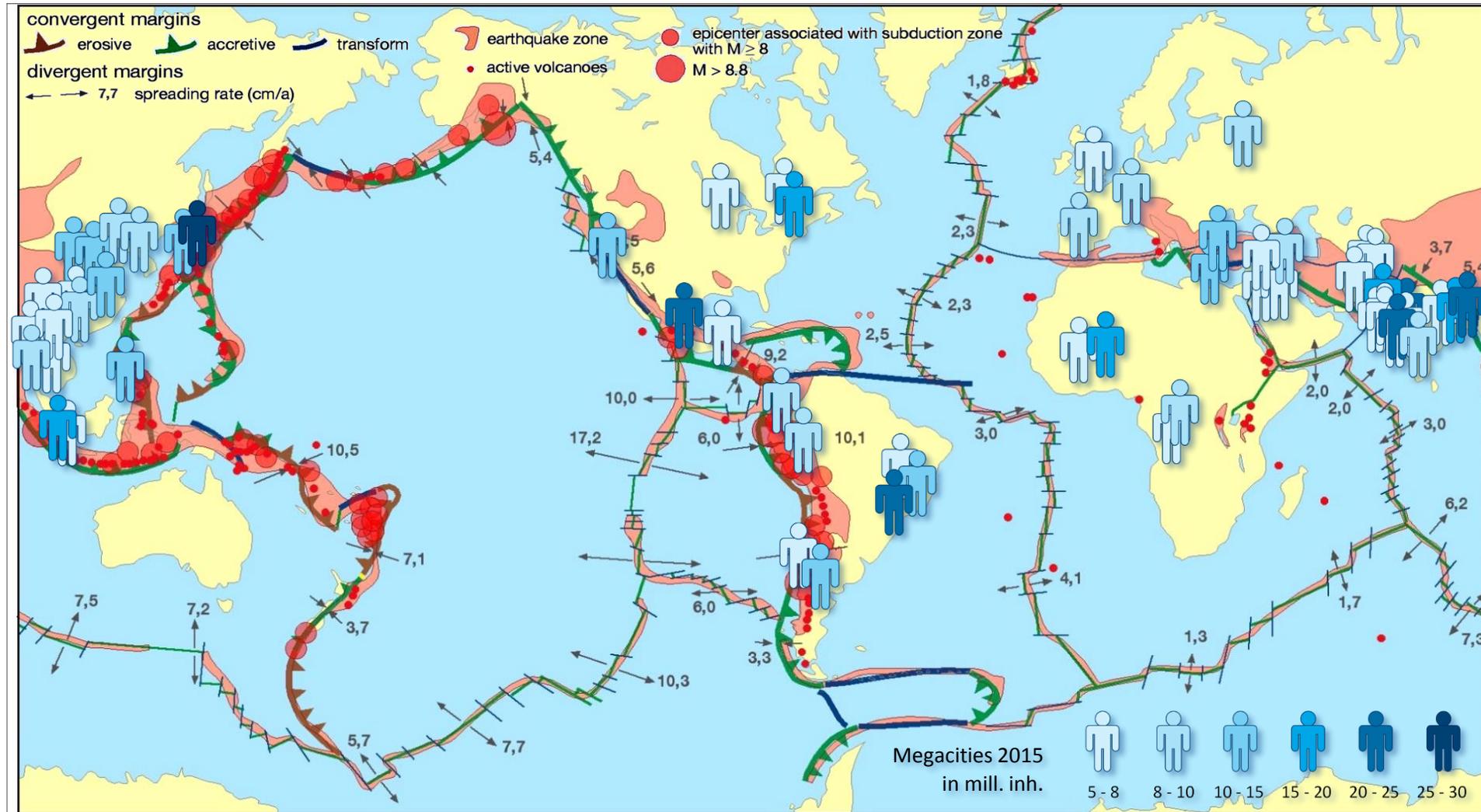
45 years later



Seismic barrier and the geological record – and the uncertainty

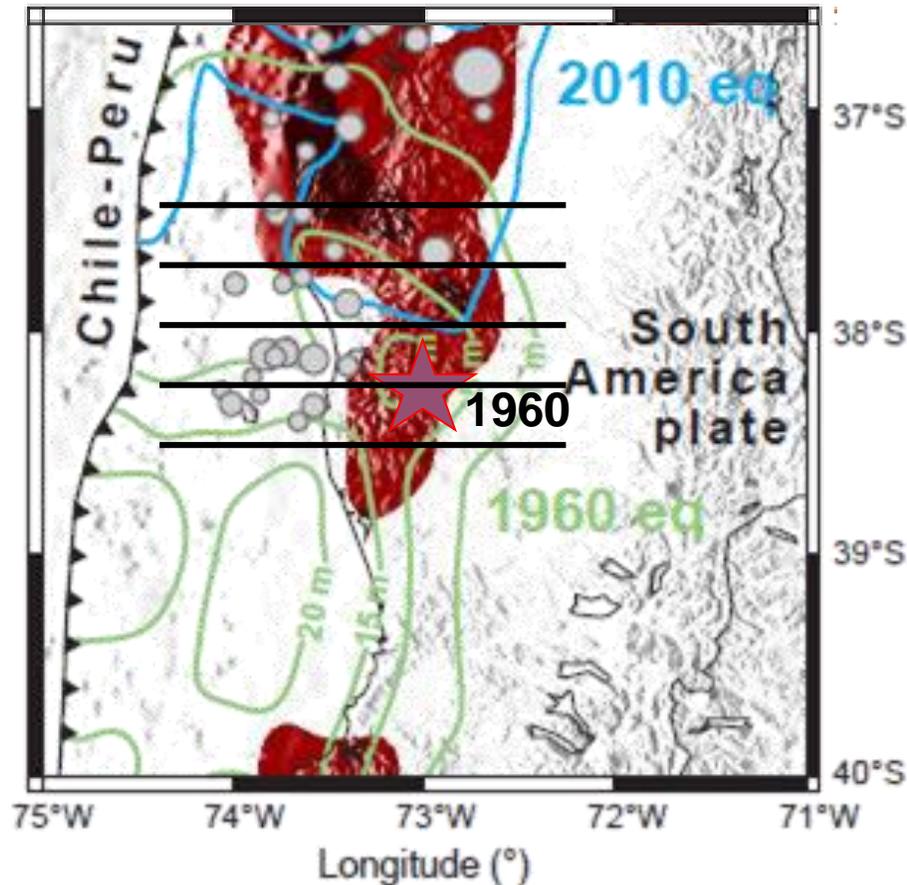


Der Erdbeben-Zoo und die Bananenschale - was wir schon immer über Subduktionszonen wissen wollten

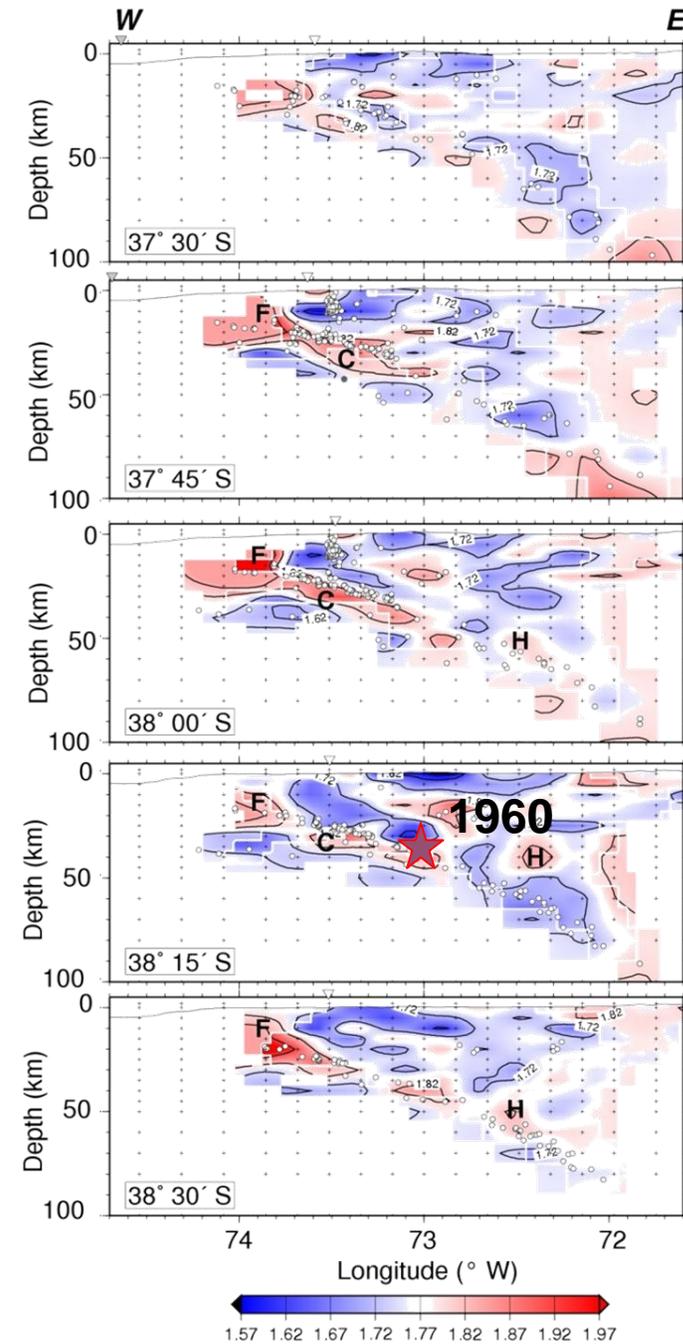


Onno Oncken, GFZ Potsdam
mit M. Moreno, J. Bedford, M. Rosenau, S. Angiboust,
B. Schurr, P. Victor, V. Mouslopoulou

Erdbebenaktivität und V_p/V_s Verhältnis korrelieren mit der Kopplung !



Moreno et al., 2010

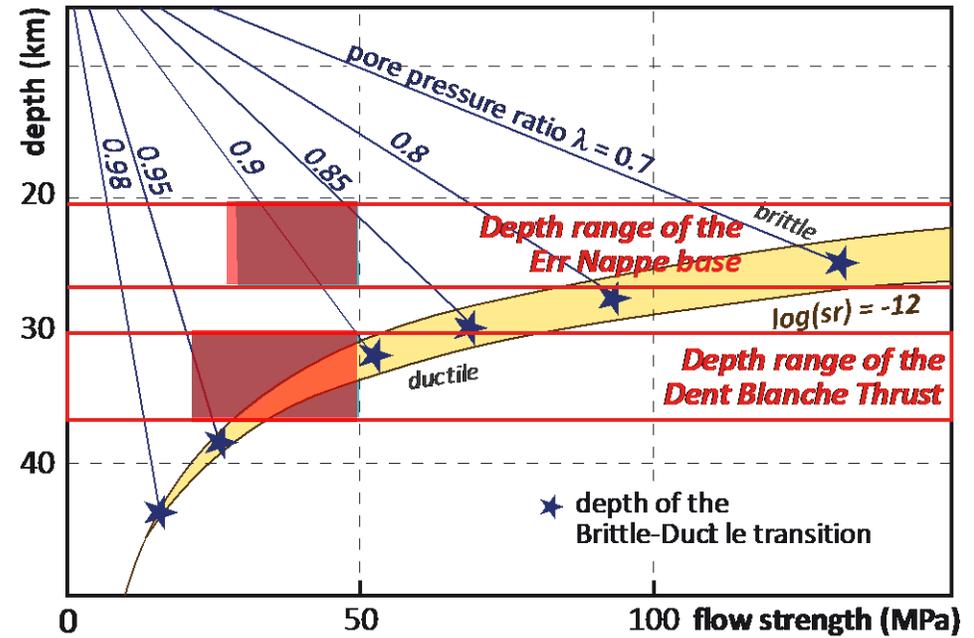
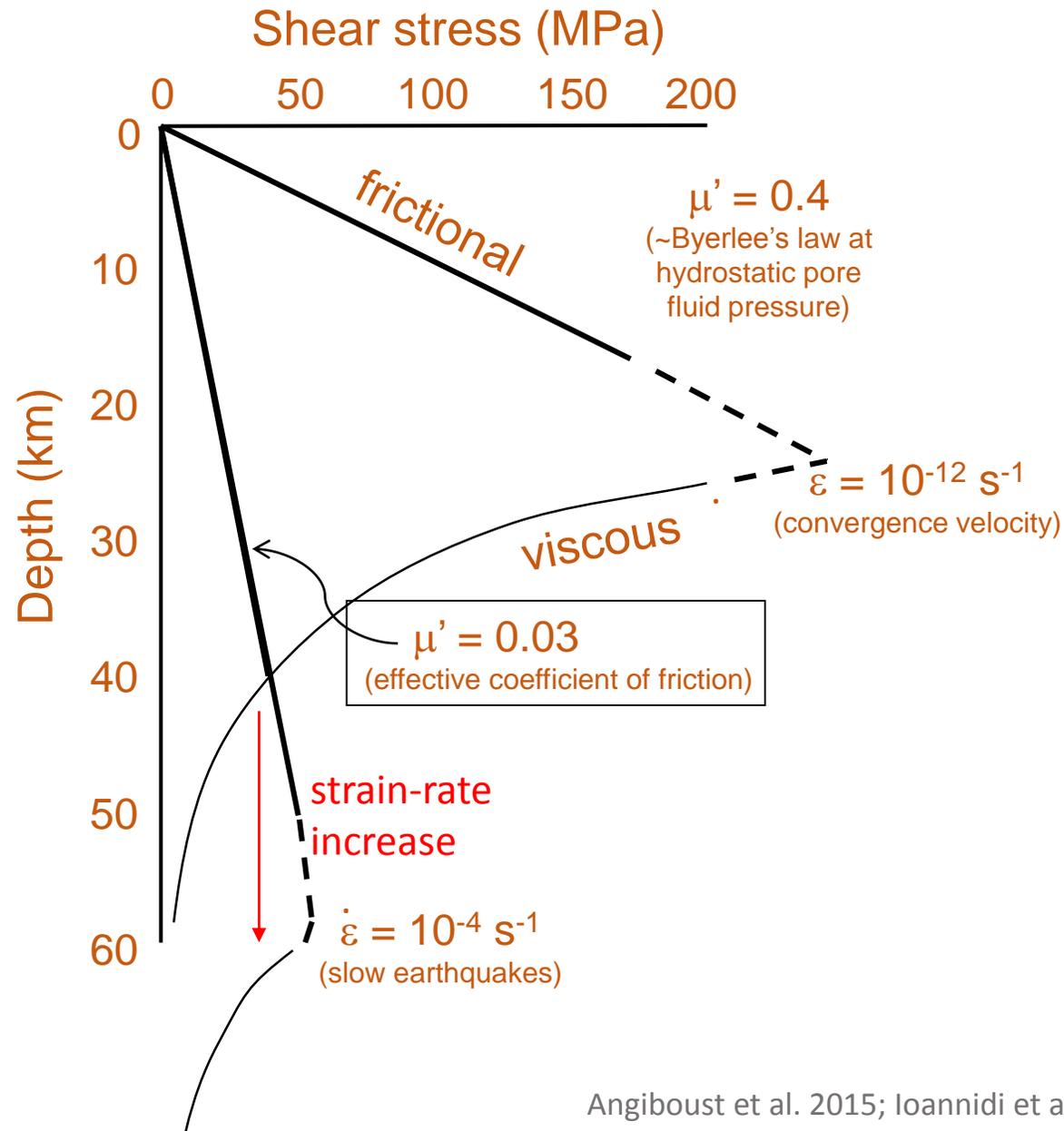


Haberland et al. 2009 v_p/v_s

2010 Maule Bruchzone

1960 Valdivia Bruchzone

Stress/fluid pressure fluctuations in transition zone



At seismic cycle time scale,
transition depth depends on

(1) Pore fluid pressure

...and on

(2) Strain rate